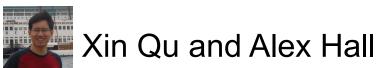
# The strength of the tropical inversion and its response to climate change in 18 CMIP5 models





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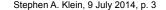




# Outline



- > Tropical inversion: Its importance
- How well do CMIP5 models simulate the present-day tropical inversion strength?
- Why do models predict the tropical inversion strength to increase in a warmer world?
- > Take away points



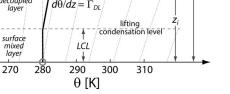
# **Tropical inversion: Its importance**

- The ubiquitous tropical trade inversion and its strength are key factors for subtropical boundary layer clouds
  - The inversion limits the height of boundary layer and its strength limits the entrainment rate of free tropospheric air
  - On seasonal-to-interannual time-scales, low cloud amount is highly correlated with the strength of the inversion
  - Modeling studies suggest the low-cloud response to climate change is sensitive to changes in inversion strength

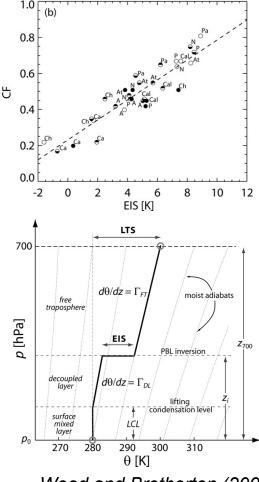
How well do climate models simulate the tropical inversion strength?

What do climate models predict for changes in inversion strength? And why?

We use the Estimated Inversion Strength:  $EIS = LTS - \Gamma_m^{850} \cdot (z700 - LCL)$ 



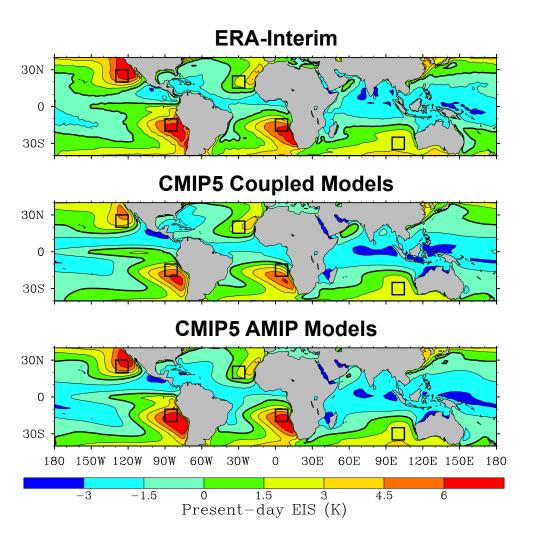
Wood and Bretherton (2006)





# Present-day tropical inversion strength



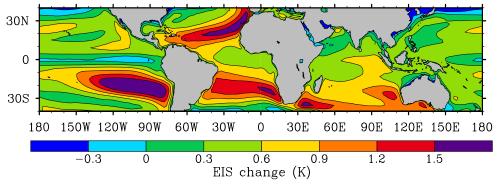


- Very good large-scale geographic correspondence
- Coupled models have a low bias (-1.6 ± 1.2 K) to the inversion strength in stratocumulus (Scu) regions
  - Low bias is mainly attributable to SST warm biases in Scu regions
- Models with prescribed SST have a much smaller low bias (-0.5 ± 0.6K)
- EIS underestimate likely contributes to low-cloud amount underestimates in Scu regions in CMIP5 models

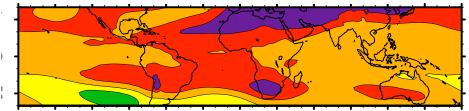
# Predicted 21<sup>st</sup> century EIS changes



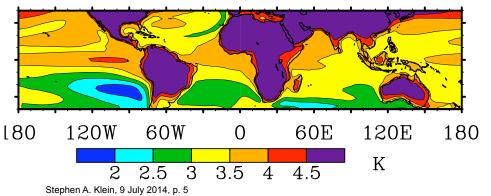
#### ΔEIS in CMIP5 Models (RCP8.5)



ΔT<sub>700</sub> in CMIP5 Models



 $1.2 \Delta T_s$  in CMIP5 Models



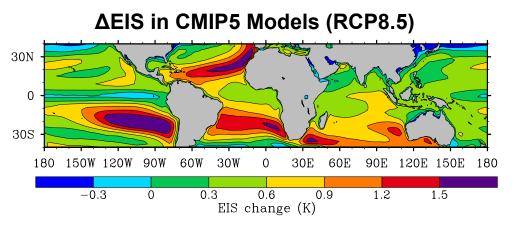
- EIS <u>robustly</u> increases in the 21<sup>st</sup> century simulations particularly in Scu regions and the Southern Hemisphere
  - In Scu regions  $\Delta$ EIS is +1.0 ± 0.2 K
  - This is contrary to simplified models of the tropics

$$\Delta EIS \approx \Delta T_{700} - 1.2 \Delta T_s$$

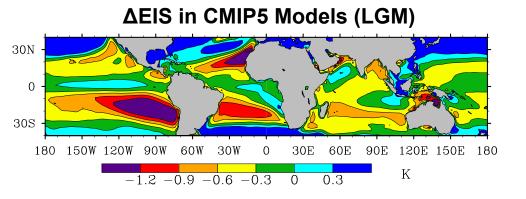
- $\succ \Delta T_{700} (Scu regions) \sim \Delta T_{700}$  $(land) > \Delta T_{700} (warm pool)$
- ΔSSTs (Scu regions) < ΔSST</li>
  (warm pool) < ΔT-surf (land)</li>

# 21st century vs. Last Glacial Maximum





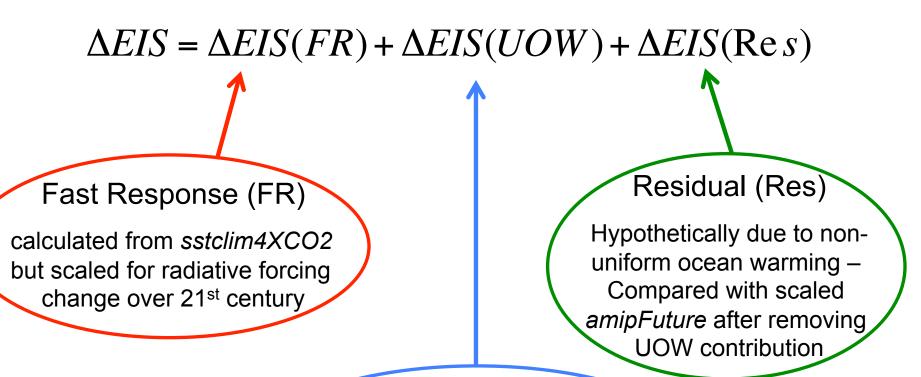
- EIS <u>robustly</u> increases in the 21<sup>st</sup> century simulations particularly in Scu regions and the Southern Hemisphere
  - In Scu regions  $\Delta$ EIS is +1.0 ± 0.2 K



Predicted EIS changes (in Scu regions) are symmetric with respect to the sign of the climate change



### What drives EIS increases?

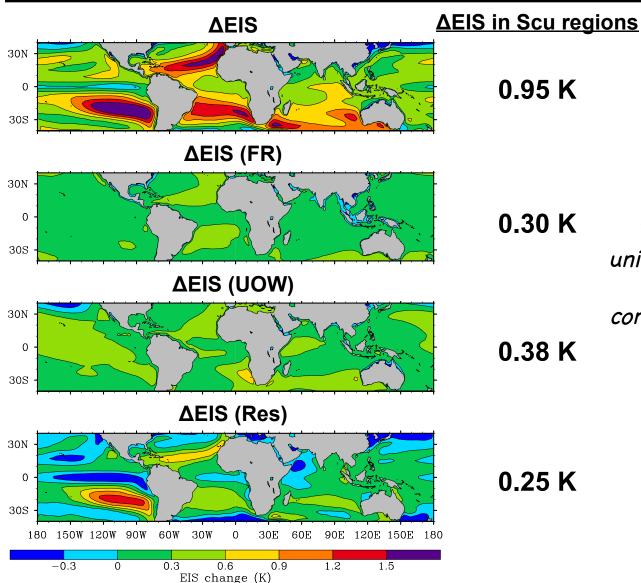


#### Uniform Ocean Warming (UOW)

calculated from *amip4K* scaled for the tropical warming over 21<sup>st</sup> century

Note the power of CFMIP/CMIP  $\rightarrow$  this study utilizes 11 experiments

### The contributions to EIS changes



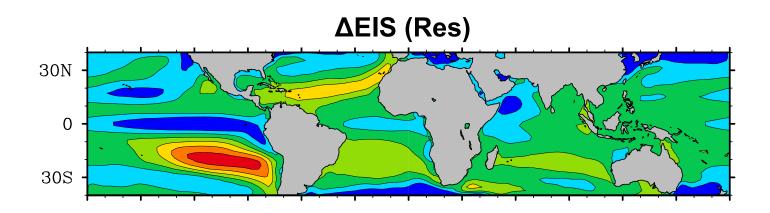
0.95 K

0.30 K The fast response, uniform ocean warming, and residual each contribute about equally

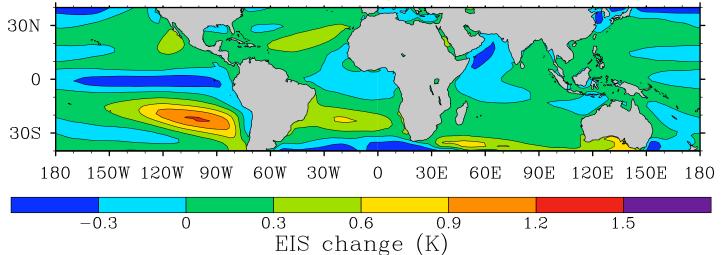
0.38 K

0.25 K



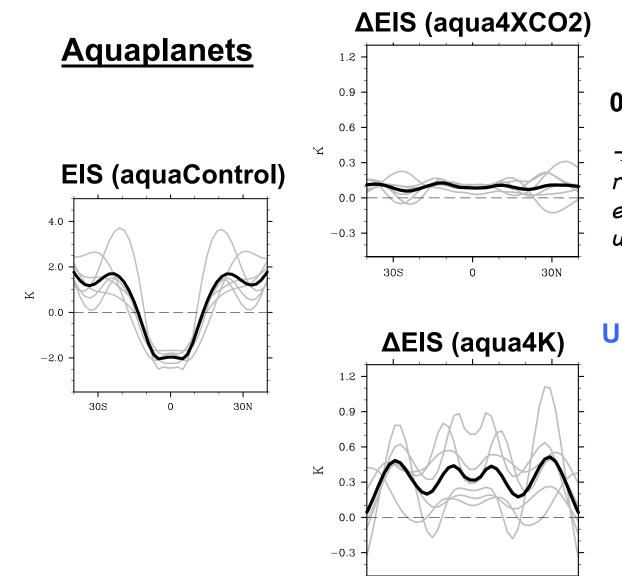


**ΔEIS from amipFuture (with UOW removed)** 



### What is the role of land?





30S

0

30N

#### **Fast Response**

#### 0.1 K << <u>∆</u>EIS (real)

→ Land plays a significant role in the fast response, as expected since land warms up but ocean does not

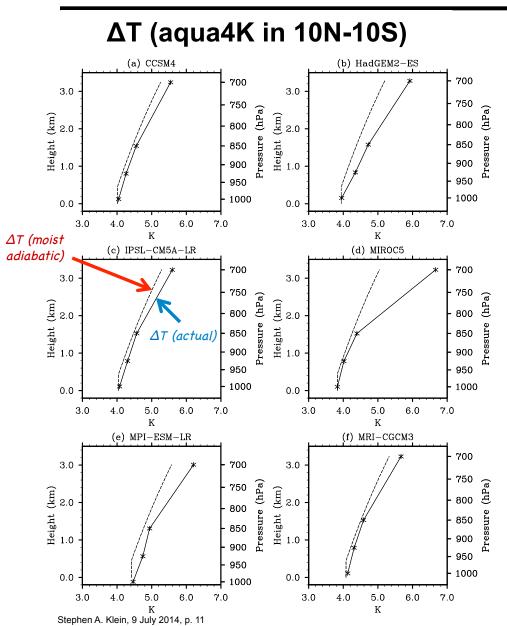
**Uniform Ocean Warming** 

#### **0.3 K** ~ $\Delta$ EIS (real)

→ Land does not play a significant role in the uniform ocean warming

 $\rightarrow$  So what does?

# Super - Moist-Adiabatic Warming



→ Warming at 700 hPa in the deep tropics exceeds that predict from a moist adiabat (assuming constant RH)

#### Why?

→ This drives the EIS increase in the aquaplanets and the uniform ocean warming simulations (we believe)



- EIS is relatively well simulated in climate models apart from a noticeable underestimate (~1.6 K in Scu regions) resulting from coastal SST warm biases
- Climate models simulate EIS increases in a warmer world and decreases in a colder world
- > EIS increases in a warmer world would encourage greater low-level cloudiness (a negative feedback)
  - This increase is expected to be smaller than the decrease directly resulting from the warming (Bretherton et al. 2013, Bretherton and Blossey 2014, Qu et al. 2014)
  - $\Delta EIS \sim 40\% \Delta SST \& d(LCC)/dEIS \sim -1 X d(LCC)/dSST$



- EIS increases are a combination of the results of the direct response to CO<sub>2</sub>, greater warming over land, a smaller than average SST rises in Scu regions, and a general warming of the planet
- > Do we believe these climate model predictions for EIS?

Response	Believable?	Physical Explanation
Direct response to CO2	~~	Less lower tropospheric radiative cooling
Greater warming over land	•	Lower heat capacity over land leads to greater warming of air aloft and which influences $T_{700}$ in nearby marine Scu regions
Increased SST gradient between Scu region and warm pool	?	Greater evaporative damping for a uniform radiative forcing
Super-moist adiabatic warming in deep convection regions	??	Name a reason – please!

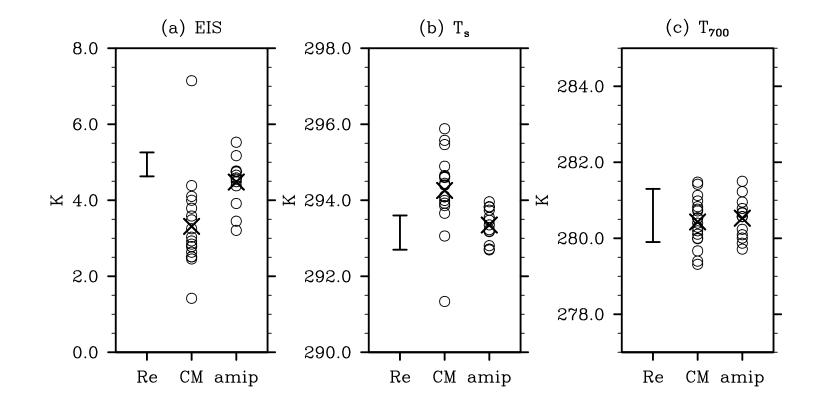
# Thank you for your attention!



### **Extra Slides**

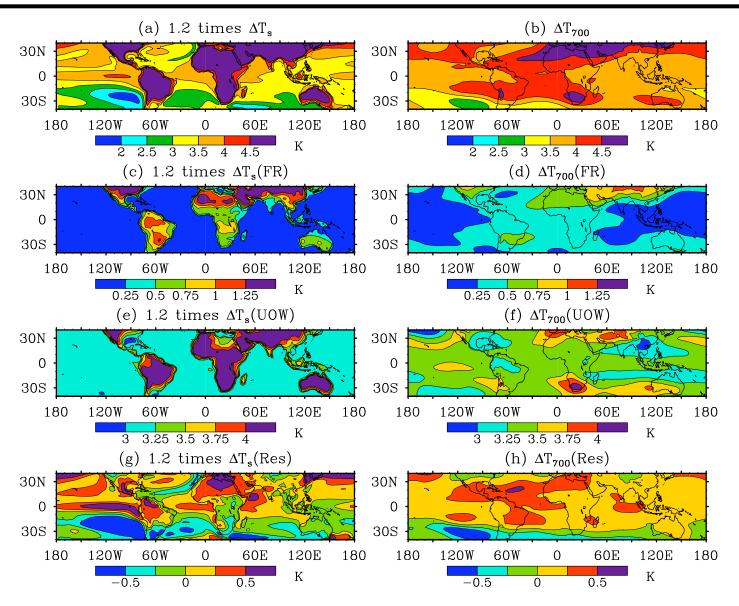
Stephen A. Klein, 9 July 2014, p. 15





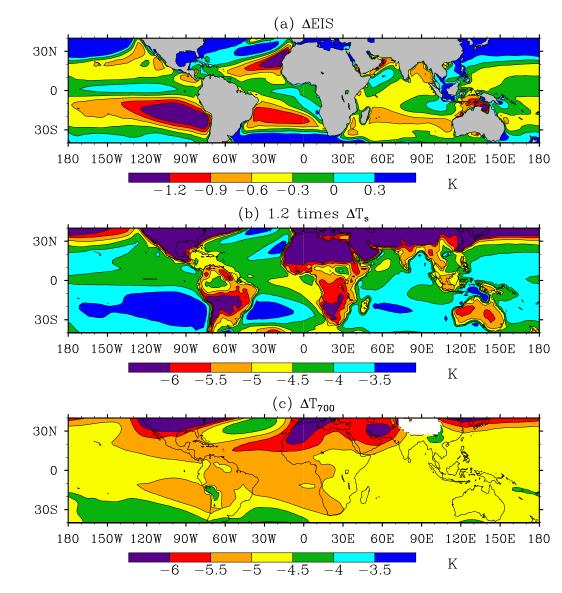
# Various contributions to Ts and T700 changes





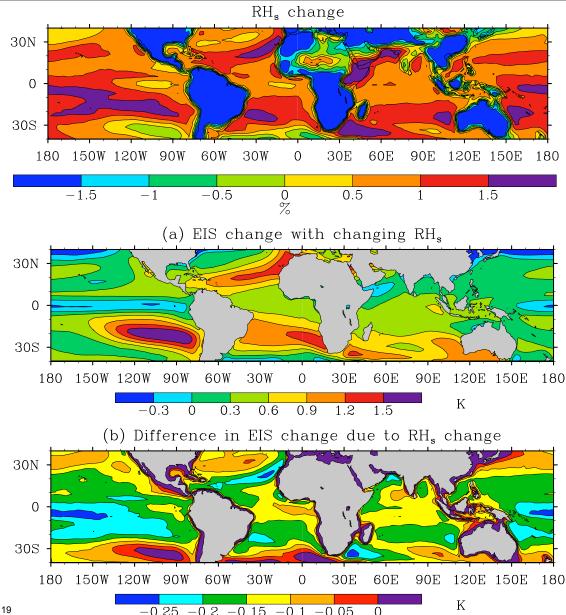
# **EIS changes in LGM simulations**





Stephen A. Klein, 9 July 2014, p. 18

### The effect of changing ocean surface RH



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$$\begin{split} EIS &= LTS - \Gamma_m^{850} \cdot (z700 - LCL) \\ \Delta EIS &= \Delta EIS(FR) + \Delta EIS(UOM) + \Delta EIS(\text{Re}\,s) \\ \Delta EIS &\approx \Delta T_{700} - 1.2\Delta T_s \\ \Delta T_x &= \Delta T_x(FR) + \Delta T_x(UOM) + \Delta T_x(\text{Re}\,s) \end{split}$$