

Effects of cloud turbulence on collision-coalescence in maritime shallow convection

W. W. Grabowski¹,

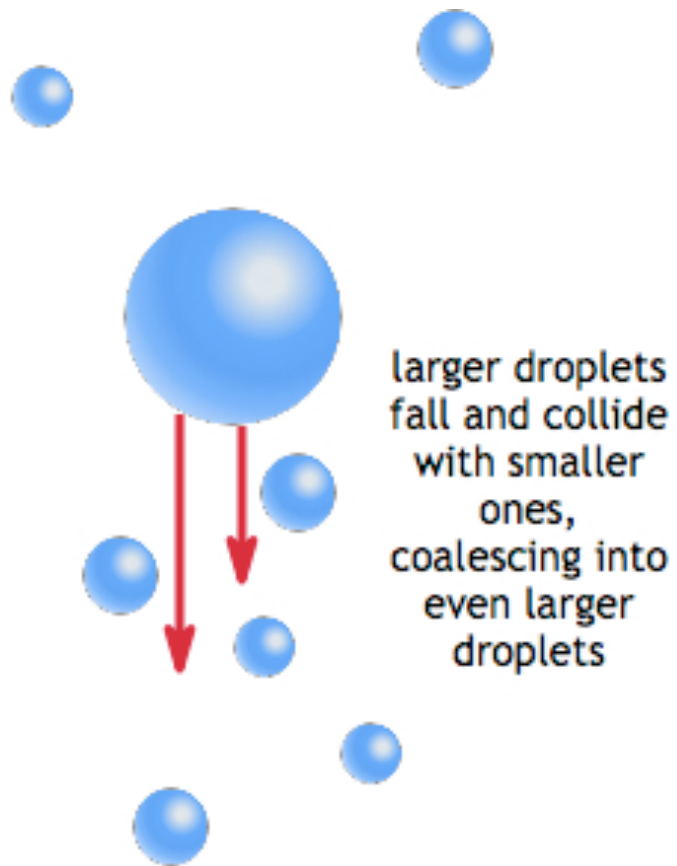
L.-P. Wang², A. Wyszogrodzki¹, and O. Ayala²

¹National Center for Atmospheric Research,
Boulder, Colorado

²University of Delaware, Newark, Delaware



Wyszogrodzki et al., 2013: Turbulent collision-coalescence in maritime shallow convection, *Atmos. Chem. Phys. Discuss.*
Grabowski and Wang, 2013: Growth of cloud droplets in a turbulent environment. *Ann. Rev. Fluid Mech.*



The textbook explanation
of rain formation in ice-
free clouds: gravitational
collision-coalescence...

Time evolution of the droplet spectrum: the Smoluchowski equation:

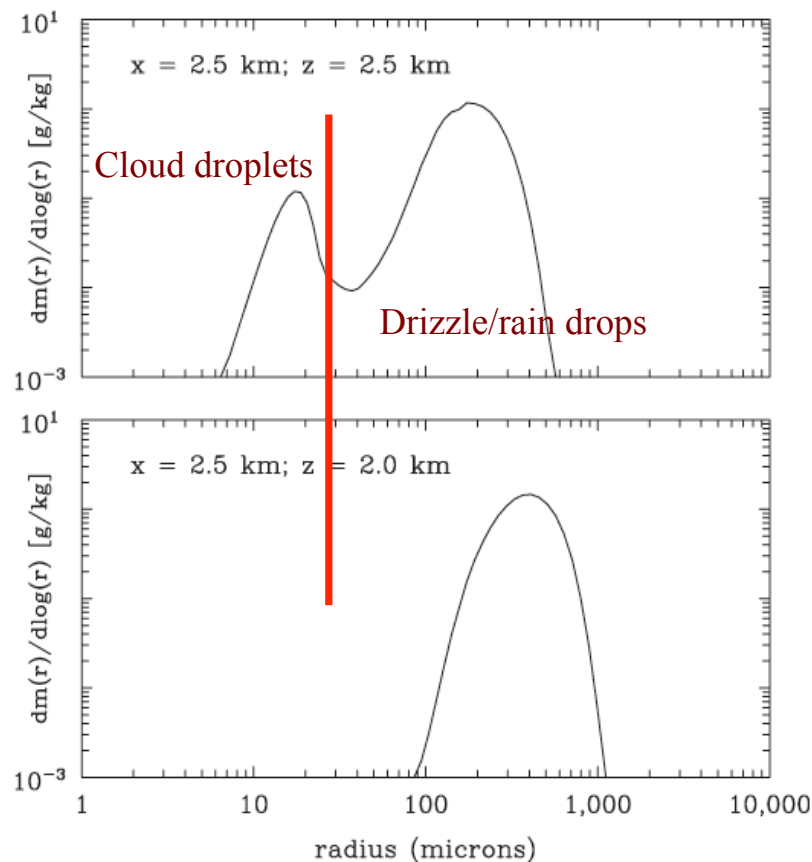
$$\frac{\partial n(x,t)}{\partial t} = \frac{1}{2} \int_0^x K(x-y, y) n(x-y, t) n(y, t) dy - \int_0^\infty K(x, y) n(x, t) n(y, t) dy$$

In a 3D cloud model:

- droplet activation

- growth by condensation of water vapor

- transport in the physical space (advection + sedimentation)



Bin warm-rain microphysics of Grabowski et al. (2011):

- Prediction of the spectral shape of cloud droplets and drizzle/rain; 112 bins;
- Prediction of the supersaturation and thus relating the concentration of activated cloud droplets to local value of the supersaturation; additional variable, concentration of activated CCN, needed.

gravitational kernel: $K_{ij} = \pi (a_i + a_j)^2 |v_i^t - v_j^t| E_{ij}^g$

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generalized kernel (gravity plus turbulence): $K_{ij} = K_{ij}^{tg} E_{ij}^g \eta_E$

Saffman and Turner
(*JFM* 1956)

Wang et al.
(*JAS* 2005)

Grabowski and Wang
(*ARFM* 2013)

$$K_{ij}^{tg} = 2\pi R^2 \langle |w_r(r = R)| \rangle g_{ij}(r = R)$$

$$R = a_i + a_j$$

$$w_r = \mathbf{r} \cdot (\mathbf{V}_i - \mathbf{V}_j) / r \quad \text{radial relative velocity}$$

$$g_{ij} \quad \text{radial distribution function}$$

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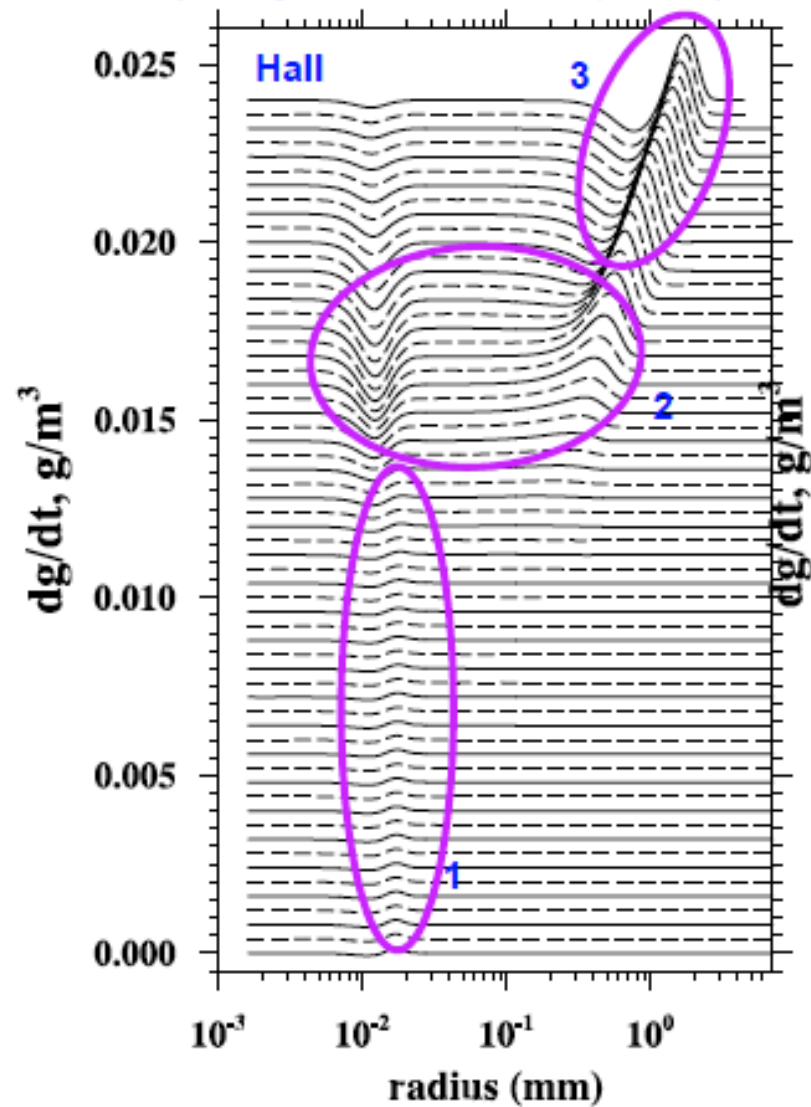
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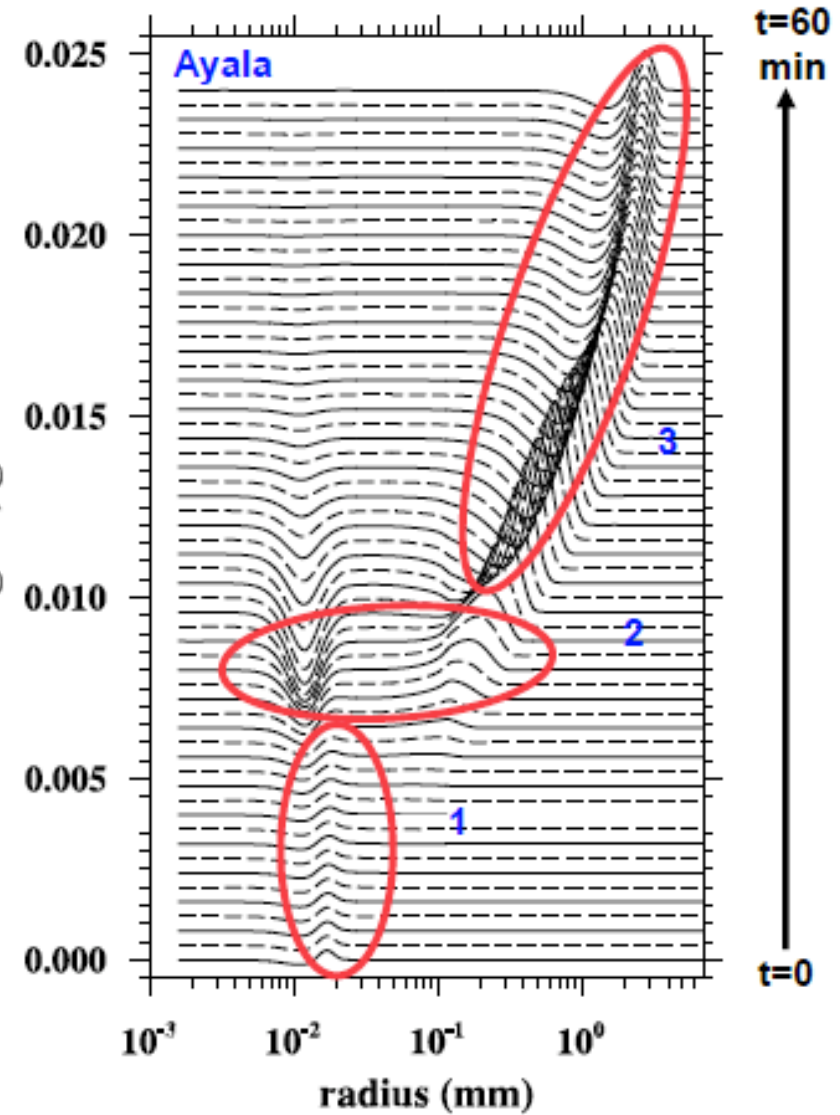
The kernel depends on two parameters describing small-scale turbulence:

- eddy dissipation rate ε
- Taylor microscale Reynolds number Re_λ

1. Autoconversion; 2. Accretion; 3. Hydrometeor self-collection
(Berry and Reinhardt, 1974)



without turbulence



with turbulence, $\varepsilon = 400 \text{ cm}^2\text{s}^{-3}$

A Large Eddy Simulation Intercomparison Study of Shallow Cumulus Convection

JAS
2003

A. PIER SIEBESMA,^a CHRISTOPHER S. BRETHERTON,^b ANDREW BROWN,^c ANDREAS CHLOND,^d JOAN CUXART,^e
PETER G. DUYNKERKE,^{f*} HONGLI JIANG,^g MARAT KHAIROUTDINOV,^h DAVID LEWELLEN,ⁱ CHIN-HOH MOENG,^j
ENRIQUE SANCHEZ,^k BJORN STEVENS,^l AND DAVID E. STEVENS^m

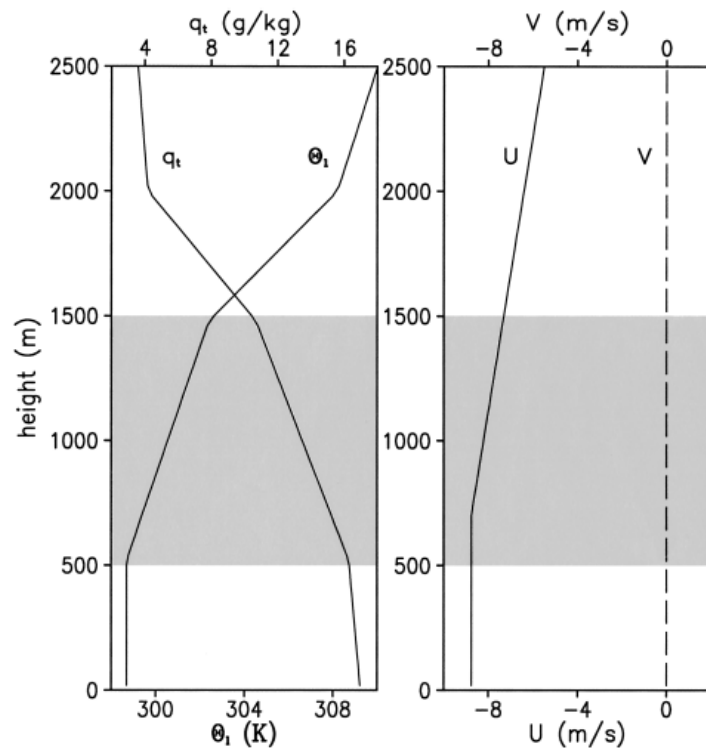
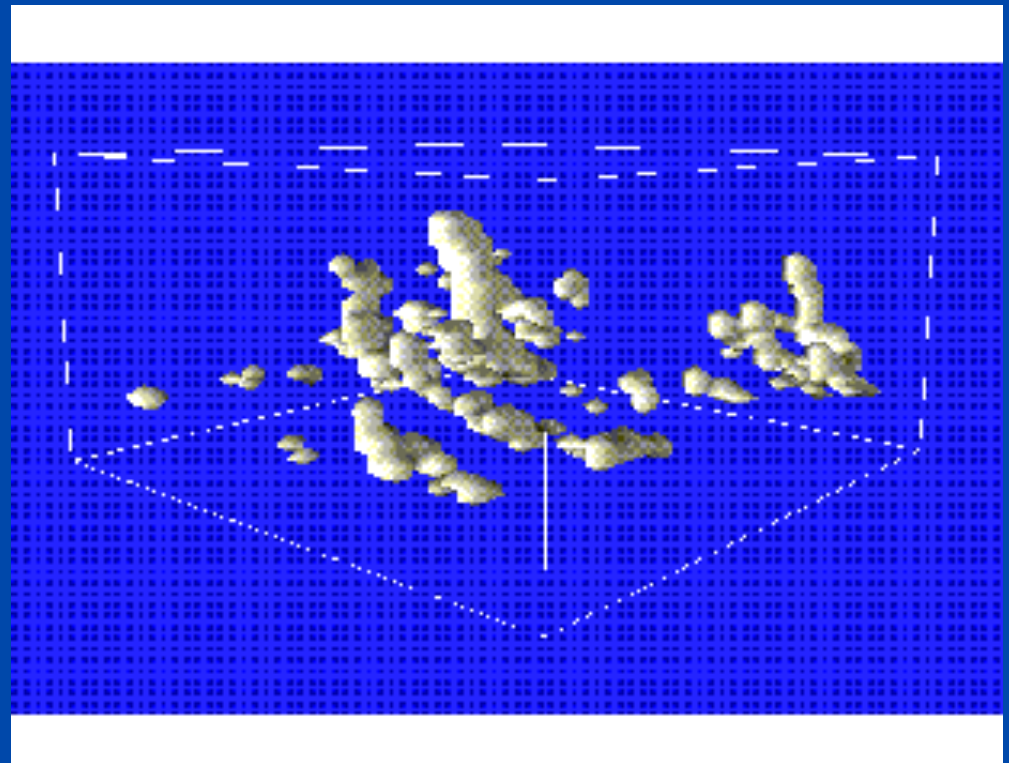


FIG. 1. Initial profiles of the total water specific humidity q_t , the liquid water potential temperature θ_l , and the horizontal wind components u and v . The shaded area denotes the conditionally unstable cloud layer.

$$\Delta x = \Delta y = 50\text{m}; \Delta z = 20\text{m}$$



The Barbados Oceanographic and Meteorological Experiment (BOMEX) case (Holland and Rasmusson 1973)

Turbulent enhancement in LES simulation:

$$\epsilon = C_{\text{eps}}(\text{TKE})^{3/2}/\Delta$$
$$\Delta = (\Delta x + \Delta y + \Delta z)/3$$
$$C_{\text{eps}} = 0.845$$

$$Re_{\lambda} = 15^{1/2}(u_{\text{rms}}/v_K)^2$$

v_K is the Kolmogorov velocity

$$u_{\text{rms}} \text{ (in ms}^{-1}\text{): } u_{\text{rms}} = 2.02 \cdot (\epsilon/400.)^{1/3}$$

$\text{cm}^2 \text{s}^{-3}$

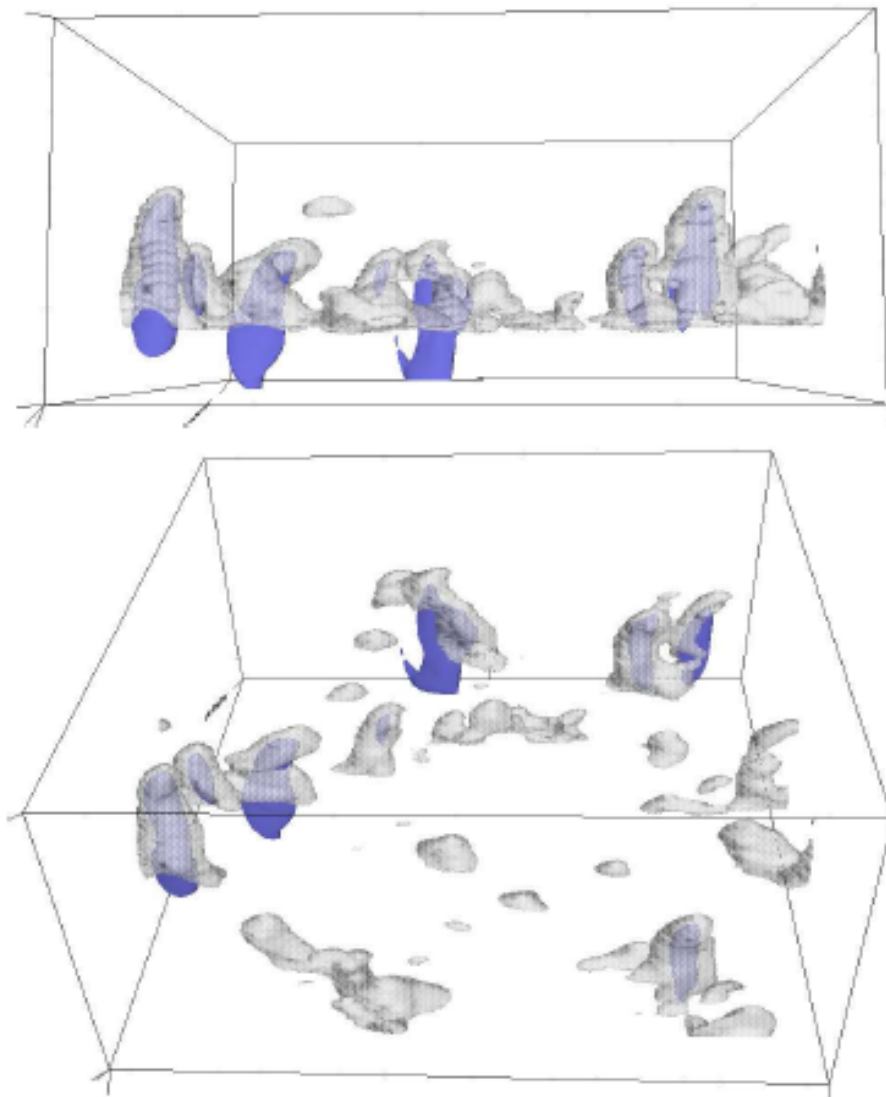
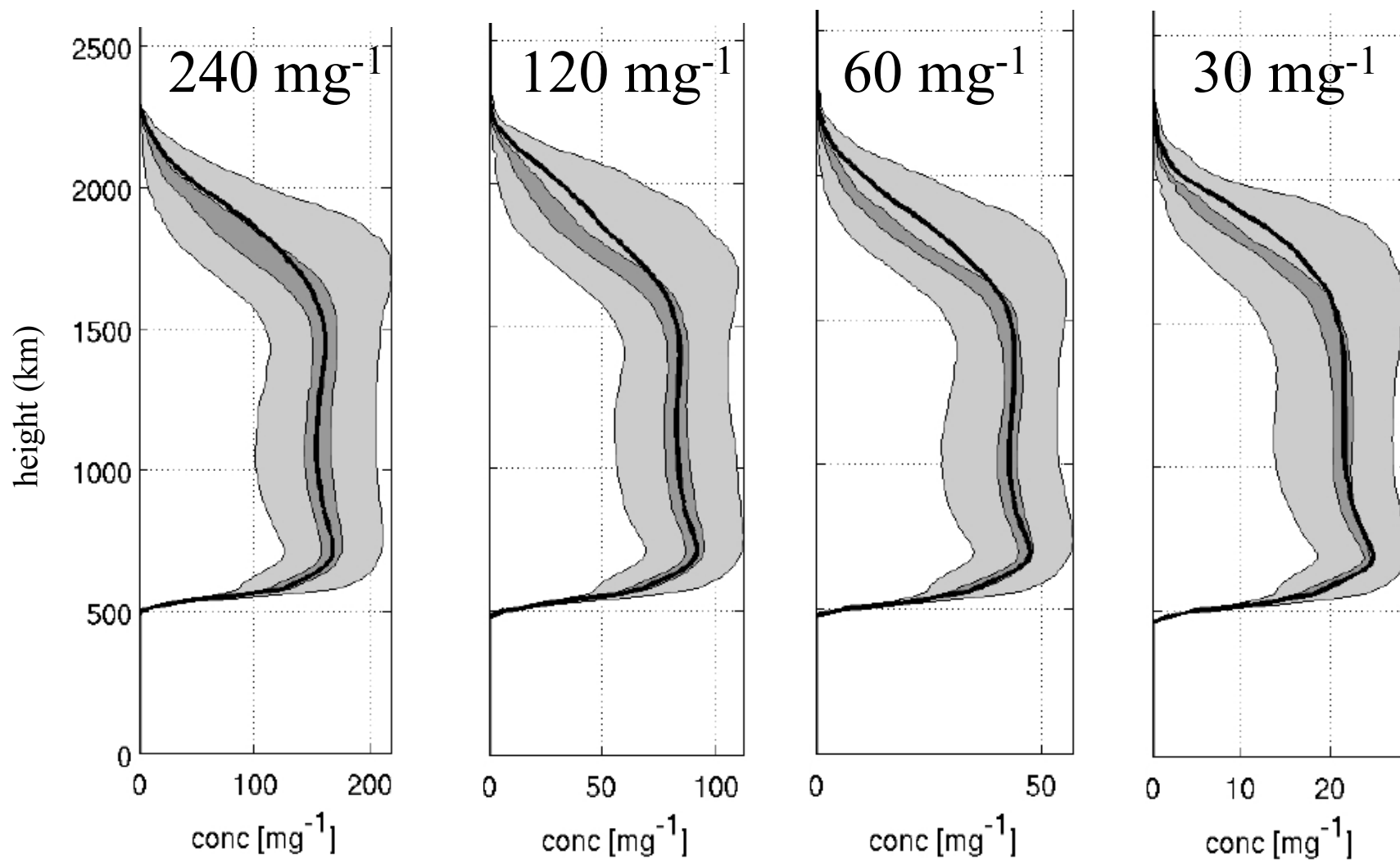
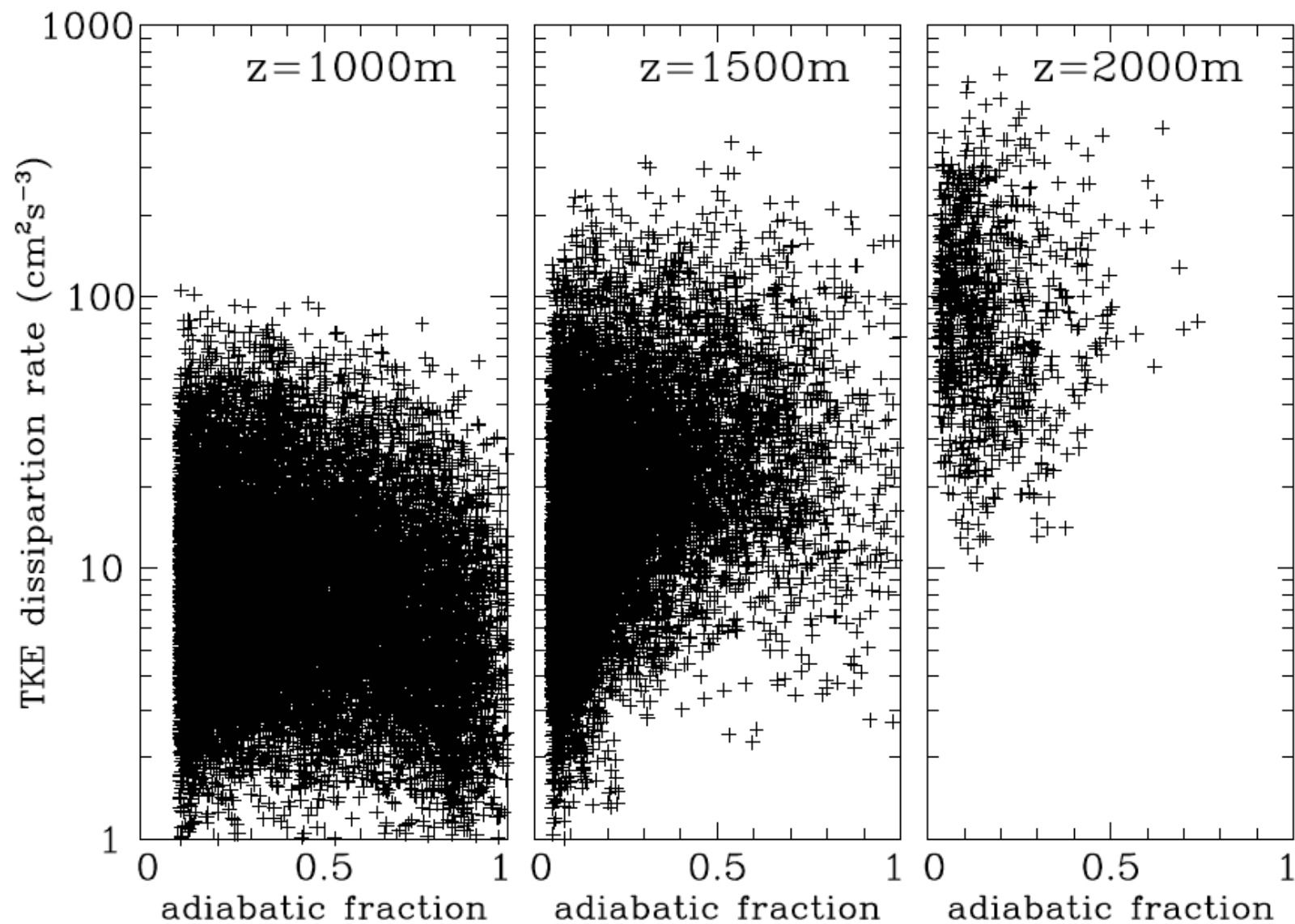


Fig. 4. Snapshots of cloud water mixing ratio (transparent gray) and rain water mixing ratio (solid blue) at the 6th hour of the simulation. The isosurfaces show values $q_c = 0.05 \text{ g kg}^{-1}$ and $q_r = 0.02 \text{ g kg}^{-1}$.

N_{CCN}^0 versus droplet concentration

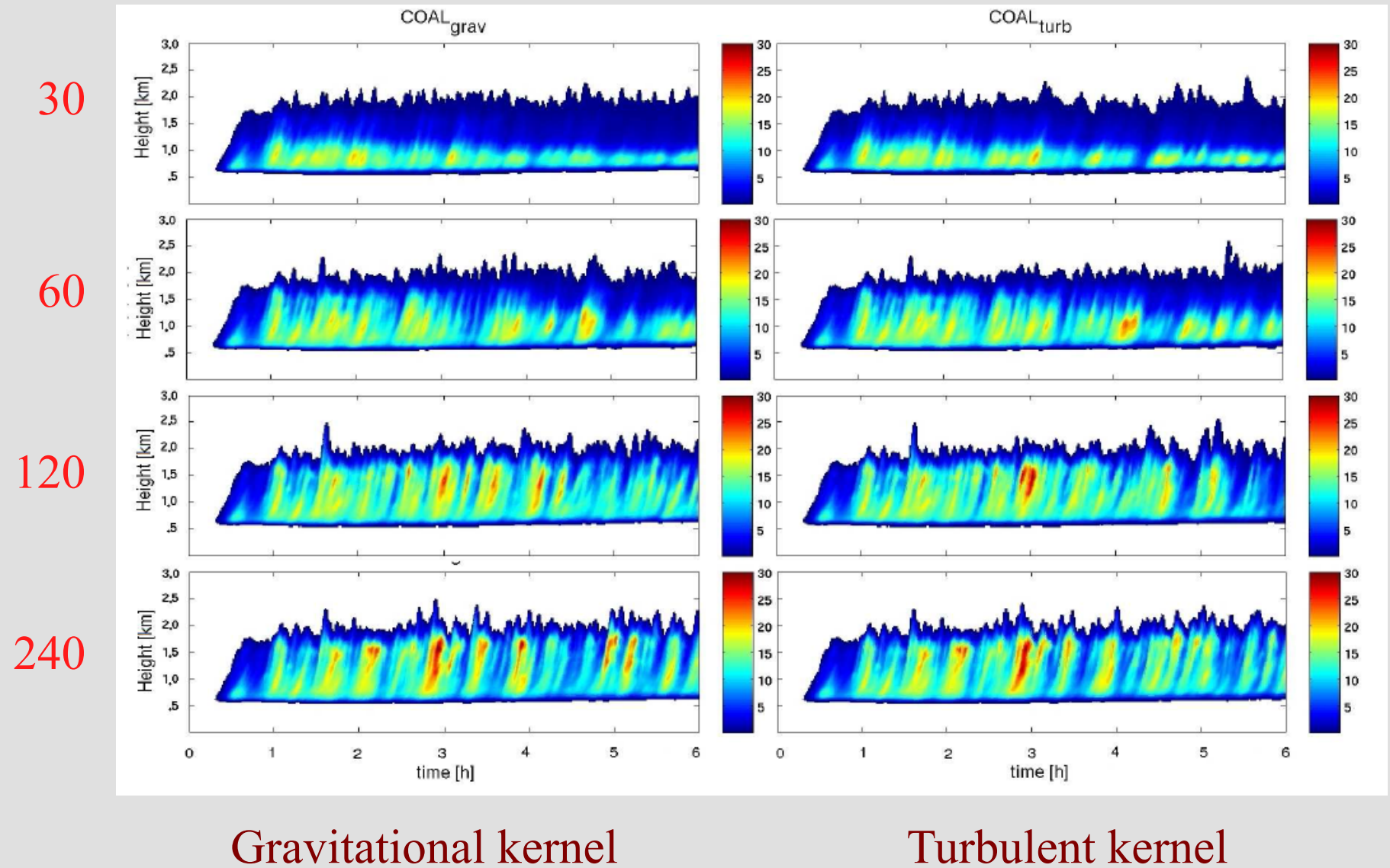


droplet concentration (mg⁻¹)

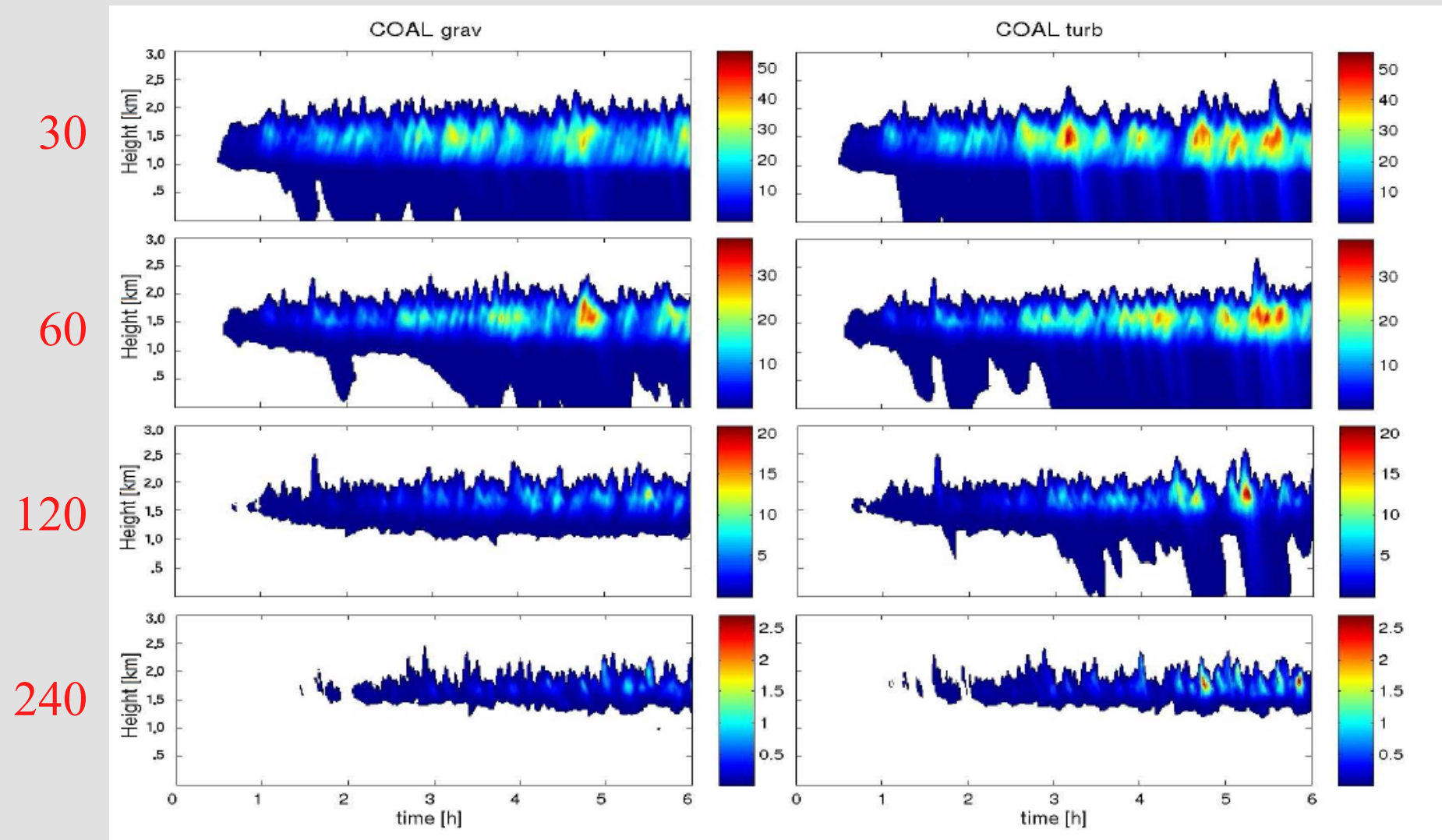


N120 simulation; all cloudy points ($q_c > 0.1\text{g/kg}$) with $\varepsilon > 1\text{cm}^2/\text{s}^3$; hours 3-6

Domain-averaged cloud water mixing ratio ($r < 25 \mu\text{m}$)



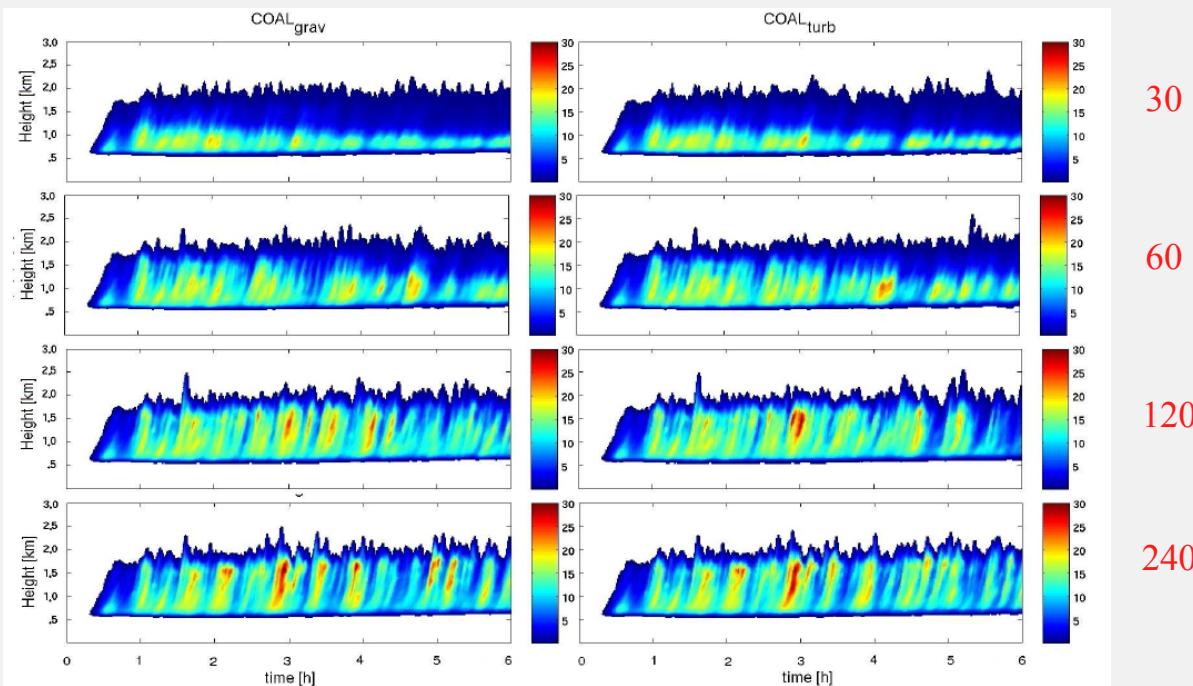
Domain-averaged drizzle/rain water mixing ratio ($r > 25 \mu\text{m}$)



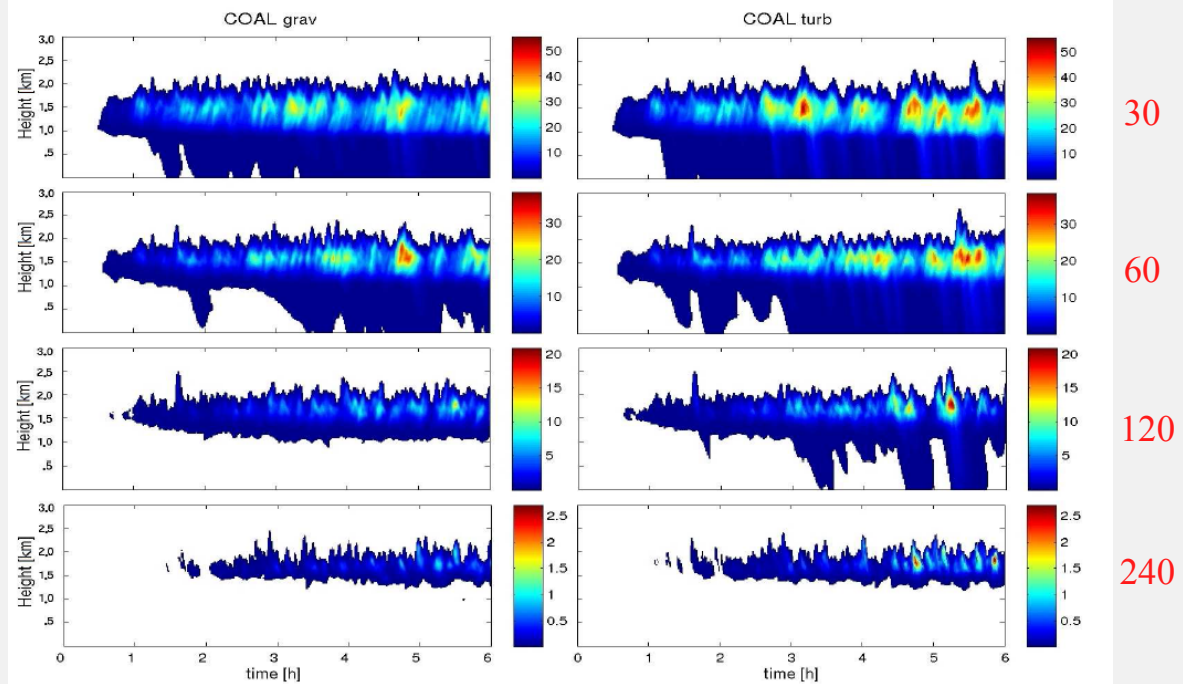
Gravitational kernel

Turbulent kernel

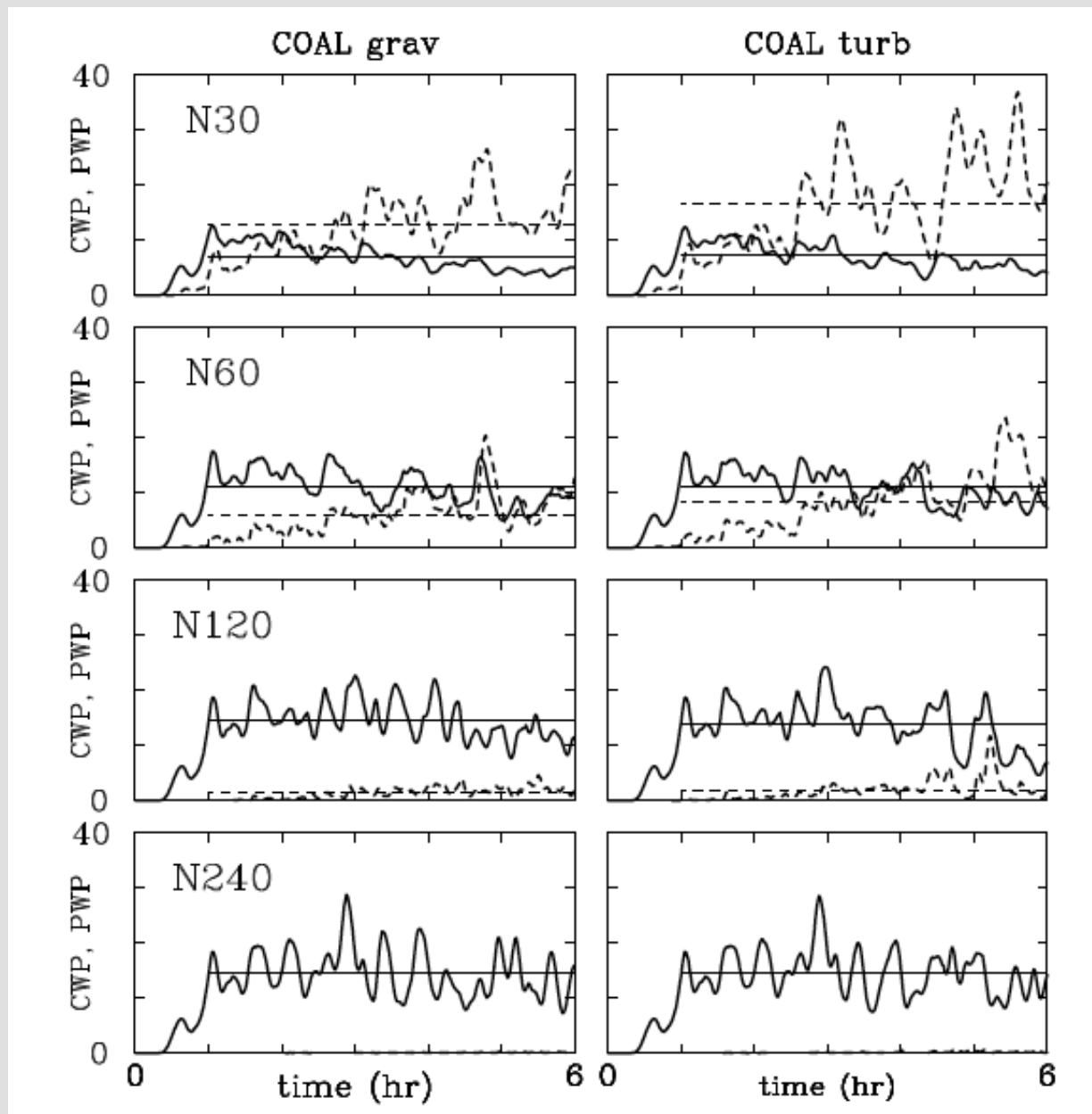
Cloud water



Drizzle/rain
water

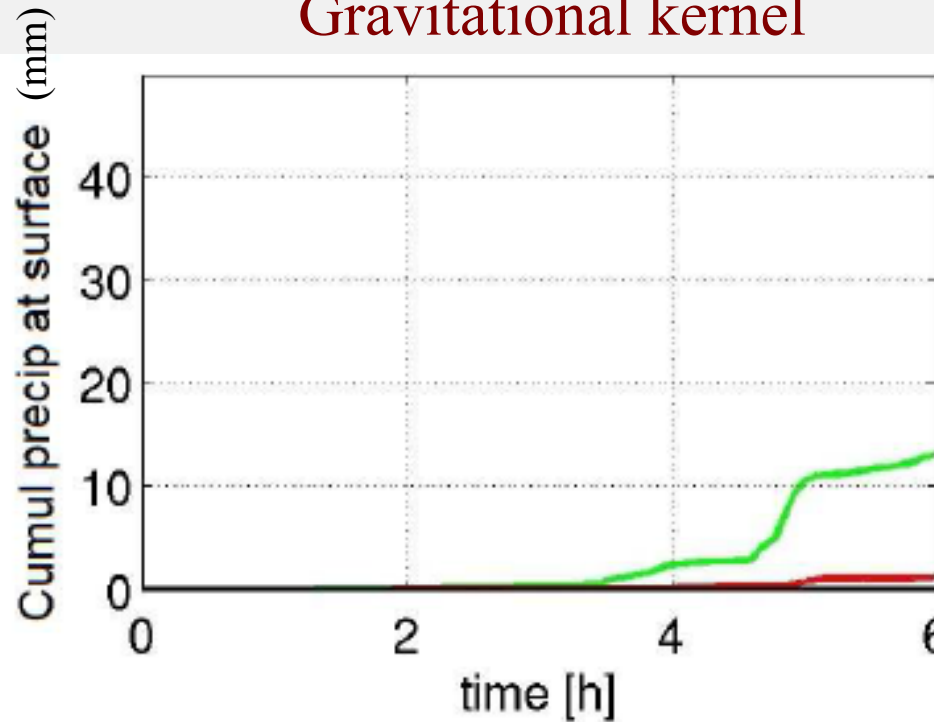


Domain-averaged cloud water path (CWP, solid line) and precip water path (PWP, dashed line)

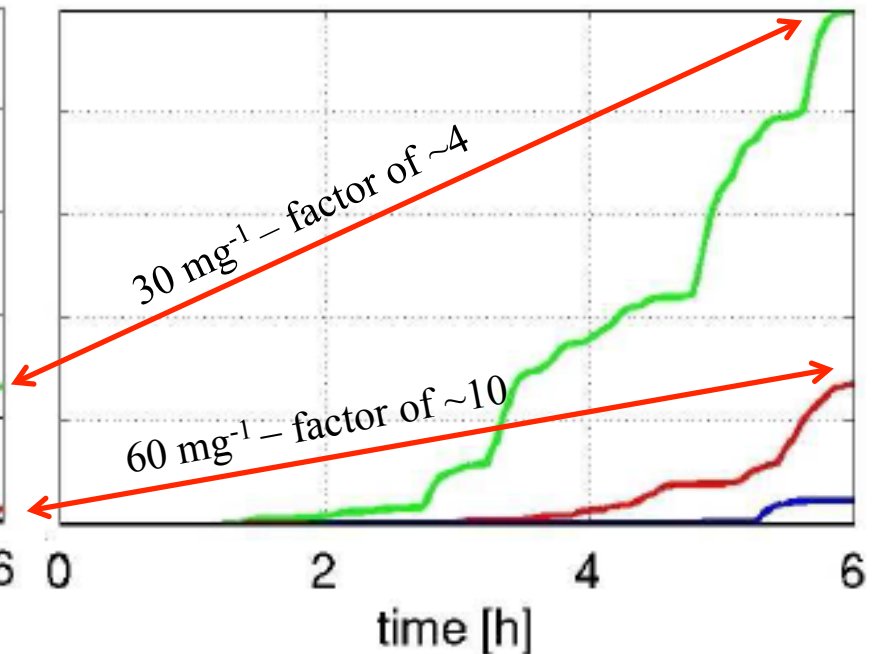


Surface rain accumulation from the cloud field:

Gravitational kernel



Turbulent kernel



Summary:

Small-scale turbulence appears has a significant effect on collisional growth of cloud droplets and development of warm rain in shallow cumuli.

Not only rain tends to form earlier in a single cloud, but also turbulent clouds seem to rain more. This is a combination of microphysical and dynamical effects.

The (perhaps surprising) magnitude of this effect calls for further observational and modeling studies to provide more support for these findings.