

Rapid cloud adjustments, cloud feedbacks and large-scale forcings

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Outline

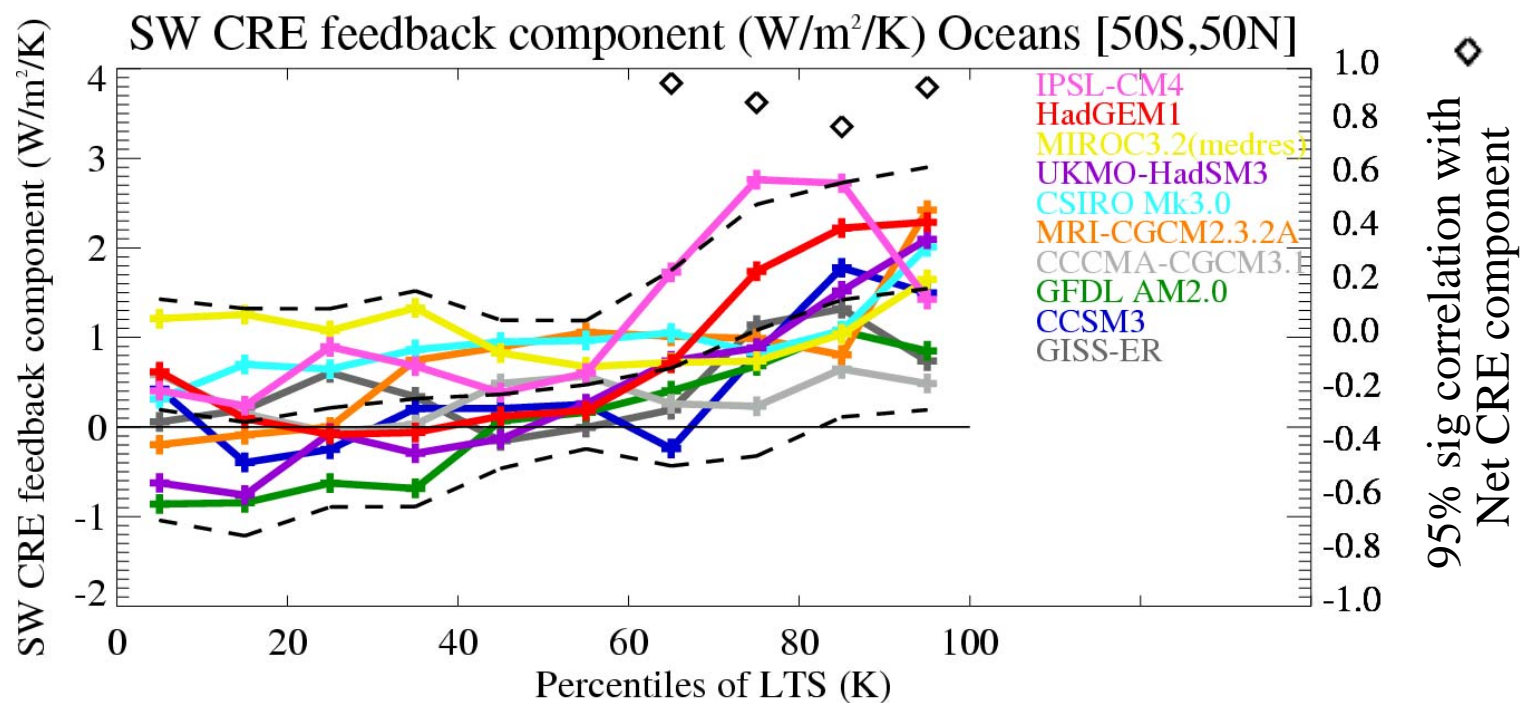
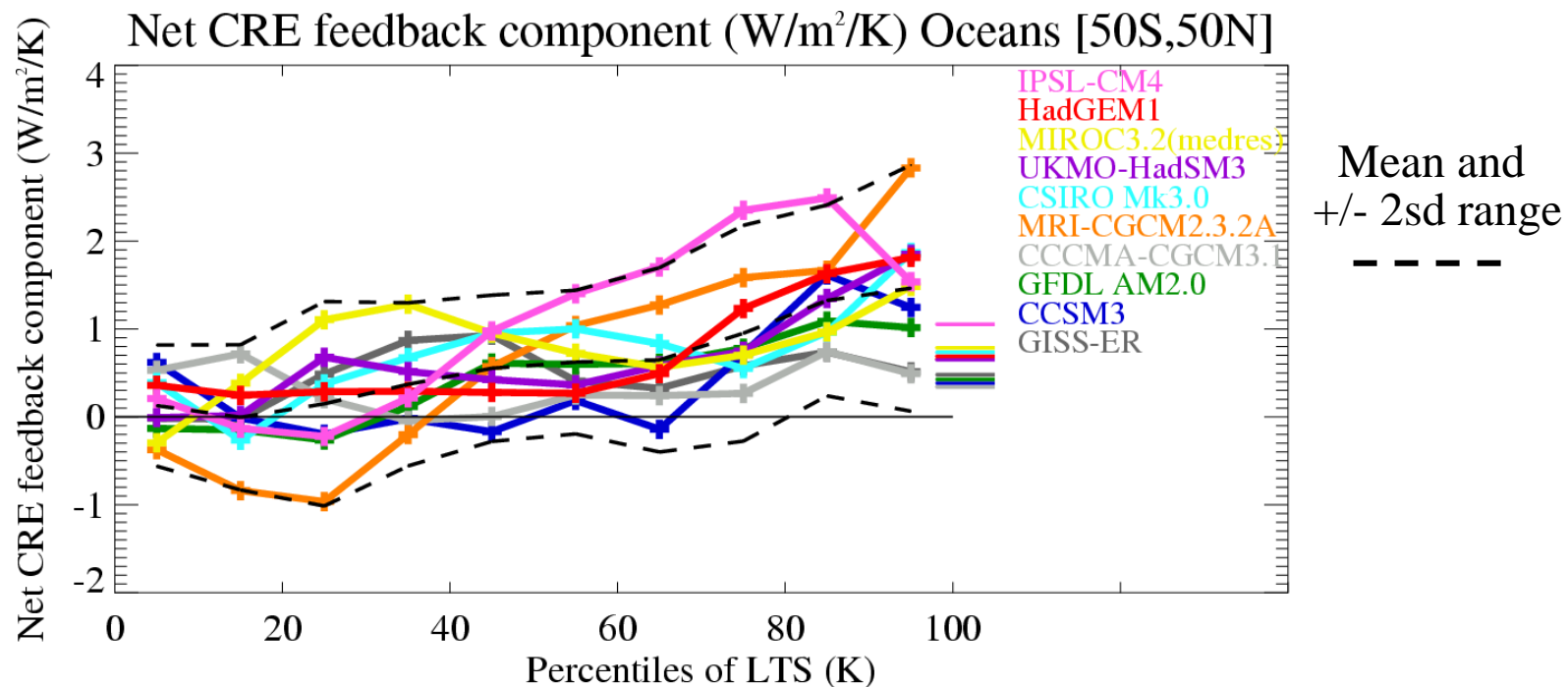
Gregory and Webb (2008) found evidence of rapid cloud adjustments in direct response to CO₂ doubling in the AR4/CFMIP-1 slab models

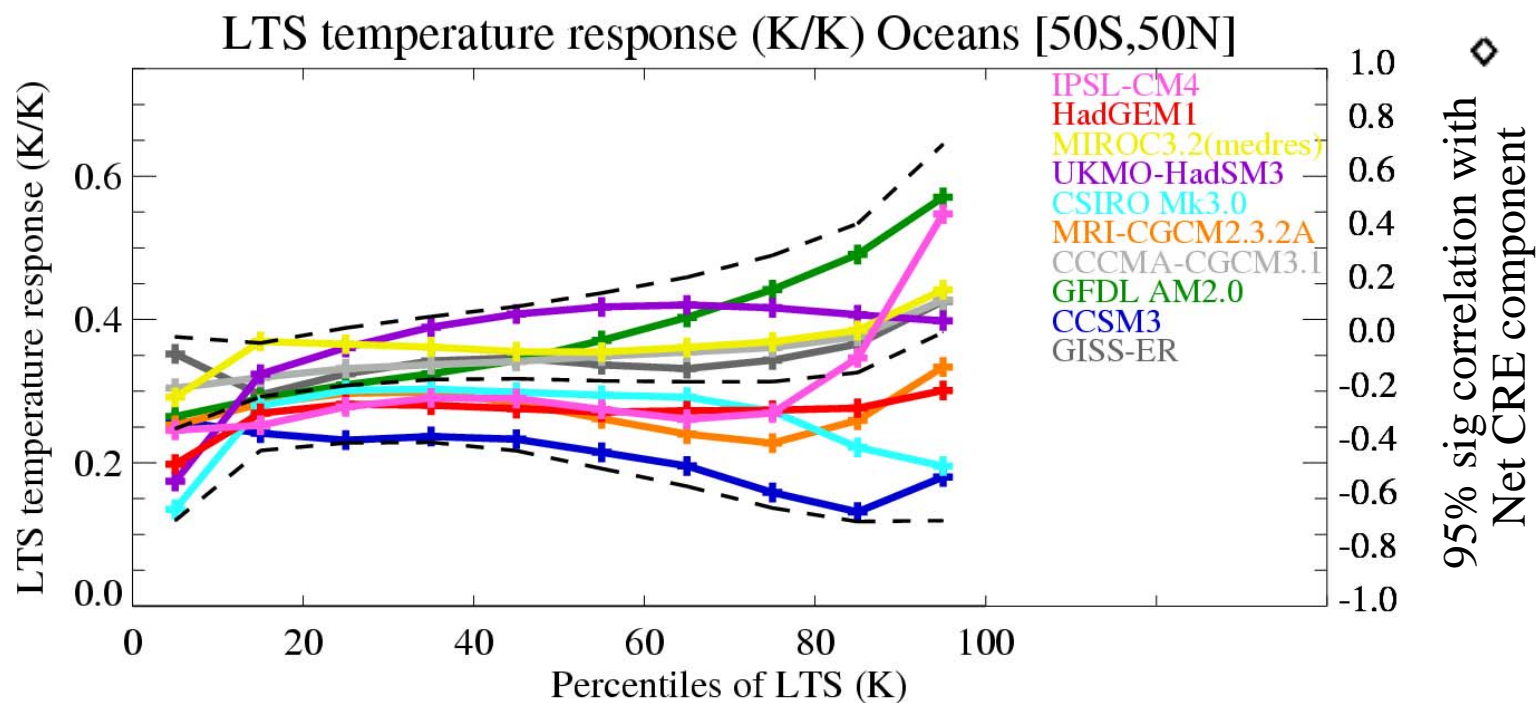
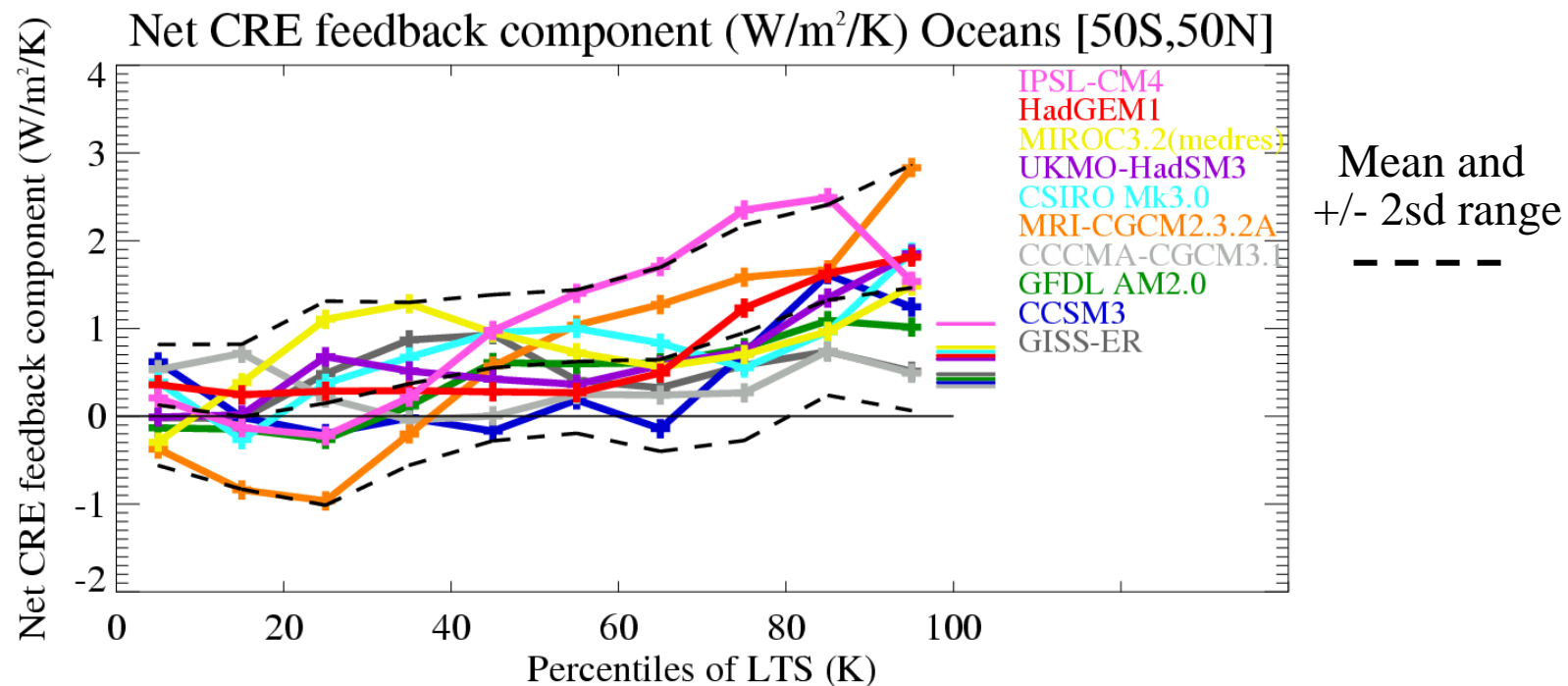
Here we diagnose forcings and feedbacks using the Gregory method in an enlarged ensemble with four additional AR4 slab models

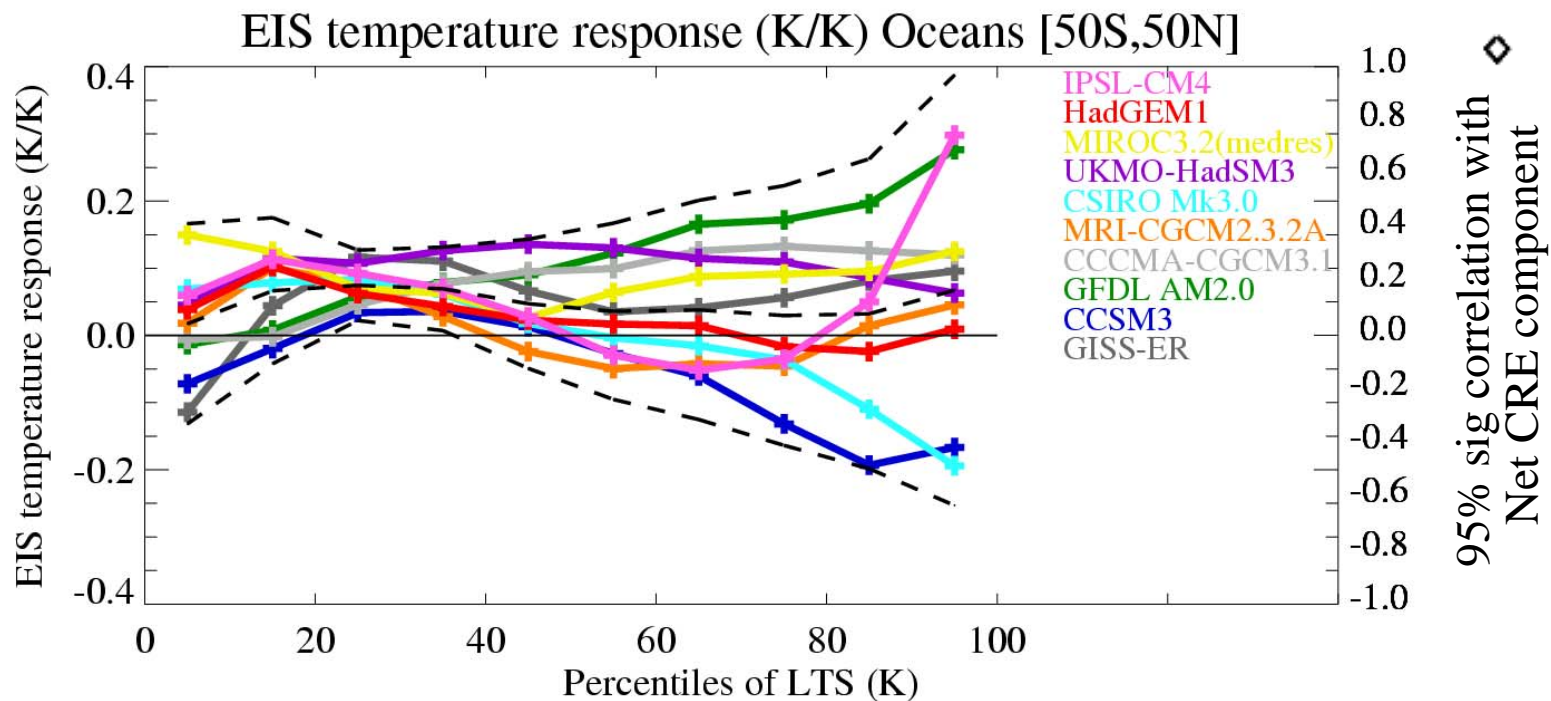
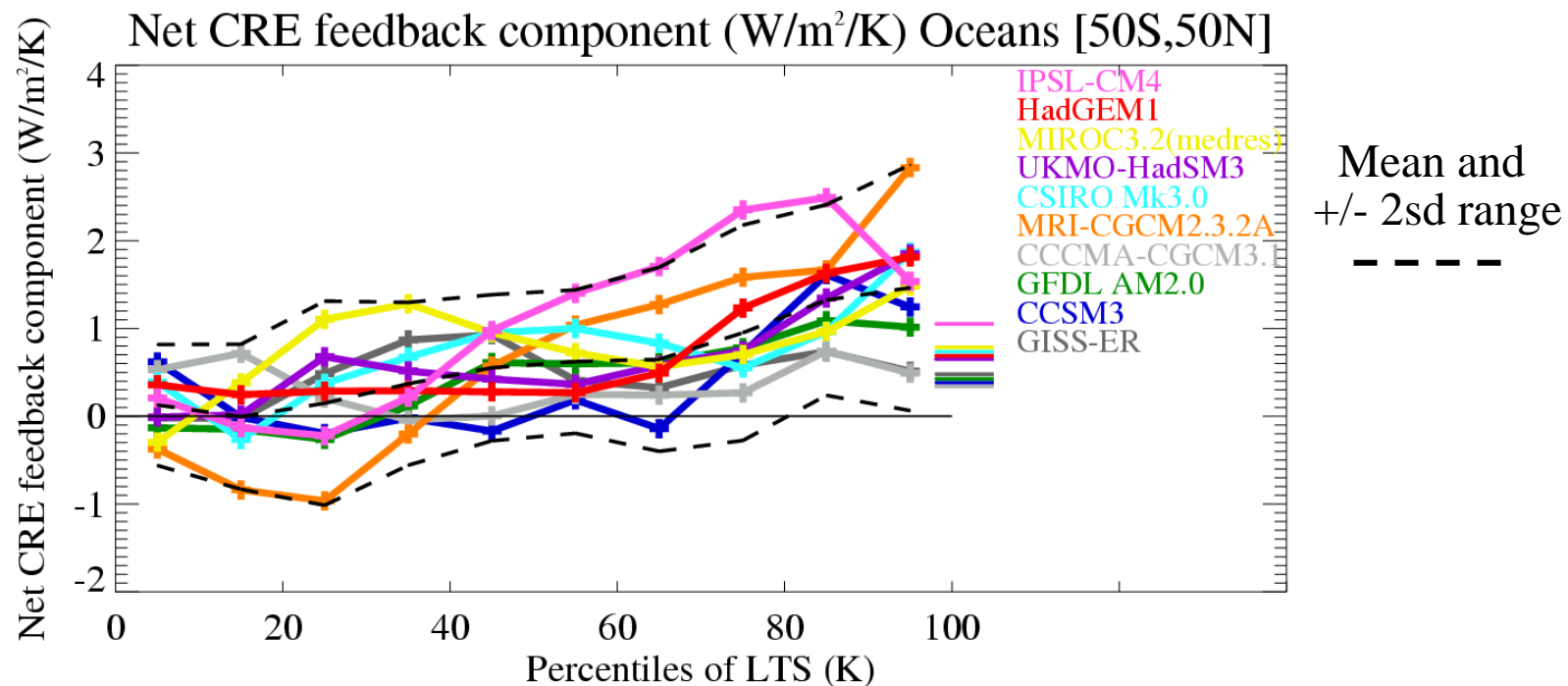
We separate shallow vs deep cloud regimes by sorting cloud components of forcing and feedback into percentiles of lower tropospheric stability (LTS) over ocean points using the method of Wyant, Bretherton and Blossey (2009)

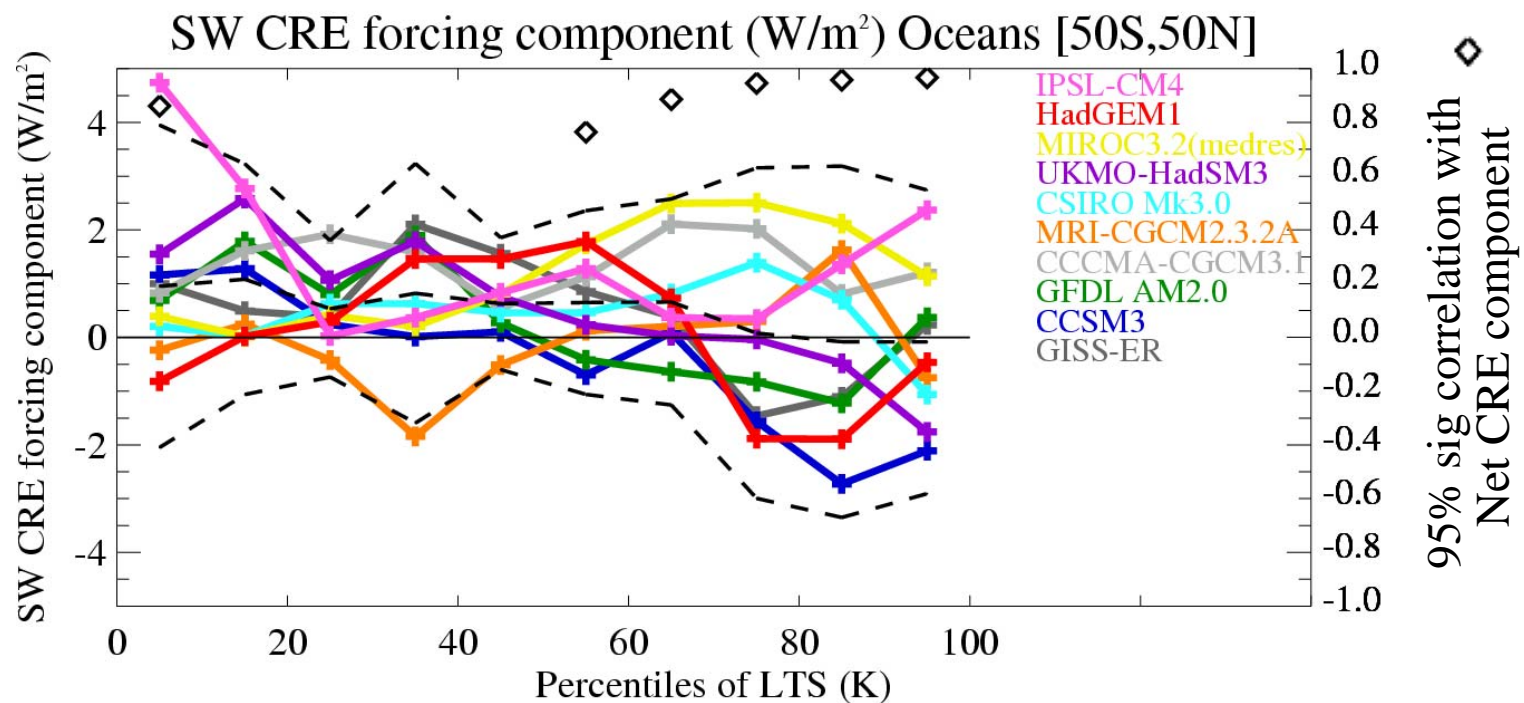
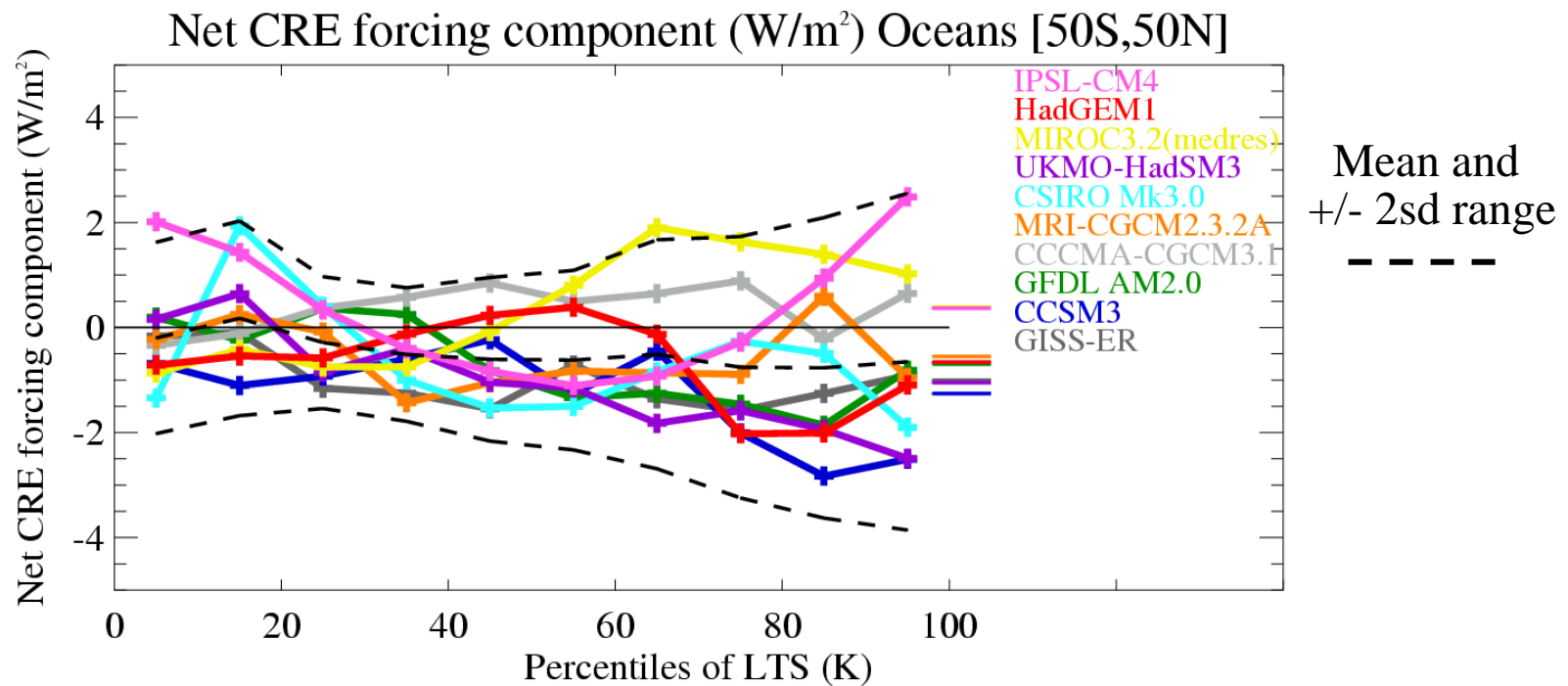
We examine CO₂ and temperature dependent responses in LTS and EIS (Estimated Inversion Strength, Wood and Bretherton, 2006) to see whether these are consistent between the models. If they are substantially different, then inter-model feedback and forcing differences are unlikely to be due solely to differences in local moist physics

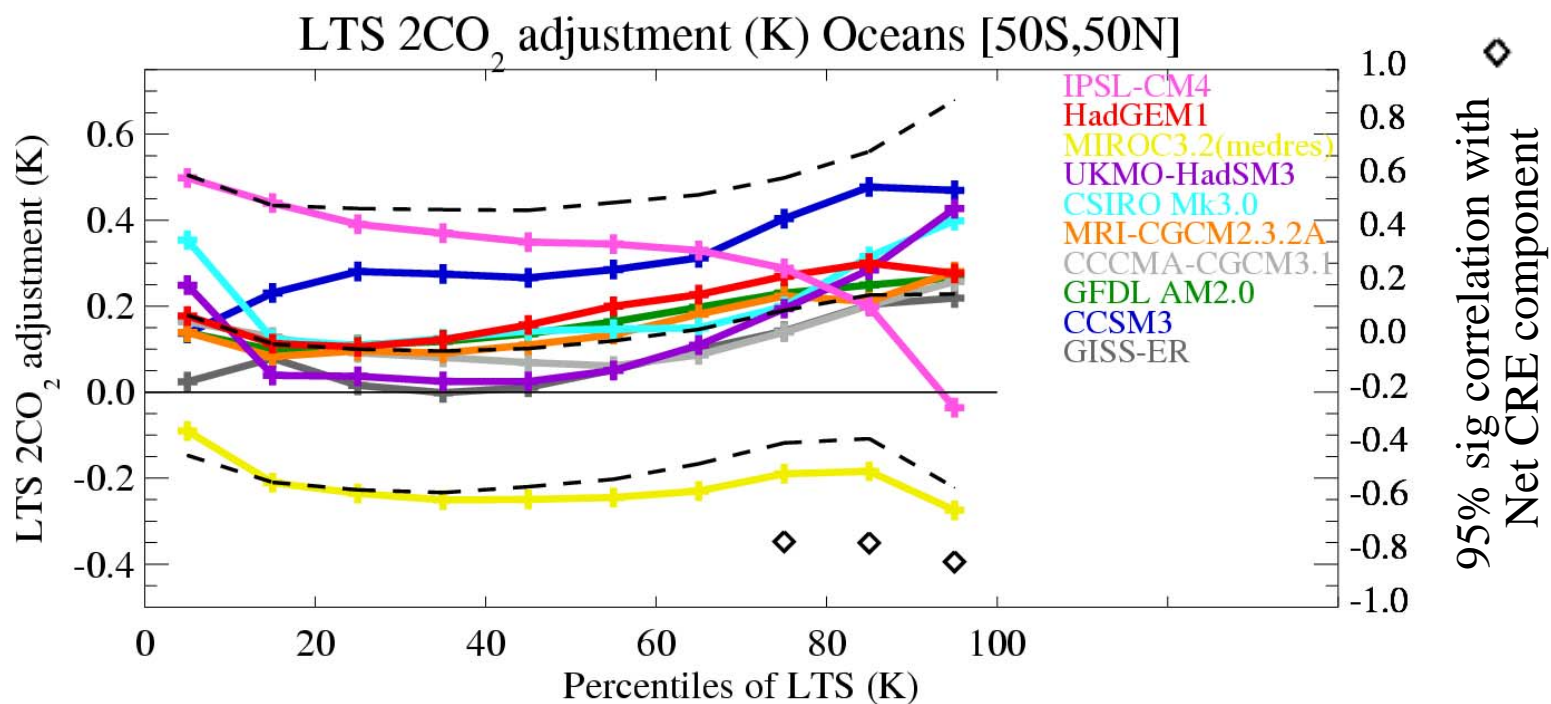
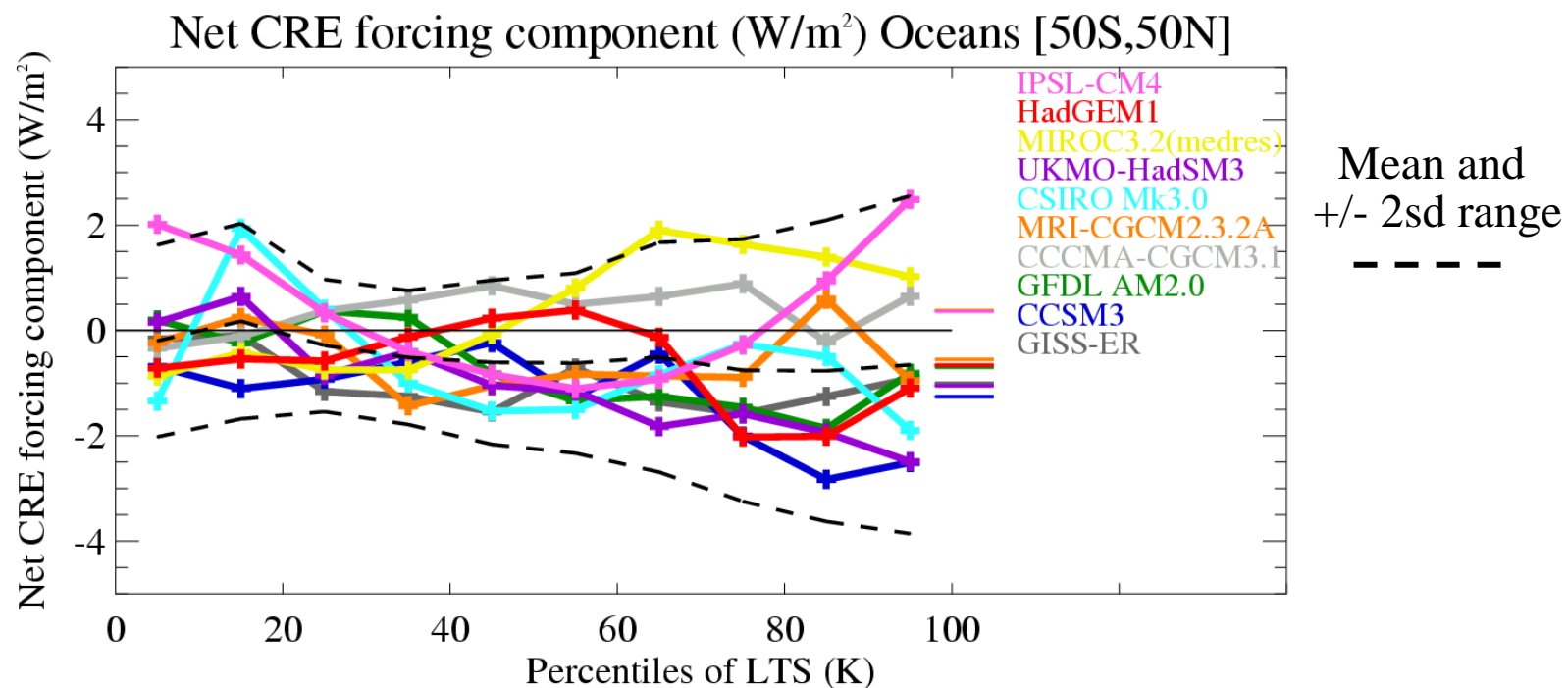
Early results from CFMIP-2 rapid cloud adjustment experiments will also be shown

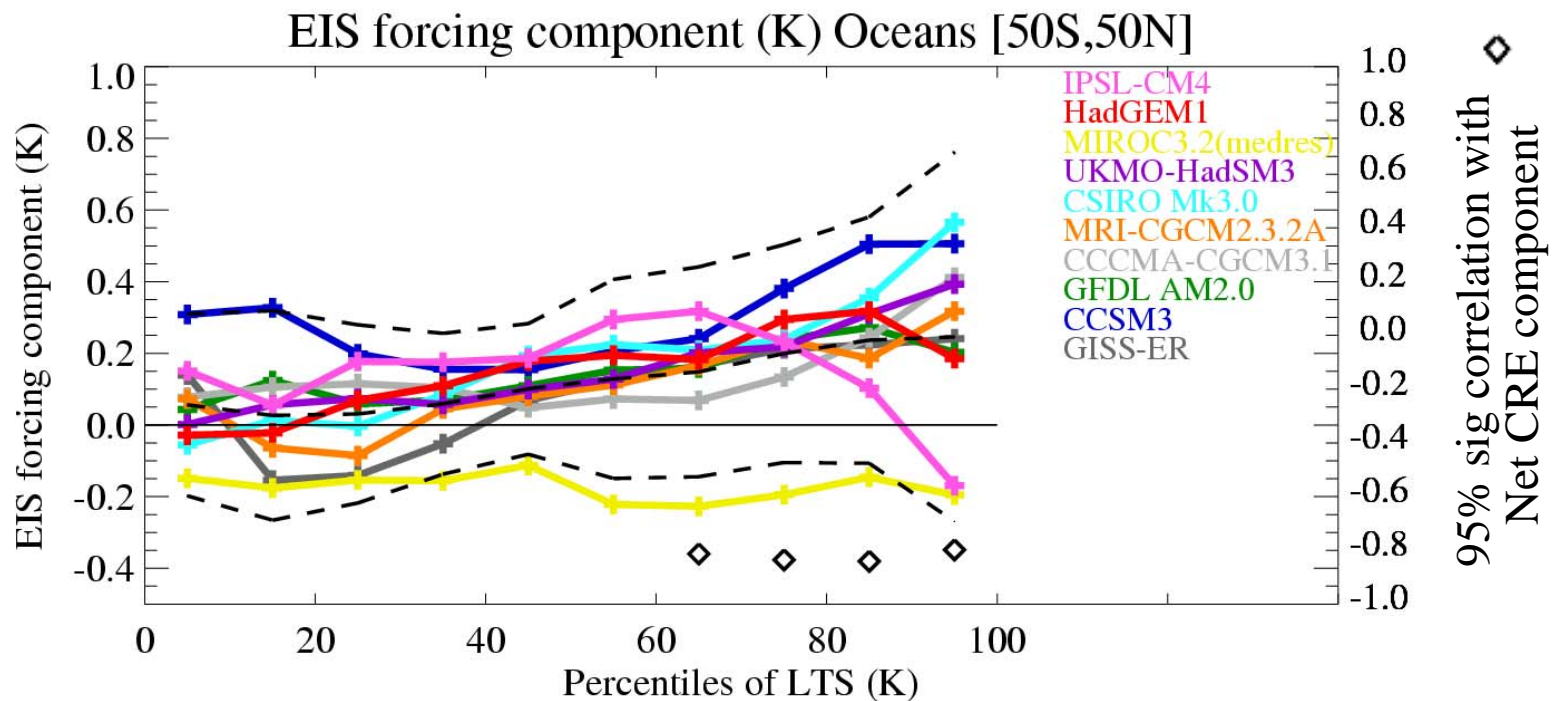
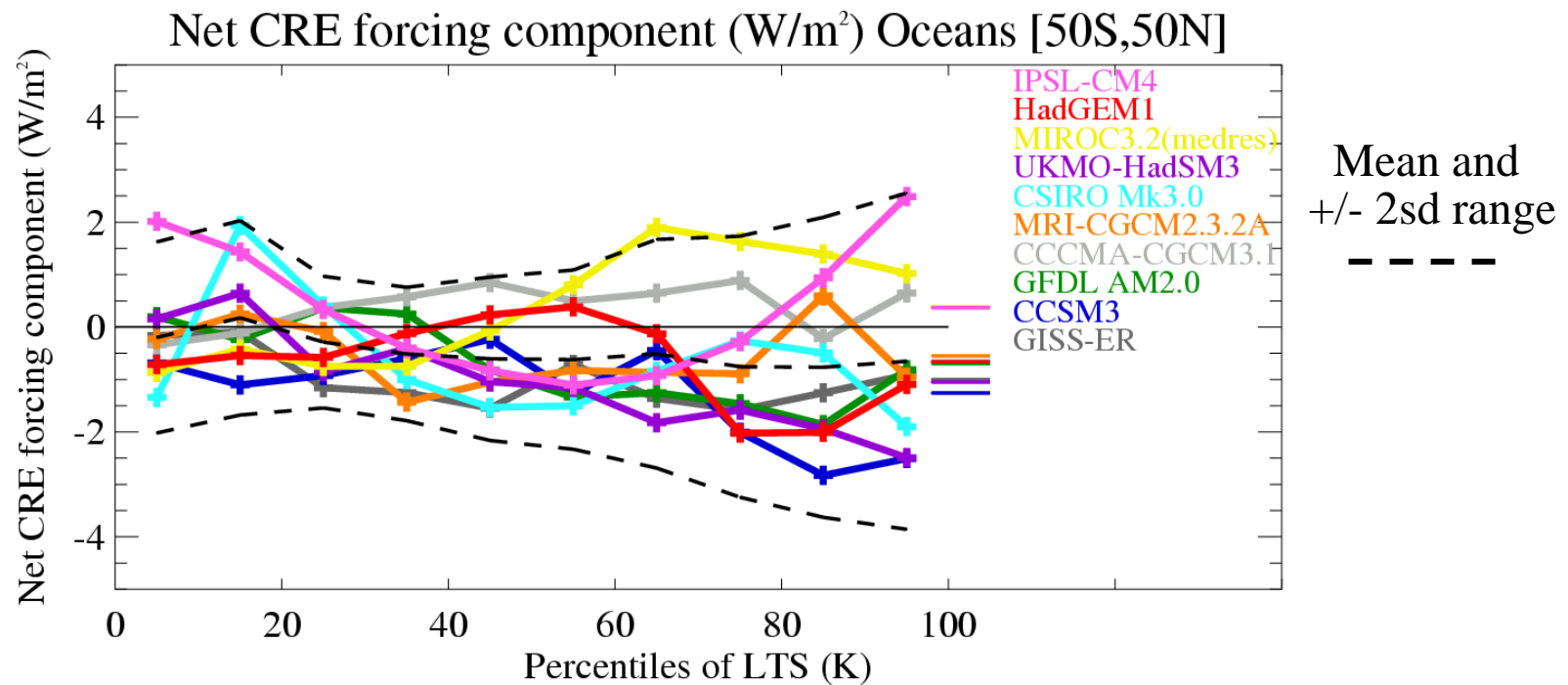




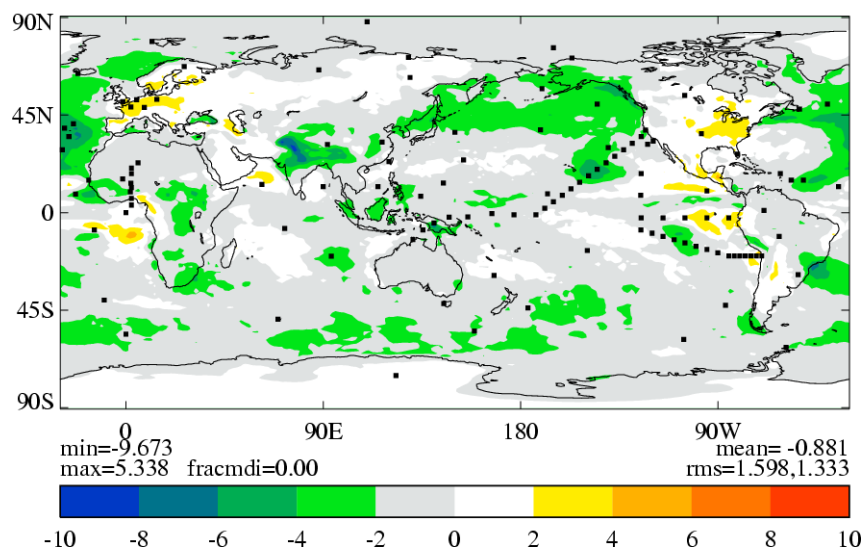




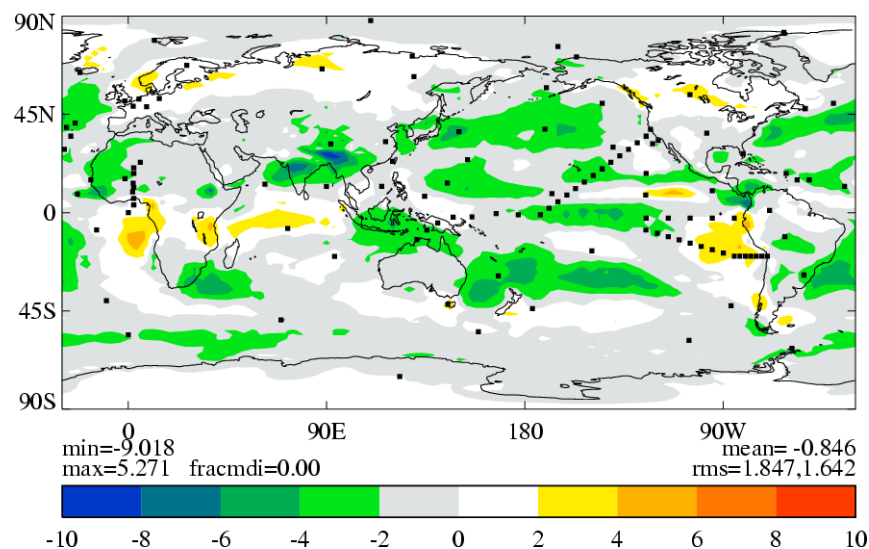




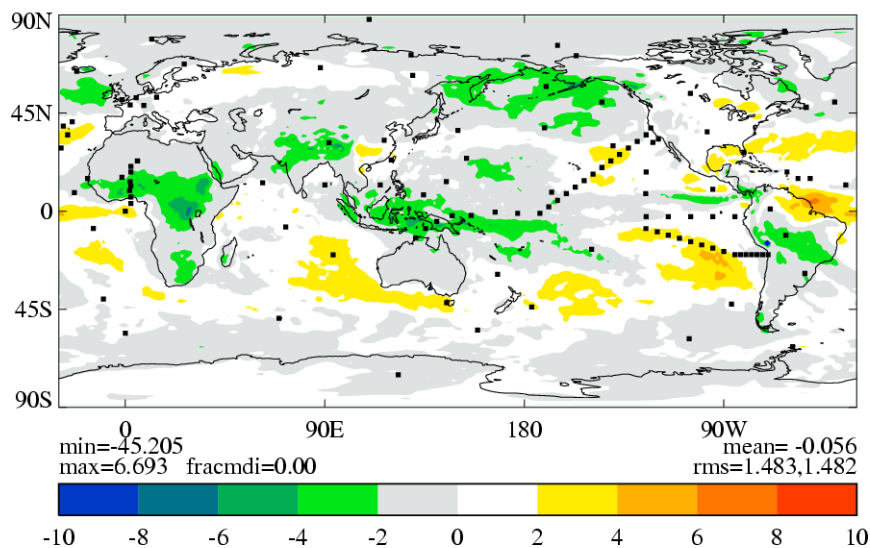
HadGEM2-A 4xCO₂ Fixed SST
Net CRE Response
Global 2xCO₂ Equivalent = -0.44 Wm⁻²



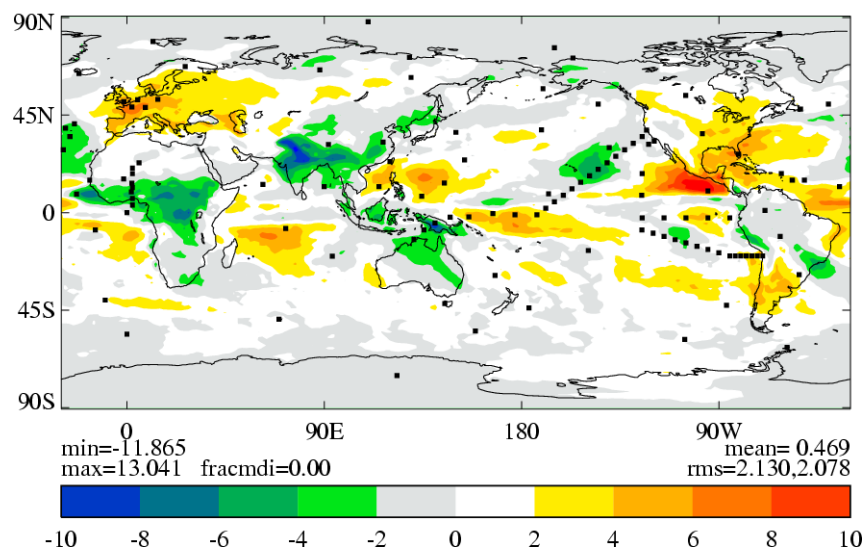
IPSL-CM5A-LR 4xCO₂ Fixed SST
Net CRE Response
Global 2xCO₂ Equivalent = -0.42 Wm⁻²



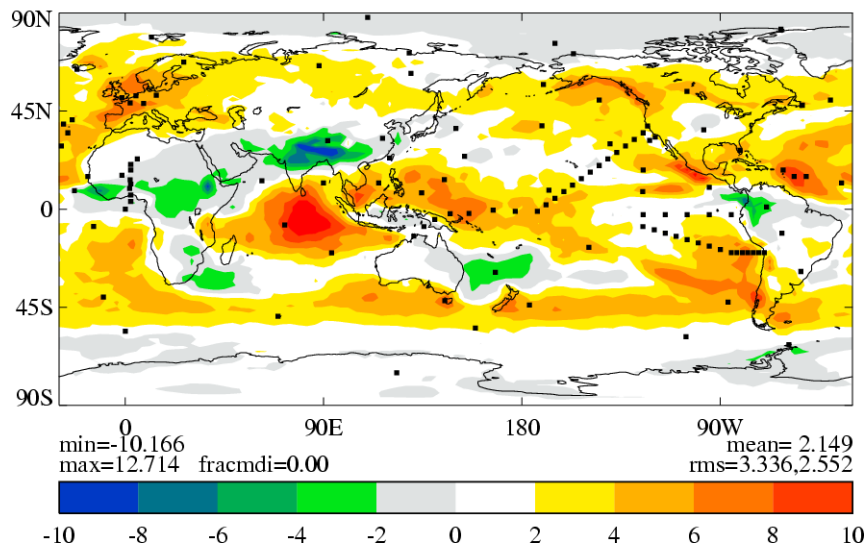
CNRM-CM5 4xCO₂ Fixed SST
SW CRE Response
Global 2xCO₂ Equivalent = +0.03 Wm⁻²



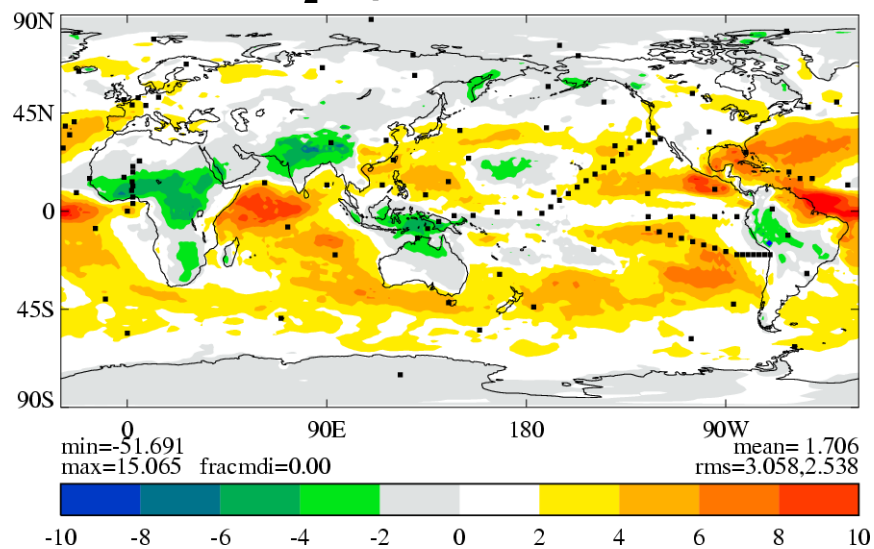
**HadGEM2-A 4xCO₂ Fixed SST
SW CRE Response
Global 2xCO₂ Equivalent = +0.23 Wm⁻²**



**IPSL-CM5A-LR 4xCO₂ Fixed SST
SW CRE Response
Global 2xCO₂ Equivalent = +1.07 Wm⁻²**



**IPSL-CM5A-LR 4xCO₂ Fixed SST
SW CRE Response
Global 2xCO₂ Equivalent = +0.85 Wm⁻²**





Conclusions (1)

Cloud components of climate feedback are positive in all models in stable oceanic areas, in spite of increases in LTS in all models and increases in EIS in most models

Spread in cloud feedback is largest in stable regions, suggesting that stratocumulus/transition clouds are a leading order cause of spread in climate sensitivity in this ensemble

Large inter-model differences in LTS and EIS responses in stable regions mean that we can not safely assume that differences in feedbacks are due to local moist physics alone

Conclusions (2)

Cloud components of CO₂ forcing take positive and negative values and again show largest differences in stable regions

CAM3 has strong increases in stability in all bins, and the strongest negative cloud component of CO₂ forcing, while MIROC shows the opposite behaviour

The cloud components of the forcings are strongly anti-correlated with changes in LTS and EIS in the stable regions

This suggests that remote processes (e.g. deep convection) may be controlling the different cloud responses via stability changes in some models

Conclusions (3)

Substantial differences in stability responses in stable regions coincident with differing cloud components of CO₂ forcing and feedback suggest a need for studies following on from CGILS to consider a range of stability forcings

Early results from CFMIP-2 experiments with 4xCO₂ and fixed SSTs show substantial positive shortwave cloud responses over oceans in low and high cloud areas.

Negative SW responses are seen over land are suggestive of a monsoon-like response.