

# Forcing, Feedbacks & Climate Sensitivity of the CMIP5 models

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### Motivation: Climate sensitivity and CMIP5

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MULTI-MODEL AVERAGES AND ASSESSED RANGES FOR SURFACE WARMING



In CMIP5, we are also likely to have a range of projections, so we will need to explain why different models respond differently to the same external forcing



- Metrics for measuring a model's response to external forcing
- Methods that can be used to evaluate these metrics with the CMIP5 models
- Some early results from CMIP5 (3 models!)
- Limitations, discussion and summary



### Climate change metrics in AOGCMS

- There are many metrics that can be used to quantify and compare a model's response to external forcing, how might they have changed since CMIP3?
- <u>Transient climate response</u> (TCR):  $\Delta T$  about yr 70 after 1% CO<sub>2</sub> increase. It is a more 'realistic' metric and can be readily computed, it can also provide some information about transient heat uptake & feedbacks (e.g. Gregory and Forster, 2008)
- <u>Eqm climate sensitivity</u>  $(\Delta T_{2x})$ : eqm  $\Delta T$  after  $2xCO_2$ . For CMIP5, this is too computationally expensive for AOGCMs, but large step forcing experiments are still a very useful 'science tool' for evaluating and comparing forcing and feedback processes
- This talk will focus on how we can use abrupt  $4xCO_2$  experiments in CMIP5 to diagnose and compare model forcing and feedback processes, as well as make a prediction of each models  $\Delta T_{2x}$

CMIP5 data					
Met Office Hadley Centre		Coupled run that is not part of CFMIP2 experiments			
		Abrupt4xCO2	sstClim4xCO2	CMIP3 $\Delta T_{2x}$	
	INM-CM4	150yr	30yr	2.1 K	Spans CMIP3
	CNRM-CM5	150yr	n/a	n/a	<ul> <li>climate</li> <li>sensitivity</li> </ul>
	HadGEM2-ES	270yr (in house)	30yr (in house)	4.4 K 🤳	range

- Also includes corresponding pre-industrial fully coupled run and preindustrial sst-climatology
- Not much of a multi-model intercomparison yet, but fortuitously these models represented the low and top end of the CMIP3 generation



## How do we quantify model response?

Following, Gregory et al., (2004) and Gregory and Webb (2008), the energy balance of the climate system can be expressed by:



 $N = F - Y \Delta T$ 



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## CMIP5 piControl & abrupt4xCO2

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#### Prediction of $\Delta T_{2x}$ has a range ~ 2 to 4.6 K, very similar range to CMIP3

• Which feedback processes give rise to ~ -0.63 to -1.50 Wm<sup>-2</sup> K<sup>-1</sup> range?



#### Differences in clear-sky feedbacks not enough to explain sensitivity range



- As defined by CRE, cloud feedback is: ~ +0.1 Wm<sup>-2</sup>K<sup>-1</sup> HadGEM2-ES ~ -0.35 Wm<sup>-2</sup>K<sup>-1</sup> INM-CM4.0
- Largest differences occur in the Pacific basin, particularly in the NH

Differences in CRE feedback is the largest contributor to sensitivity range



## Discussion: i) is it really linear?

 In AOGCMs, separating forcing and feedback can be complicated, as shown below for abrupt 4xCO<sub>2</sub> in HadCM3





## Discussion: ii) how many different timescale responses are there?





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- Climate sensitivity helps us understand the causes of uncertainty in climate model projections.
- We can now look at this in some CMIP5 models as data becomes available.
- Preliminary results show that the range of eqm climate sensitivity has not reduced from the previous generation of models.
- Differences in cloud feedback, once again, appear to be the largest single cause of this uncertainty.
- There are limitations on the methods used, such as linearity. As more models become available this will indicate the extent of the problem.
- Finally, are we too  $CO_2$  focused when performing sensitivity experiments?





## Additional slides



HadGEM2-ES - TOA LW clear-sky (Wm<sup>-2</sup>)







 In HadGEM2-ES, there is also a "cloud adjustment" that comes about due to plant-CO<sub>2</sub> physiological effects, reducing transpiration and hence drying and warming the boundary layer (Doutriaux-Boucher et al., 2009)