Can we constrain regional summer near-surface temperature projections?

A North American case study

<u>Hervé Douville</u> Météo-France/CNRM-GAME herve.douville@meteo.fr



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Response of European heat waves in CMIP5



RCP scenarios from 19 models Schoetter et al. (to be submitted)



Motivations

- Uncertainties in regional temperature change are even stronger than in global warming (e.g. Cattiaux et al. 2013);
- Changes in extreme temperature events are strongly related to seasonal mean regional temperature changes (e.g. Peings et al. 2013, Schoetter et al., submitted);
- A regional and seasonal approach might be more effective for constraining climate change (e.g. Hall and Qu 2006) given the spatial heterogeneity of (i) model biases, (ii) non-GHG radiative forcings and (iii) internal climate variability;
- Besides radiative feedbacks, changes in summer land surface temperature are also controlled by surface heat fluxes (e.g. Seneviratne et al. 2013).



A North American case study

Projected JJAS T2M anomalies 2071-2100 (RCP8.5) vs 1979-2008 (Douville et al., in preparation)

- Large inter-model spread in both the biases and anomalies of summer near-surface temperatures
- North America is a « hot spot » for land-atmosphere coupling (e.g. Koster et al. 2006)

BCC ANO=4.44





CCCMA ANO = 6.97

CNRM ANO=4.87



00 GFDL ANO=3.51



GISS ANO=3.84

IPSL ANO=6.65

D=6.65 |PSI

IPSLB ANO=5.15



INM ANO = 3.21

MIROC ANO=5.75





MPI ANO=5.43



MRI ANO=3.63









NCC ANO=5.88



2 2.5 3 3.5 4 4.5 5 5.5 6 6.5 7

Data (Focus on JJAS mean climate)

- 15 CMIP5 models: AMIP & Historical (1979-2008), RCP8.5 (2071-2100) and 1%CO2 (yrs 19-48 & 111-140) simulations
- 1 realization per model but internal climate variability is estimated using 5 members of CNRM-CM5 (Historical and RCP8.5)
- ✓ Gridded observations & reanalyses: T2M (CRU_TS3), CRF<u>SW</u> (CERES & SRB), EF=LE/(LE+H) (MTE Jung et al. 2009 & ERAI), RH2M (NCEP2) and SMI (normalized Soil Moisture Index, ERAI-Land Balsamo et al. 2013)
- ✓ 3 sets of metrics: climatology, interannual variability (correlations based on detrended seasonal anomalies), and recent trends (1979-2012 linear fits)
- Poor man's constraint: exclude the least realistic model out of the 15 independent models for each metric showing a *significant* (p-value < 5% => R>0.51) statistical link with the projected temperature anomalies



Scatterplots of climatologies: AMIP & HIST



□ Large errors in both cloud and land surface processes contribute to summer temperature biases (even more obvious in CMIP)

- Temperature biases are generally stronger in AMIP than in CMIP
- SW CRE and EF are very uncertain
- □ What is the main driver of temperature biases?

T2M biases in the CNRM AGCM



□ 4 AMIP simulations:

- ARPv5 (old physics)
- ARPv5 with improved (MODIS) surface albedo
- ARPv6 (new physics)
- ARPv6 nudged towards ERA-Interim dynamics

□ Increased precipitation intensity in ARPv6 leads to stronger runoff, lower evapotranspiration and higher JJAS temperature + dynamical amplification due to a thermal low effect

Scatterplots of anomalies: 1%CO2 & RCP8.5



Both SW CRE and EF also seem to contribute to the inter-model spread in temperature anomalies (note that models disagree about the sign of the EF response)

□ The statistical links are robust: (i) statistically significant (p-value < 5%), (ii) found in both RCP8.5 scenarios and 1%CO2 experiments and (iii) consistent with the spatial variability of T2M anomalies within most CMIP5 models

Biases vs anomalies: 1%CO2 & RCP8.5







Robust links with both cloud and land surface processes

□ Consistent between RCP8.5 & 1%CO2

One realization per model is enough

R(X,Y) vs anomalies: 1%CO2 & RCP8.5



□ Interannual correlations are robust metrics (1 realization per model is enough) but (not surprisingly) they do not necessarily represent strong constraints on projected summer temperatures

□ Yet, they are useful to detect « unusual behaviours » (e.g. GISS that show a too weak influence of SW CRE on T2M and a wrong-sign correlation between soil moisture and EF).

Trends vs anomalies: 1%CO2 & RCP8.5



Observed trend not much affected by the global warming hiatus Recent linear trends in T2M show significant correlations with projected T2M anomalies and HadGEM2ES is a possible outlier for two reasons:

(i) A strong and disputed (e.g.
Alkama et al. 2011) CO2 stomatal closure effect on plant transpiration
(ii) A strong and disputed (e.g. Zhang et al. 2013) indirect aerosol effect

Both effects are more easily detected (and falsified) at the regional scale but are probably also important at the global scale

One realization per model might however not be enough

Constraining projections by removing « outliers »



°C



□ 4 out of the 15 CMIP5 models are potential outliers

Removing outliers leads to a substantial reduction of the inter-model spread (that here includes the uncertainty due to internal variability)

Global vs regional emergent constraints?



Constraining the global warming would be extremely useful to constrain regional climate changes but seems to be very difficult

□ Constraining the regional warming might be easier and would be also very useful to constrain the global warming

NB: We only have one globe to constrain regional warming but we have multiple regions to constrain global warming

Conclusions

- Biases / uncertainties in present-day / future summer temperature over North America (as well as over Europe) are due to the poor simulation of both cloud and land surface processes;
- Constraining regional and seasonal mean temperature changes is feasible... but using multiple and physically-based metrics (rather than a single magical index);
- Assessing the models ability to simulate trends that cannot be explained by internal climate variability (D&A) will provide more and more efficient metrics.



Implications

- Running global off-line land surface simulations (e.g. GSWP) driven by observed atmospheric forcings and CO2 concentrations would be useful for understanding the role of land surface biases and feedbacks in CMIP models (cf. H2020 IMPULSE);
- Improving a sustainable (but not necessarily global) observing system is as important as improving the climate models for narrowing uncertainties in climate projections;
- Having multiple (5?) members in the forthcoming CMIP-DECK experiments (at least for the historical simulations) would be useful to distinguish between model uncertainty and internal variability and for the development of robust emergent constraints.



Prospects: Looking at two « canonical » regions?

BCC ANO=4.13CCCMA ANO=5.74 CNRM ANO=4.67

GFDL ANO=2.91





GISS ANO=3.56



INM ANO=2.56

IPSL ANO=5.91

IPSLB ANO=4.53









CSIRO ANO=3.77

MOHC ANO=7.16







NCC ANO=4.99



MRI ANO=3.40



NCAR ANO=4.50



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Climatol. vs anomalies: 1%CO2 & RCP8.5



R(X,Y) vs anomalies: 1%CO2 & RCP8.5



Outliers: MIROC (and GISS?) models



Trends vs anomalies: 1%CO2 & RCP8.5



Outlier: CCCMA (and IPSL?) models



Prospects: Looking at other regions?





T2M (left) & CRFSW (right) biases: HIST



T2M (left) & CRFSW (right) biases: AMIP



Trends vs anomalies: 1%CO2 & RCP8.5



Recent linear trends in T2M show significant correlations with projected T2M anomalies (especially in RCP8.5) and HadGEM2ES (with a strong indirect aerosol effect and a strong CO2 physiological effect on surface evapotranspiration) appears as an outlier

□ Also true for CRFSW and EF in 1%CO2 experiments (less clear in RCP8.5 due to model-dependent radiative effects of anthropogenic aerosols) but large uncertainties in both « observed » and simulated trends (one realization per model is not necessarily enough)

Scatterplots of anomalies: 1%CO2 & RCP8.5



Both SW CRE and EF contribute to the inter-model spread of T2m anomalies 24

□ The statistical links are robust (i.e. statistically significant and found in both RCP8.5 and 1%CO2 expts)

□ Also true for scaled anomalies



Effect of scaling on the intermodel spread



JJAS warming over Europe in CMIP5



Inter-model spread in European warming as a function of...



RCP8.5 scenarios from 33 models Cattiaux et al. 2013



Seasonal dynamics of T2M biases



□ The MME warm bias appears in June, in line with a lack of SW cloud radiative cooling

□ It is then amplified by an underestimation of the evaporative fraction

