

# Overview of LES Results for S11 Forcings

**Peter Blossey**<sup>(#)</sup> and Chris Bretherton, U Washington

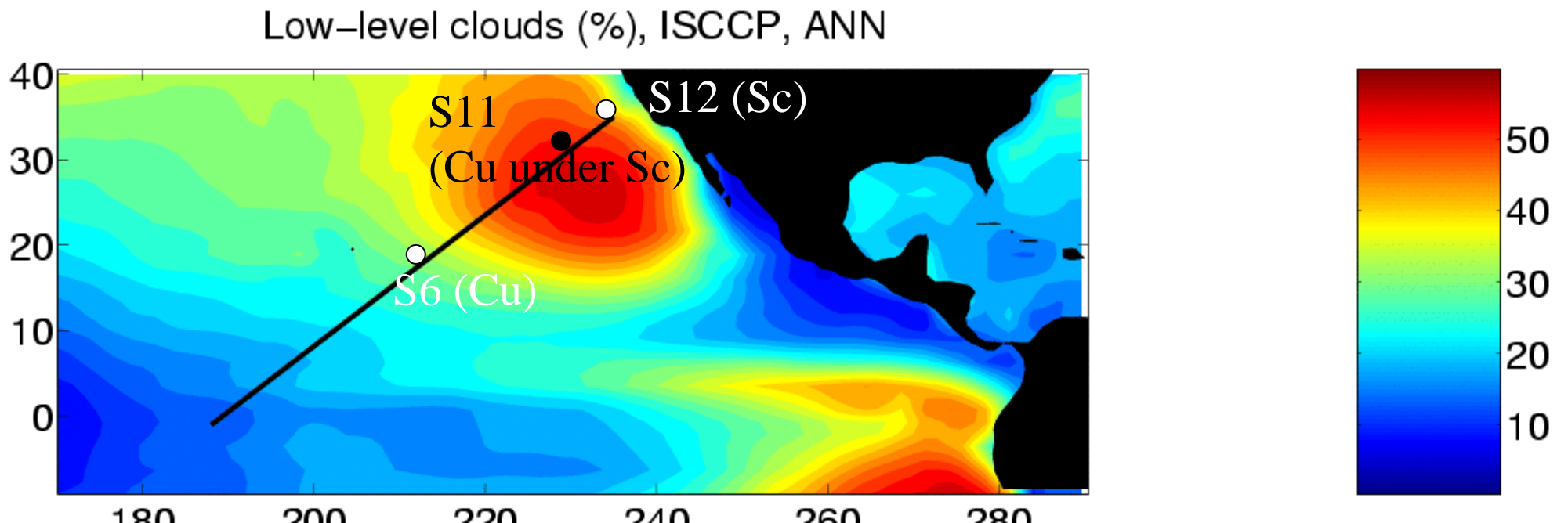
Minghua Zhang and Marat Khairoutdinov, Stony Brook U.

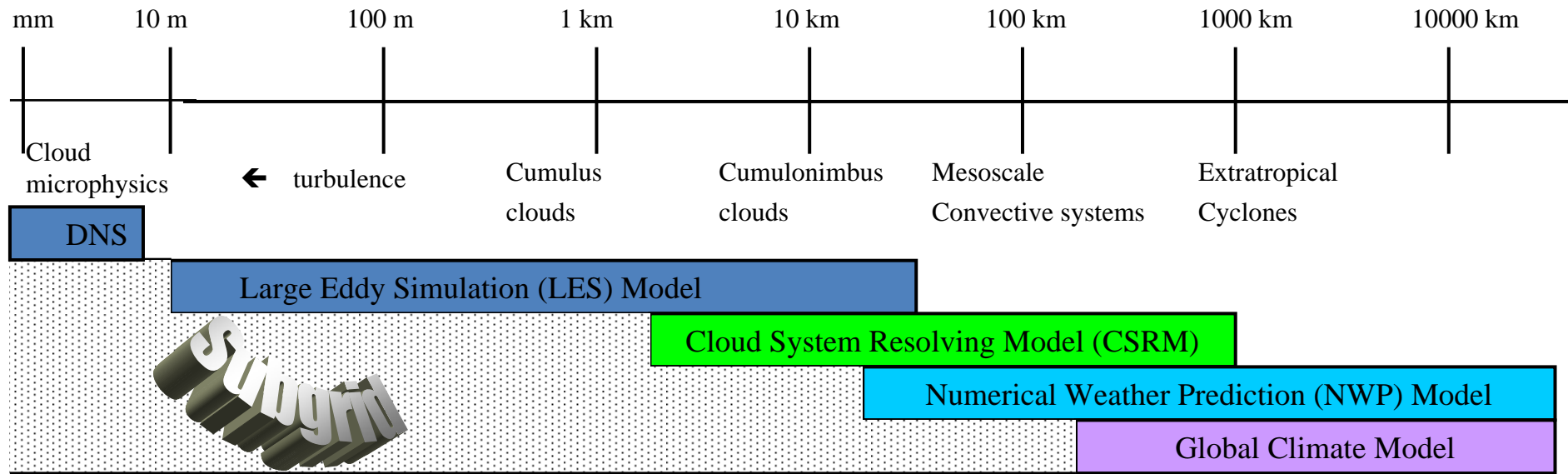
Anning Cheng, NASA LARC

Adrian Lock, UKMO

Stephan de Roode, TU Delft

Irina Sandu and Thijs Heus, MPI



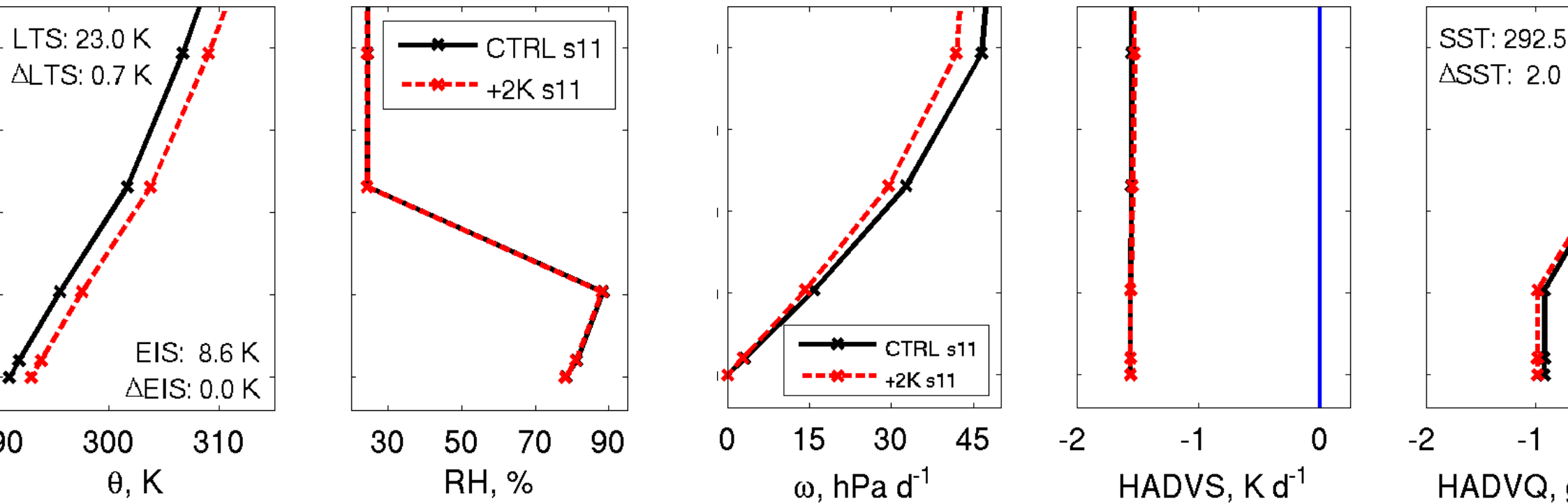


courtesy Pier Siebesma

Length scales that dominate the turbulent transport are explicitly resolved by LES

Subgrid scales ( $< \sim 50\text{-}100$  m) are parameterized

Hope: LES as a higher-fidelity benchmark for single column model (and GCM) cloud response



Control Forcings: JJA climatology from 3 **GCSS Pacific Cross-section** points

to steady state with **diurnally-averaged** insolation

Vertical profiles of temperature, humidity and ozone above the LES domain are used for radiative transfer

## **Simulation time**

10 days

Adaptive time step, dtmax = 10 secs

Radiation time step = 60 secs

## **Domain size**

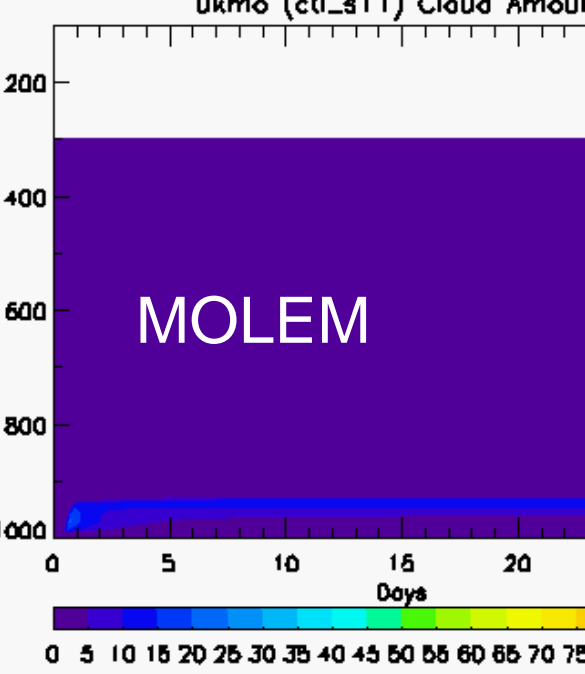
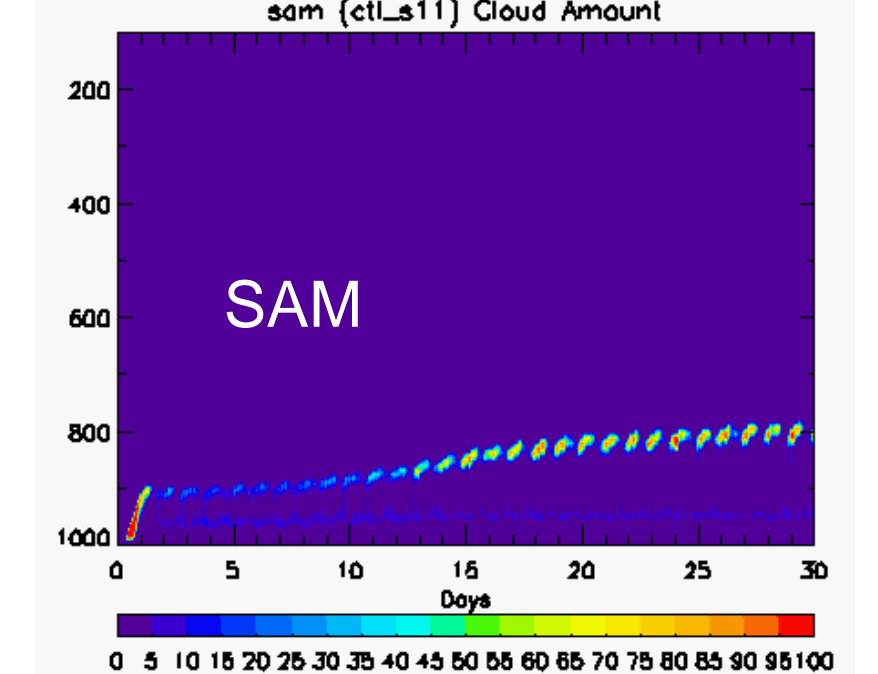
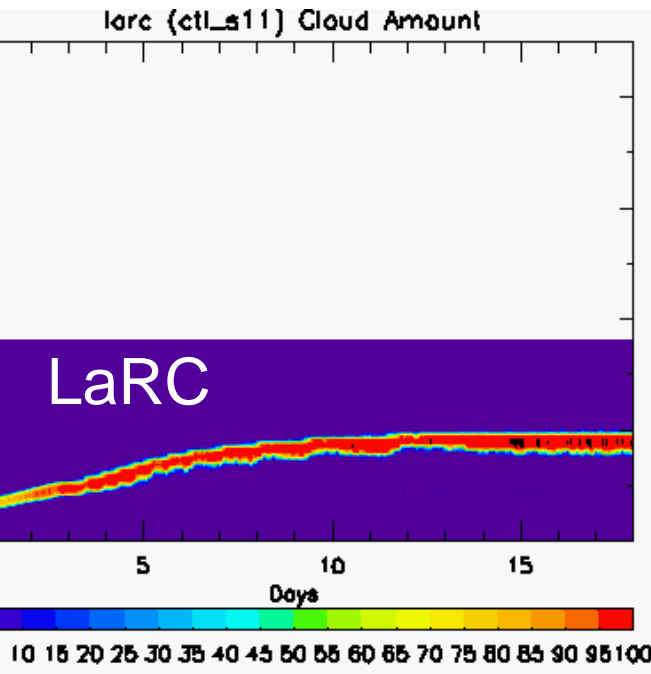
4.8 x 4.8 x 4 km<sup>3</sup>, 96 x 96 x 128 grid points ( $\Delta z = 25$  m in lower part)

Additional layers above LES domain for radiation

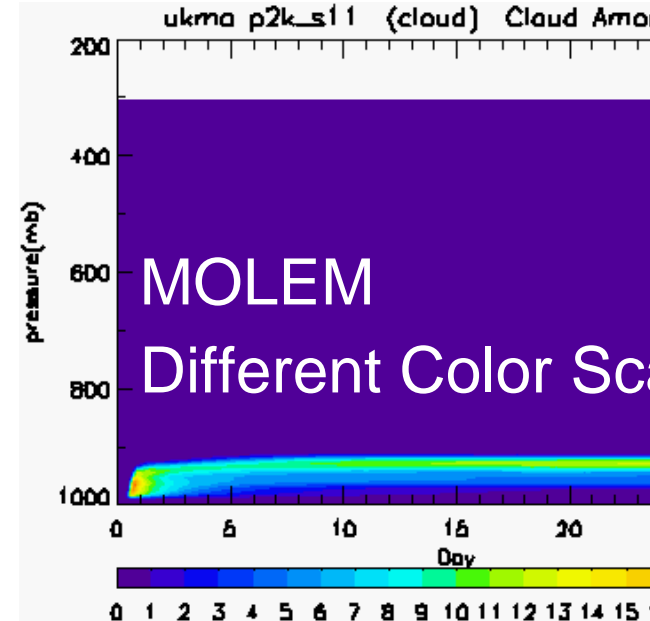
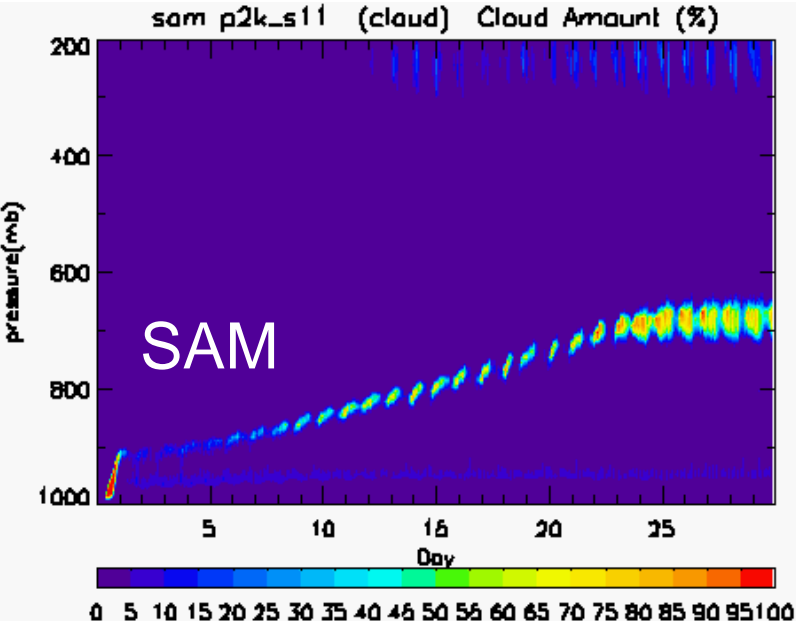
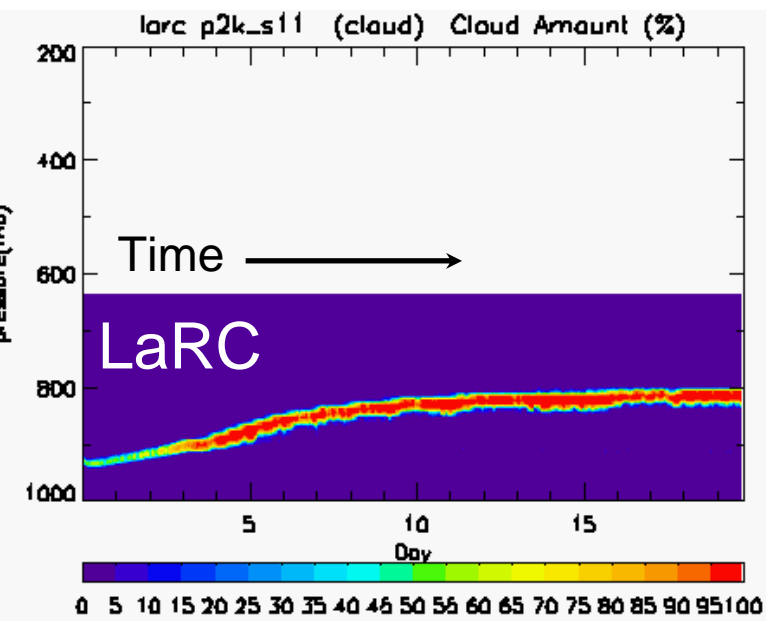
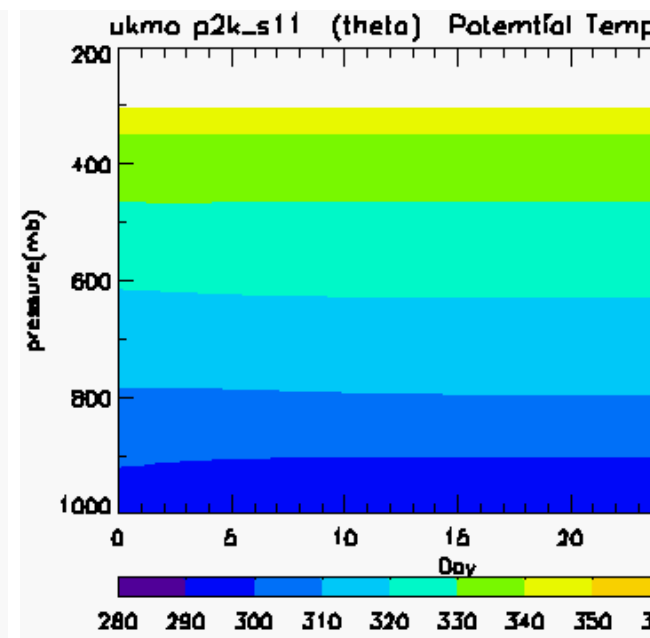
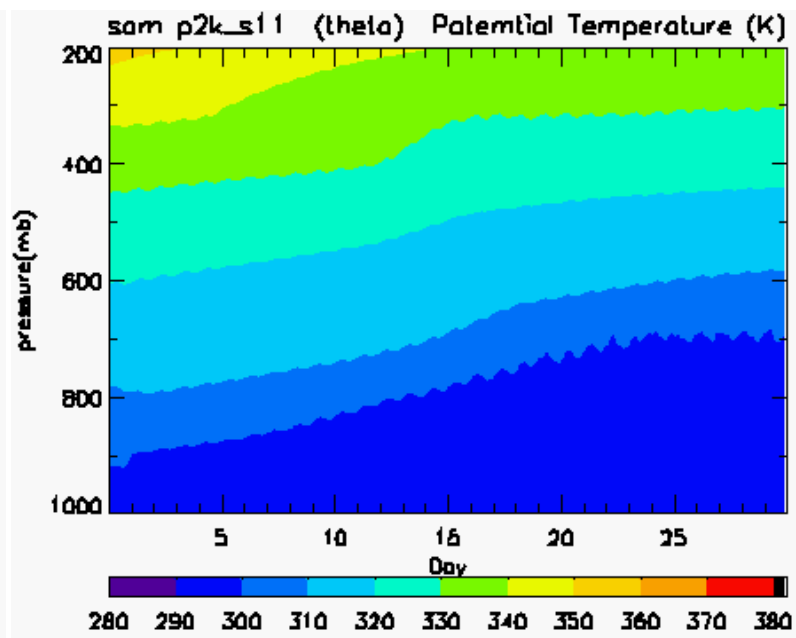
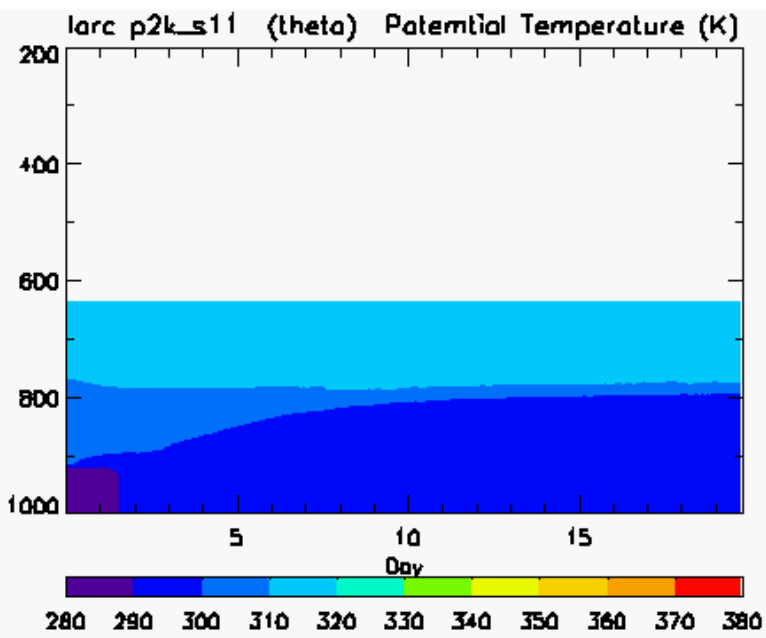
## **Total CPU hours on 32 processors**

2700 hours

## Cloud Amount (red color = 100% cloud fraction)

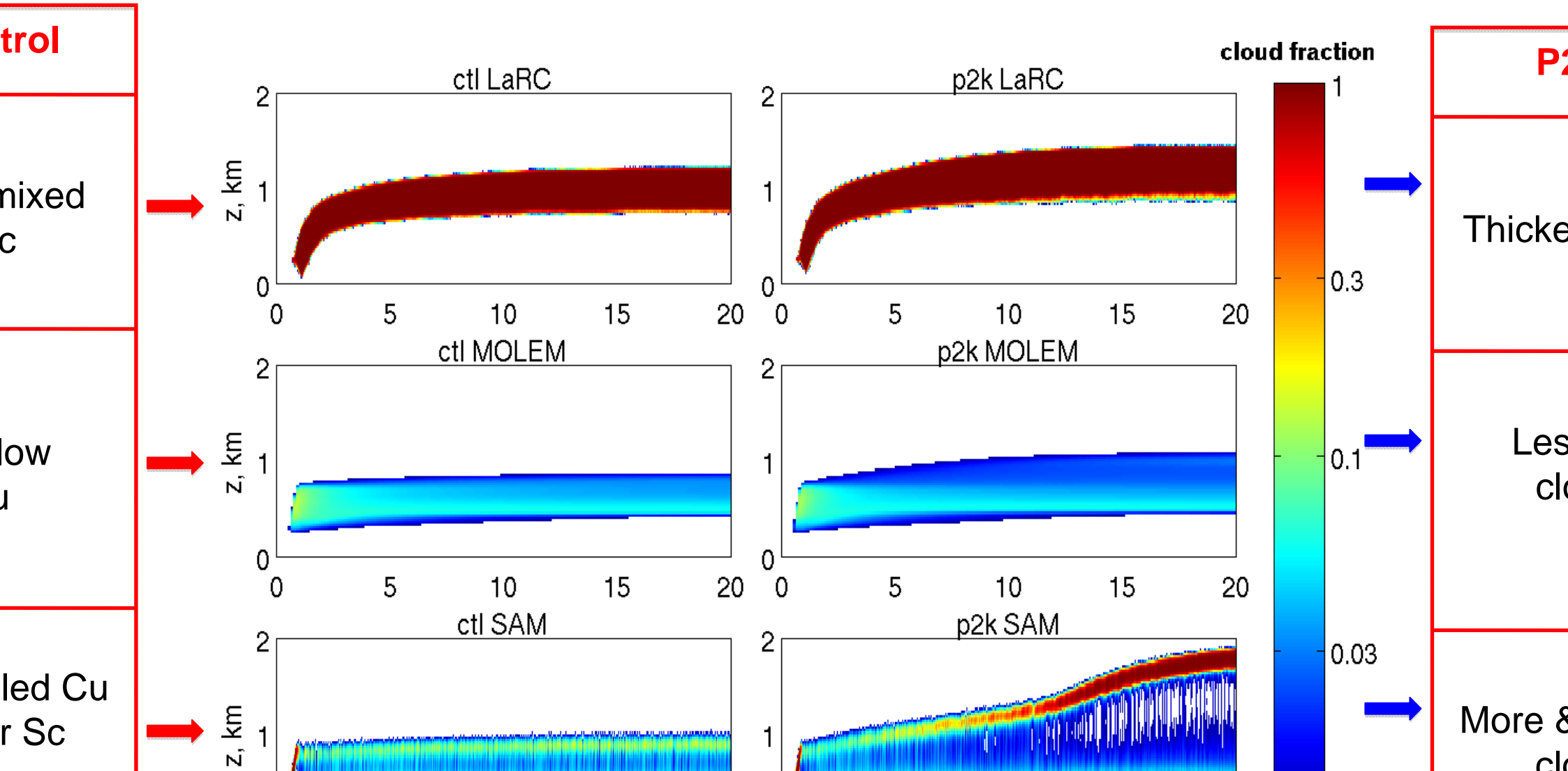


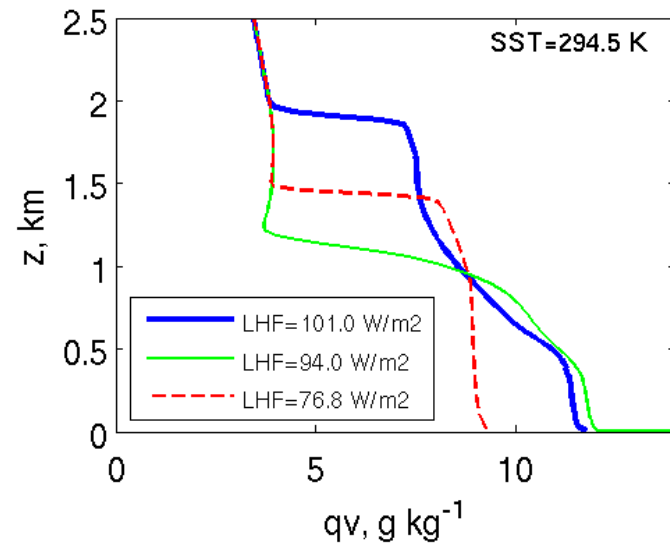
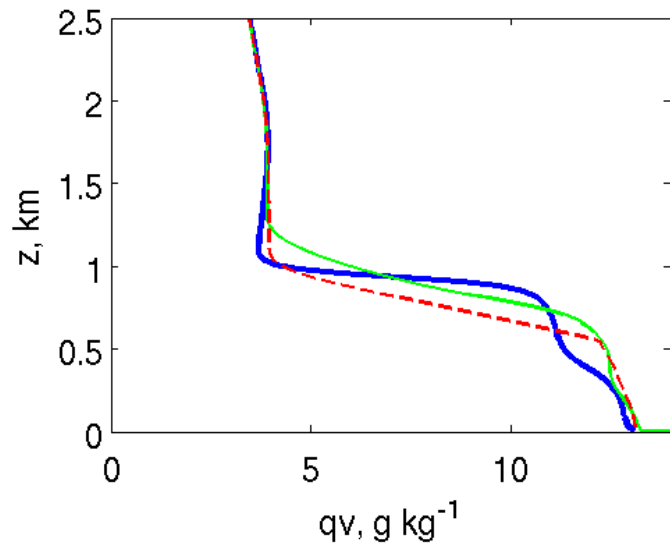
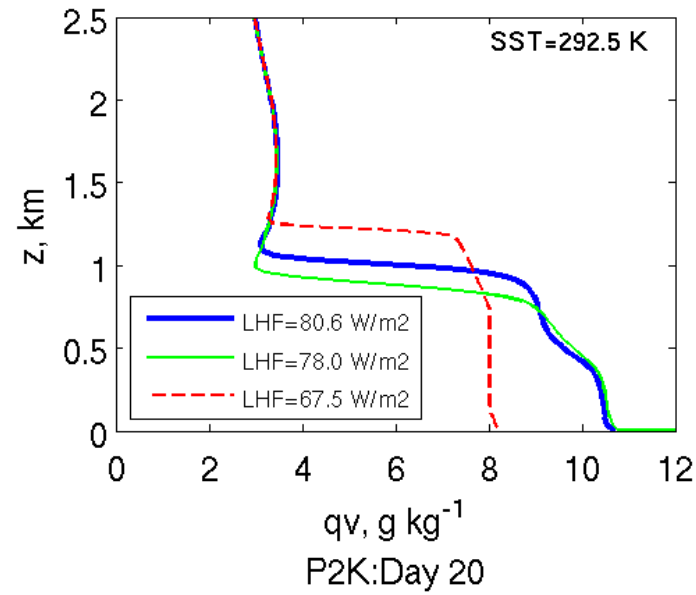
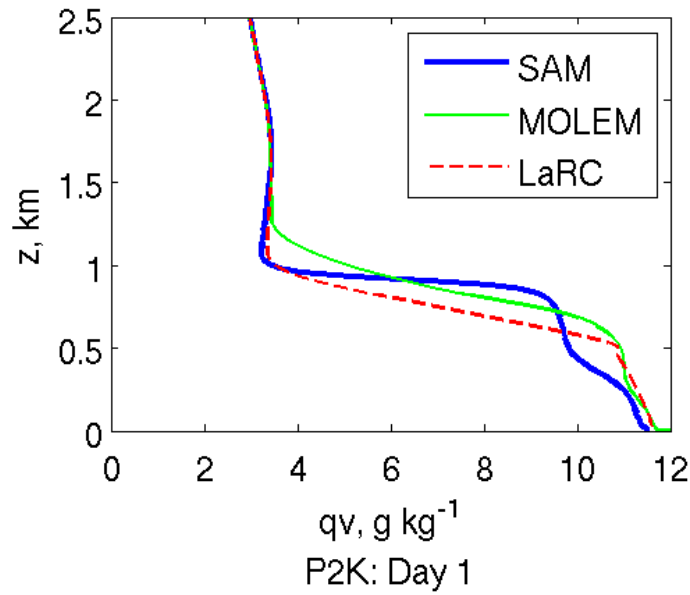
Time →



## Round 2: Case modifications

Nudge temperature above 2500m. Subsidence should help maintain free-trop temperature  
Have models use same radiation scheme (RRTMG)





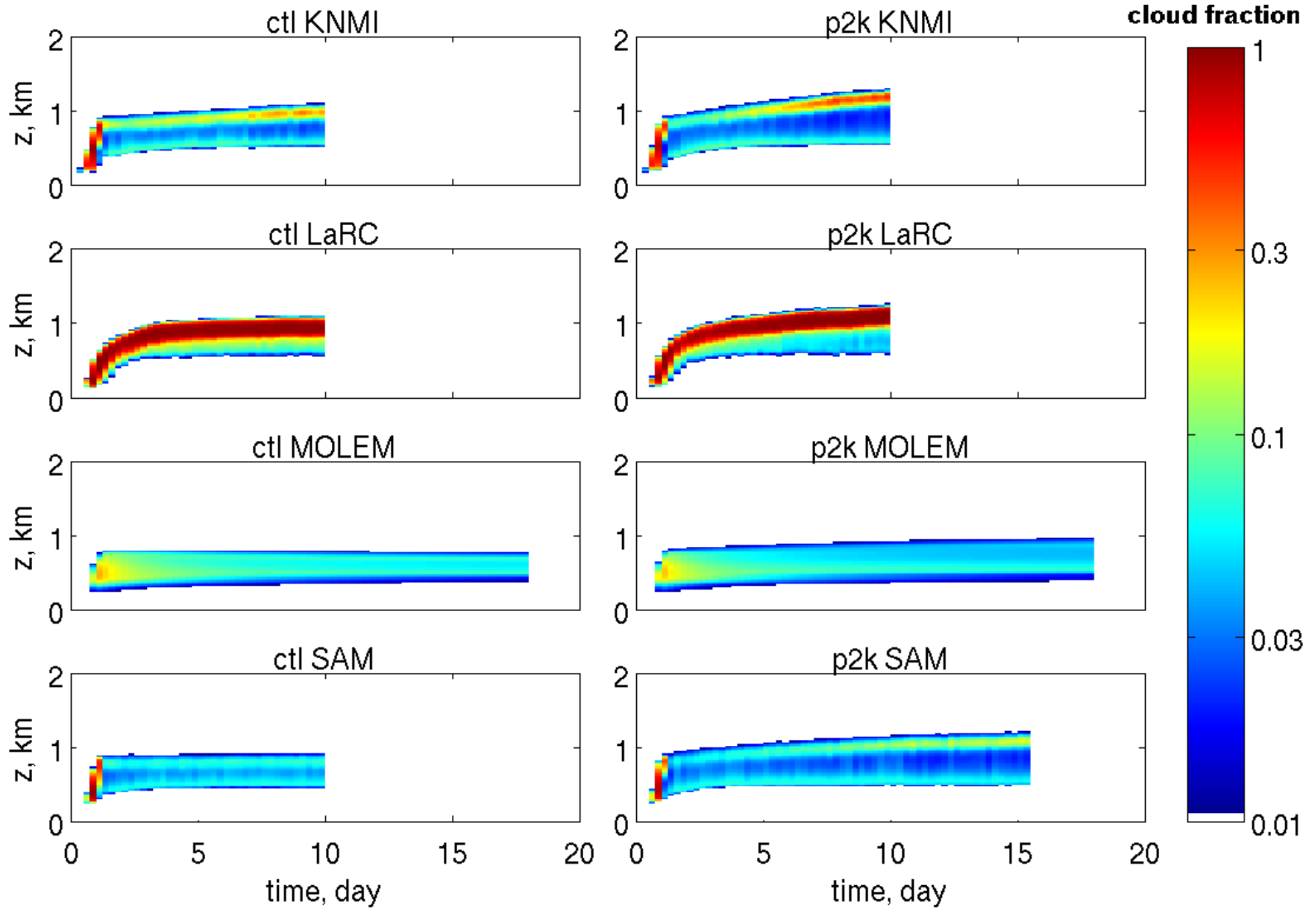
### Round 3 modifications

specified surface transfer coefficient:

$$\overline{w'q'} = C_h(C_p SST - s_{surf}^{air})$$

$$\overline{w'q'} = C_q(0.98q_{sat}(p_{surf}, SST) - q_{surf}^{air})$$

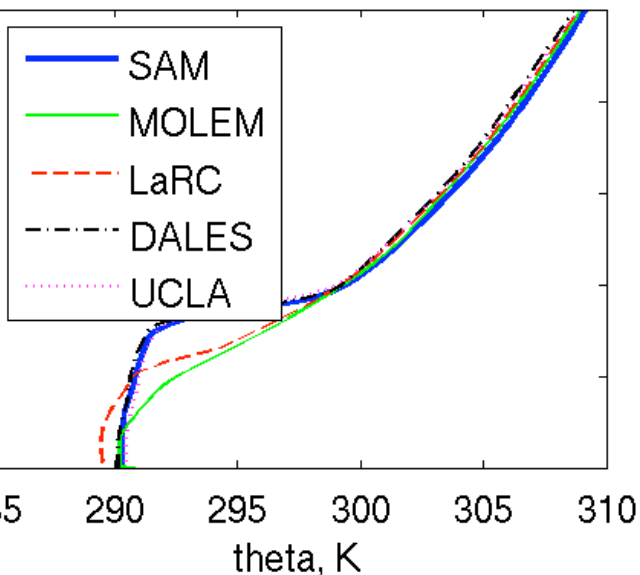




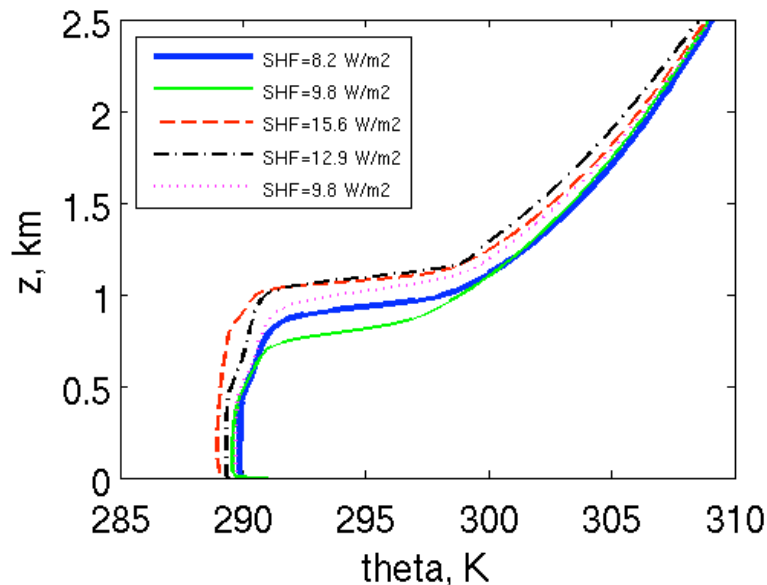
ten day runs requested while we're working out the final kinks in the setup.

MI deepening steadily through day 10, with more inversion cloud than SAM, MOLEM.

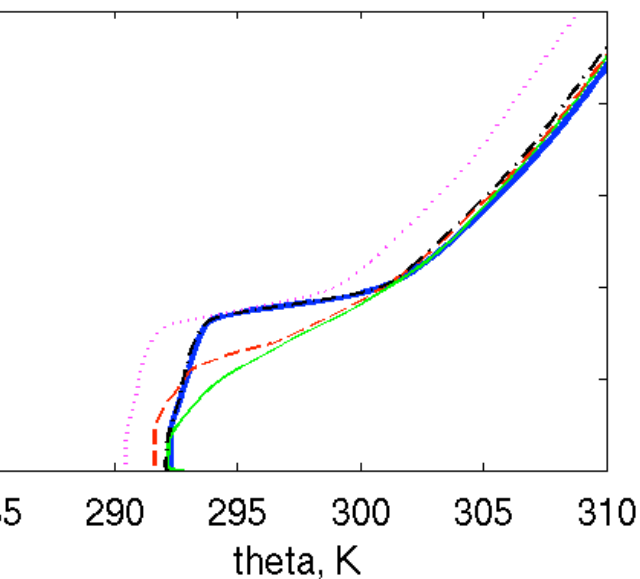
CTL: Day 1



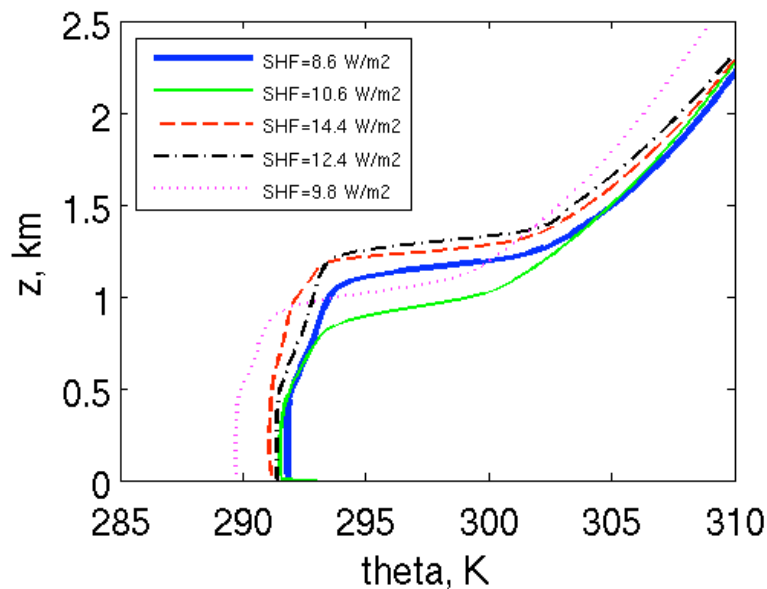
CTL: Day 10



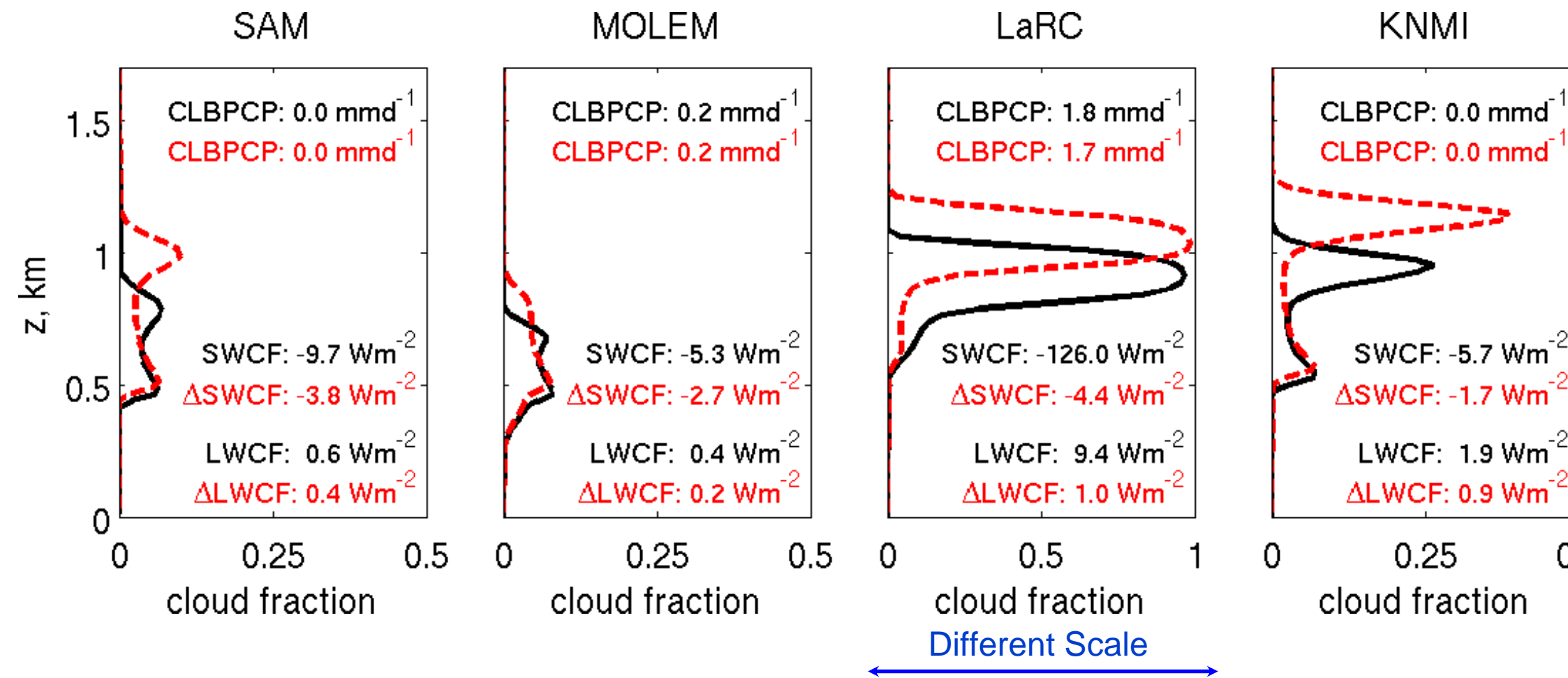
P2K: Day 1



P2K: Day 10



- Near-surface conditions more consistent with surface transfer coefficients
- Still a diversity of inversions at different heights, cloud state
- SAM, MOLEM slightly warmer in free troposphere than DALES, LaRC
- Caused by above-cloud cooling? Radiation scheme? DALES uses Fu-Liou (Pincus & Stevens, 2009).

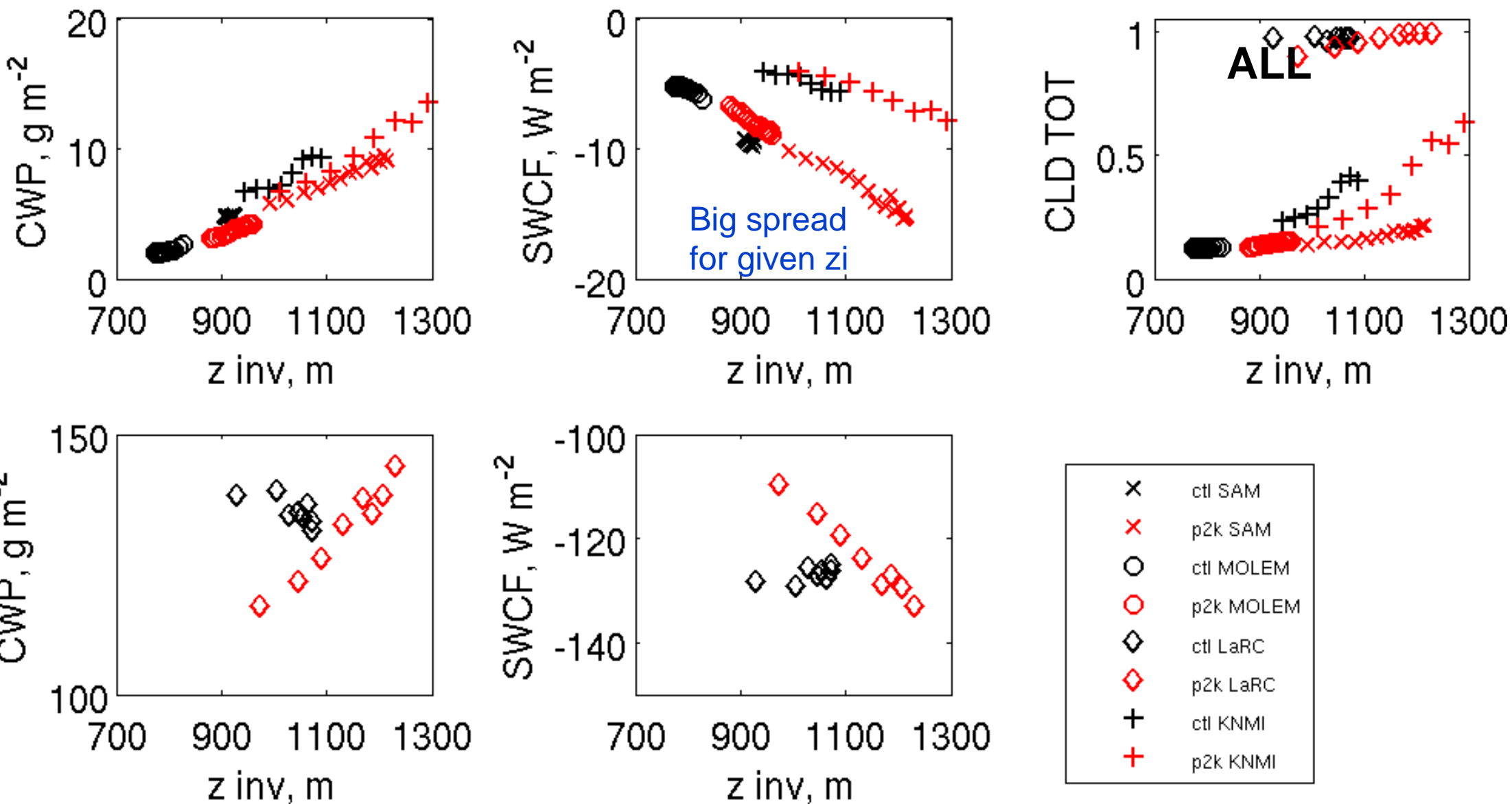


**Note:** Not equilibrium solutions: day 7.5-10 averages

Inversion deepens in all models

Below-inversion cloud increases in KNMI and SAM

Progress here (certainly compared with last June) However still a ways to go (not yet equilibrium)

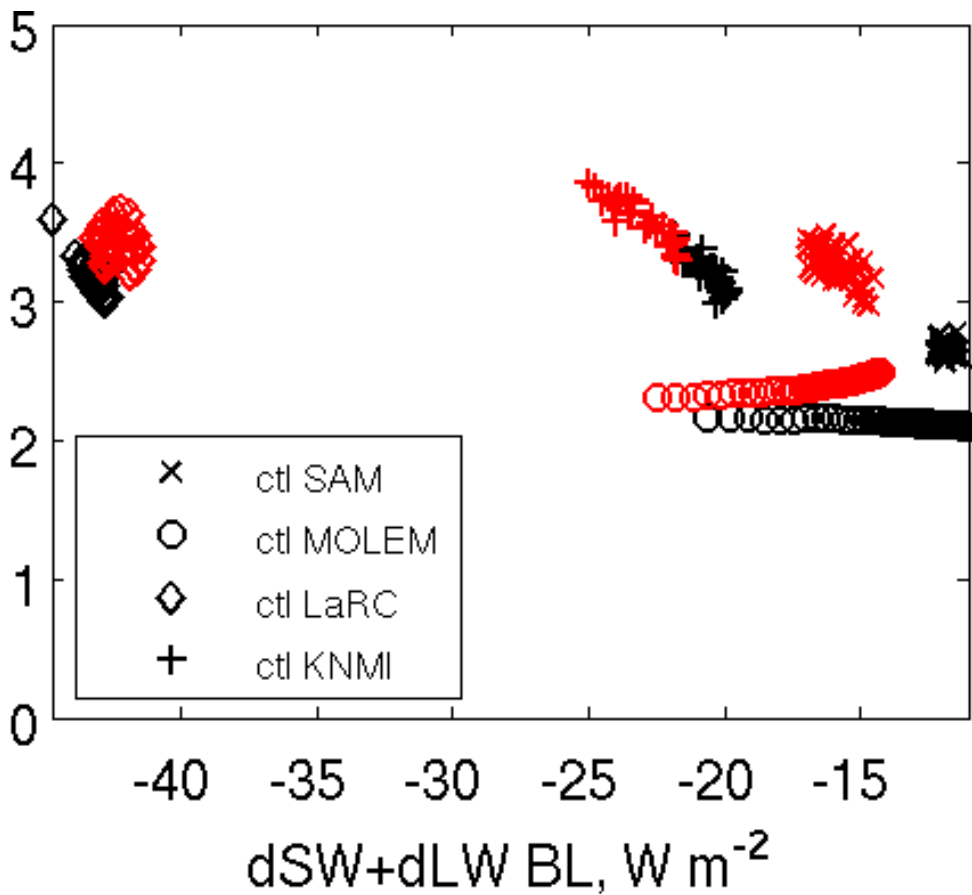


2K runs generally similar for similar  $z_{inv}$  after initial transient (Day 3 on)

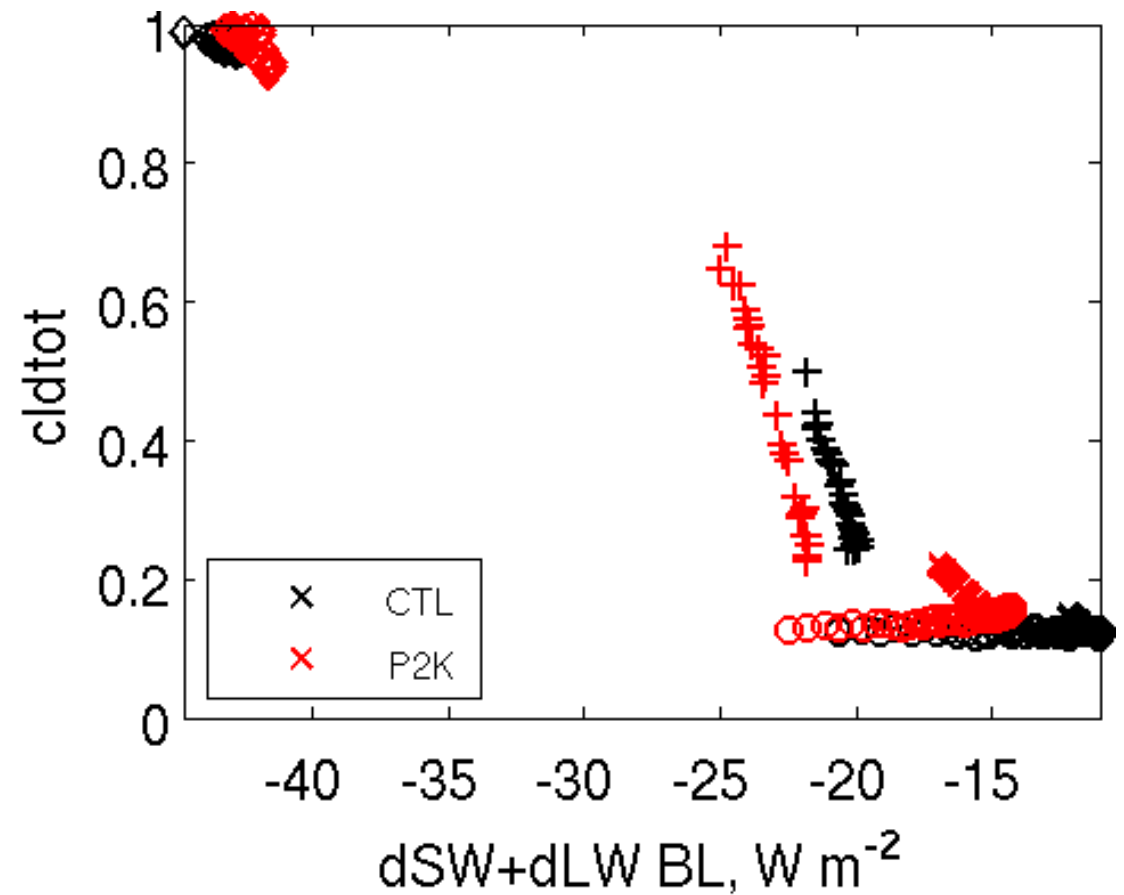
SWCF, CWP changes driven in part by increase in inversion height

deepening of inversion in P2K runs exaggerated by artificially large decrease in low-level subsi

Entrainment rate  $\propto$  radiative cooling (except LaRC)



Radiative cooling  $\propto$  cloud fraction



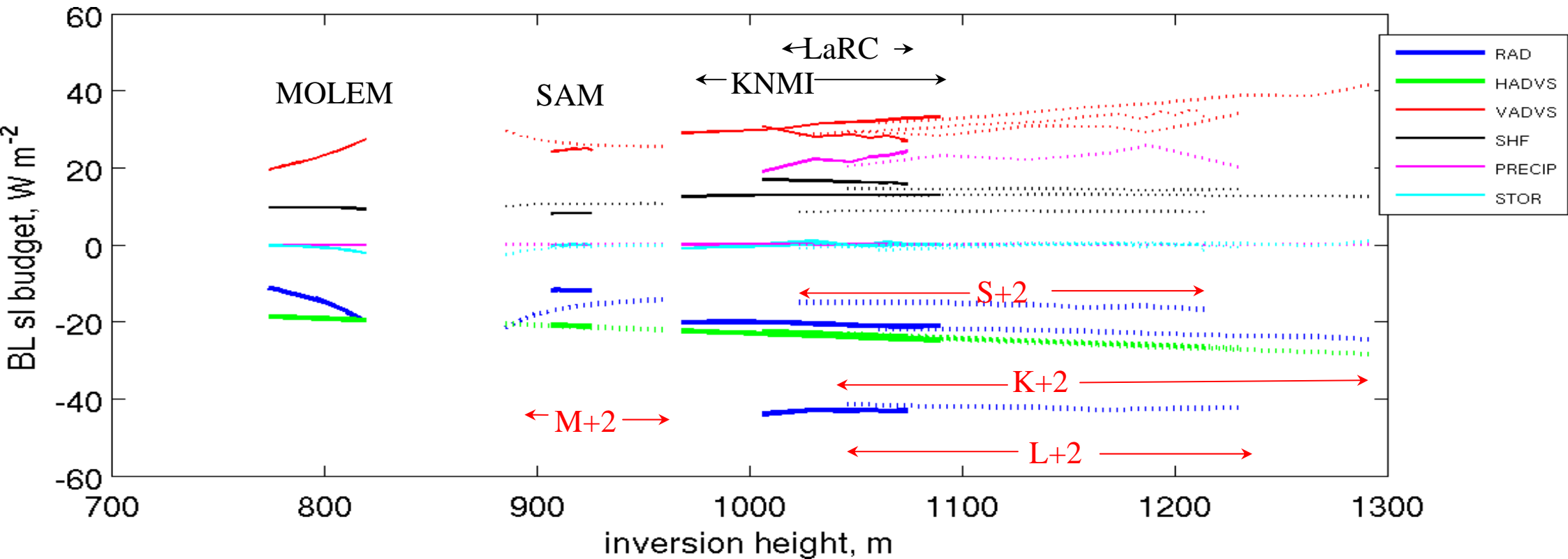
Larger cloud fraction  $\rightarrow$  More BL radiative cooling  $\rightarrow$  More entrainment

$$w_e = dz/dt - w(z_i), \quad z_i(t) = \text{height at which mean RH} = 50\%$$

$$\begin{aligned}
 & \underbrace{\rho \frac{\partial s_l}{\partial t}}_{\text{Storage}} dz + SHF - \underbrace{\overline{\rho w' s_l}}_{\text{Entrainment}} \Big|_{z_i+50} - \int_{z_i-50}^{z_i+50} \rho \frac{\partial s_l}{\partial t} dz - \int_0^{z_i+50} \rho w \frac{\partial s_l}{\partial z} dz - \int_0^{z_i+50} \rho (\mathbf{u}_h \cdot \nabla_h s_l) dz - F_{rad} \Big|_{z_i+50} + \dots
 \end{aligned}$$

Storage
'Entrainment'
Horizontal advection
Radiation
Late

Entrainment warming + SHF  $\approx$  Radiative Cooling + Hor. Adv.  
(VADVS) (-RAD) (-HADVS)

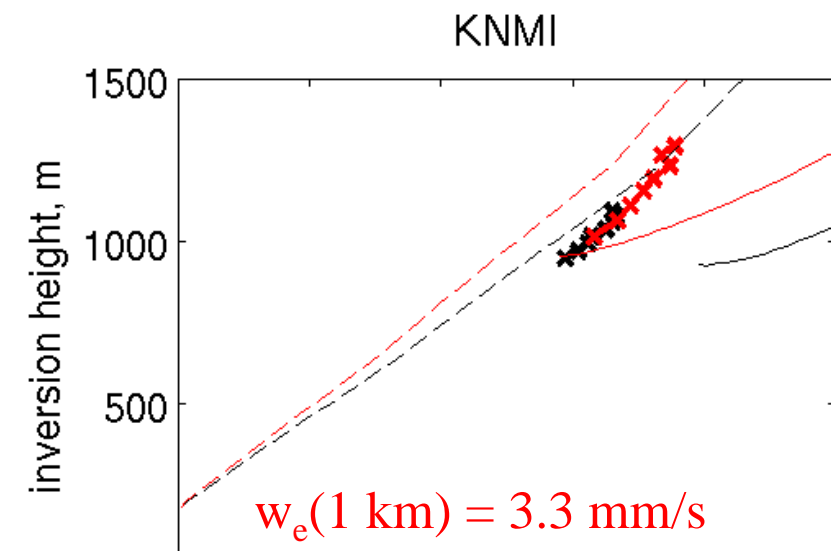
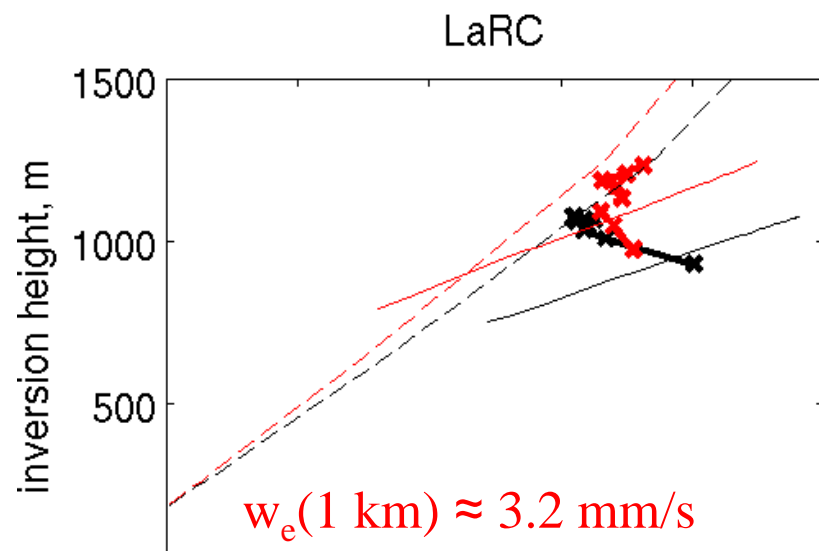
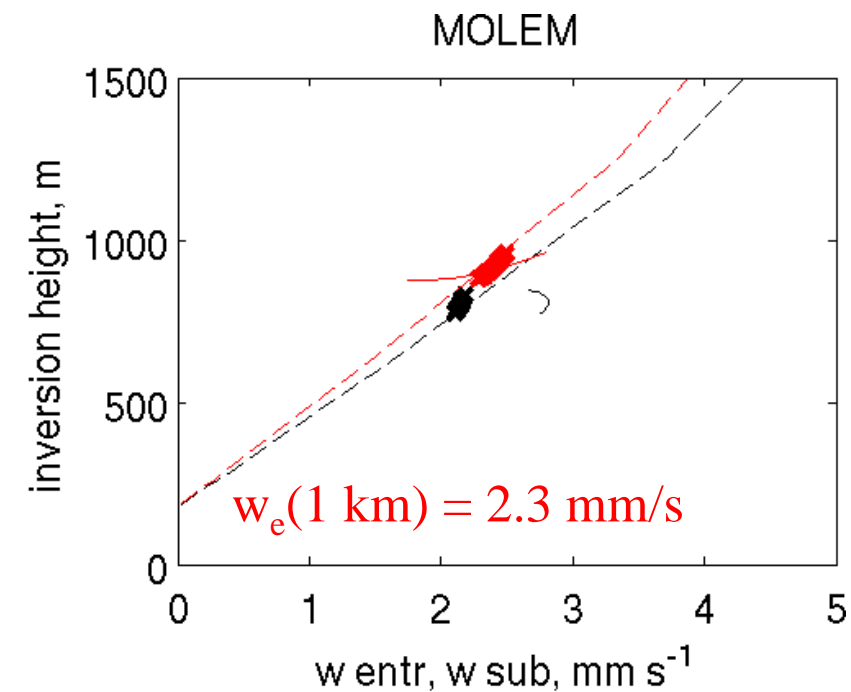
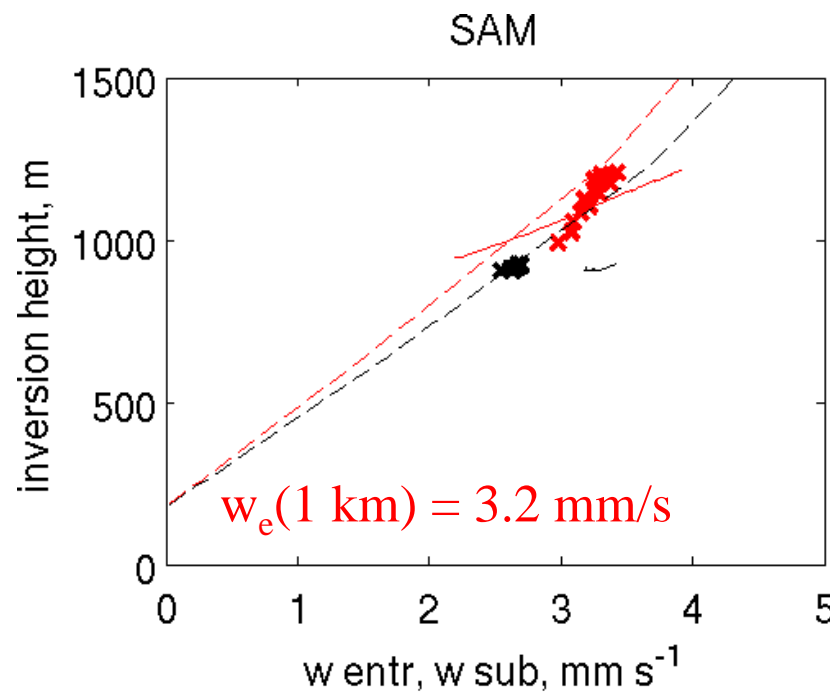


More cloud  $\rightarrow$  more red cooling  $\rightarrow$  more entrainment  $\rightarrow$  deeper DPL

Only LaRC has a strongly stable steady state.

Other LESs have  $dw_e/dz_i$  almost as large as  $-dw/dz_i$  (barely stable) because cloud fraction, radiative cooling, adv. cooling increase with

The p2K and ctl runs have surprisingly similar  $w_e(z_i)$ .



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LES setups have been harmonized:

temperature nudged aloft,

same surface scheme,

mostly same radiation,

considerable debugging of setup, inter-model differences.

Despite diversity of mean BL state, models show similar cloud response.

Cloud response appears to be driven by increase in  $Z_{inv}$

How much should  $\omega$  decrease in warmer climate?

Models still show tendency to deepen gradually...



in changes to the forcing data are:

Temperature and moisture fields from the ECMWF analysis for July 2003 are used as the control case to facilitate comparison with observations. For the warmer p2K temperature and water vapor profiles, moist adiabatic changes and constant relative humidity are used

Vertical resolution in the forcing data is increased to 62 levels as in the ECMWF analysis

recomputing omega consistent with RRTM radiation calculations, with omega going to zero at the surface rather than at 1000 hPa

The setup of the large-eddy simulations is standardized as follows:

run ten days of ctl and p2k for each case.

S5: nudge T, q above 4000m,  $\Delta x = \Delta y = 100\text{m}$ ,  $\Delta z = 40\text{m}$  up to 4km.

L1: nudge T, q above 2500m,  $\Delta x = \Delta y = 50\text{m}$ ,  $\Delta z = 25\text{m}$  up to 2500m.

L2: nudge T, q above 1200m,  $\Delta x = \Delta y = 25\text{m}$ ,  $\Delta z = 5\text{m}$  up to 900m.

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CLIPSE actively involved in CGILS-LES experiments

MO (Lock)

D (De Roode)

PI (Sandu/Heus/Rieck)

LES model: Warsaw

30 September: new results will be presented during the EUCLIPSE/GCSS meeting