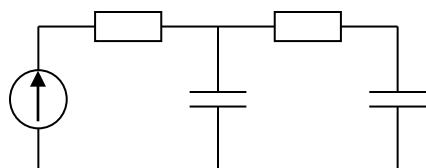


Adjusted radiative forcing, equilibrium climate sensitivity and transient radiative feedback in a two-box energy balance model

Olivier Geoffroy, David Saint-Martin, Aurore Volodire, Gilles Bellon,
Dirk Oliviè, Aurélien Ribes, Sophie Tytéca



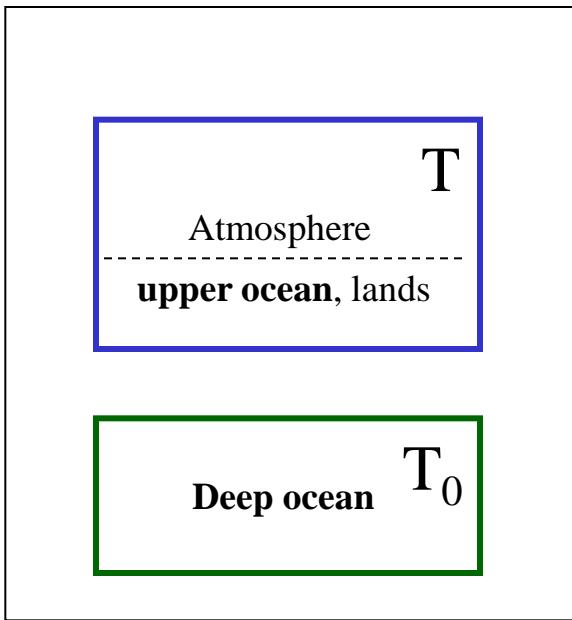
Objective

- Quantify the physical parameters that drive the first-order transient response of a given AOGCM.

Plan

- I) 2-box Energy Balance Model (EBM)
- II) 2-box EBM with efficacy factor of deep ocean heat uptake
- III) Conclusion and perspective

2-box EBM (version with $\varepsilon=1$)



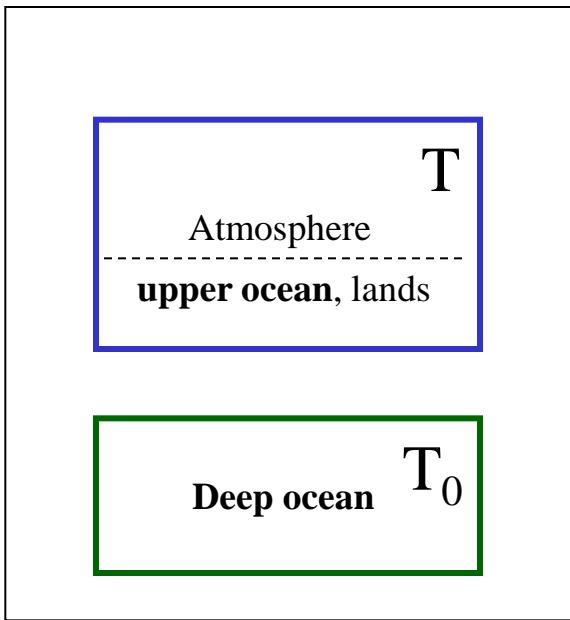
(Held et al, 2010)

T : change in mean surface air temperature

T_0 : change in deep ocean temperature

$$\begin{cases} C \frac{dT}{dt} = N - H \\ C_0 \frac{dT_0}{dt} = H \end{cases}$$

2-box EBM (version with $\varepsilon=1$)



(Held et al, 2010)

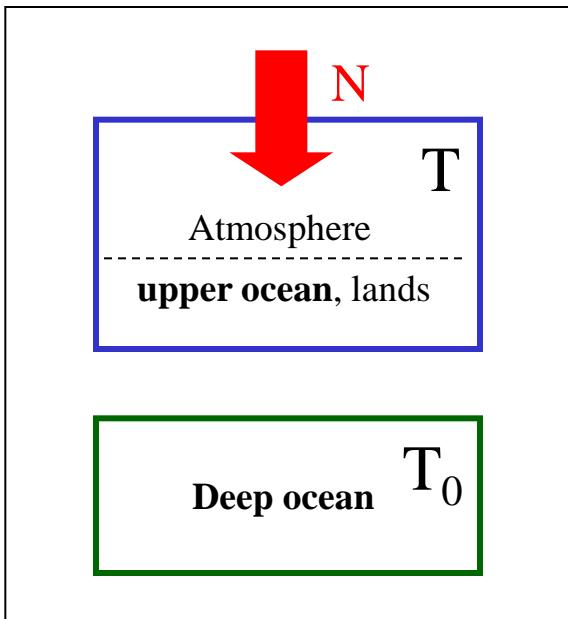
T: change in mean surface air temperature

T₀: change in deep ocean temperature

Change of heat content
of upper ocean

$$\left\{ \begin{array}{l} C \frac{dT}{dt} = N - H \\ C_0 \frac{dT_0}{dt} = H \end{array} \right.$$

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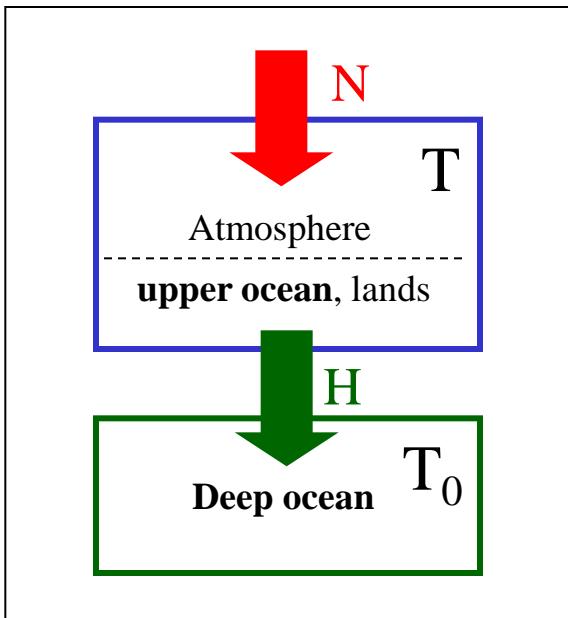
Radiative
imbalance

$$\left\{ \begin{array}{l} C \frac{dT}{dt} = N - H \\ C_0 \frac{dT_0}{dt} = H \end{array} \right.$$

with:

$$N = F - \lambda T$$

2-box EBM (version with $\varepsilon=1$)



(Held et al, 2010)

T : change in mean surface air temperature

T_0 : change in deep ocean temperature

Change of heat content
of upper ocean

Radiative
imbalance

$$\left\{ \begin{array}{l} C \frac{dT}{dt} = N + H \\ C_0 \frac{dT_0}{dt} = H \end{array} \right.$$

Deep ocean
heat uptake

Change of heat
content of deep ocean

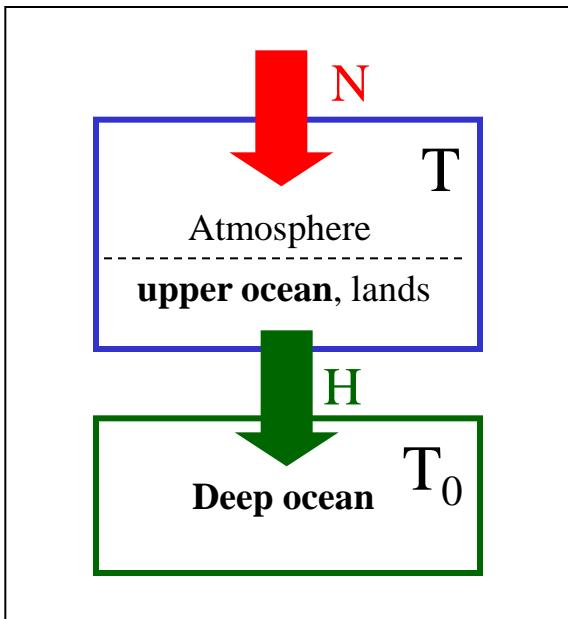
with:

$$N = F - \lambda T$$

$$H = \gamma \cdot (T - T_0)$$

5 parameters: F_{2xCO_2} , λ , C , C_0 , γ

2-box EBM (version with $\varepsilon=1$)



(Held et al, 2010)

T : change in mean surface air temperature

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Change of heat content
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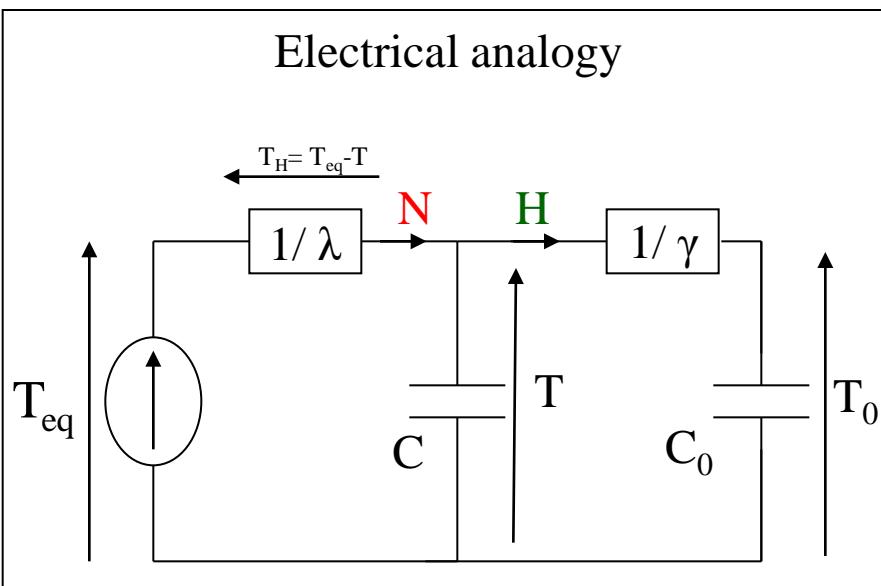
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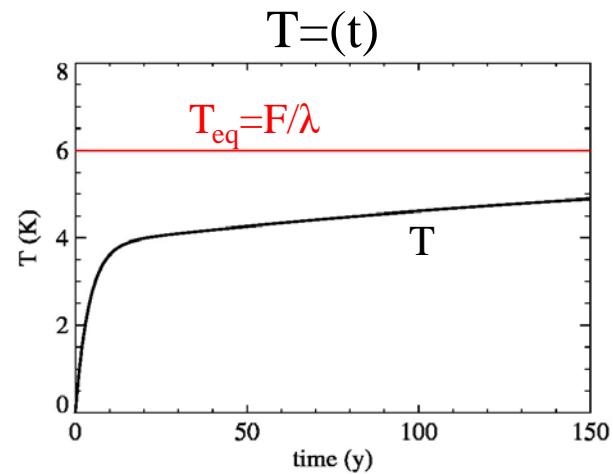


5 parameters: F_{2xCO_2} , λ , C , C_0 , γ

Parameter calibration, method

Analytical solution
(step forcing case)

$$T = \frac{F}{\lambda} - \frac{F}{\lambda} a_f e^{-t/\tau_f} - \frac{F}{\lambda} a_s e^{-t/\tau_s}$$

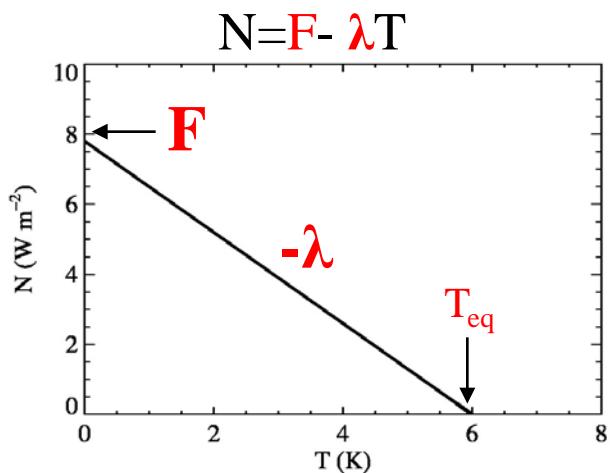


Parameter calibration, method

Analytical solution
(step forcing case)

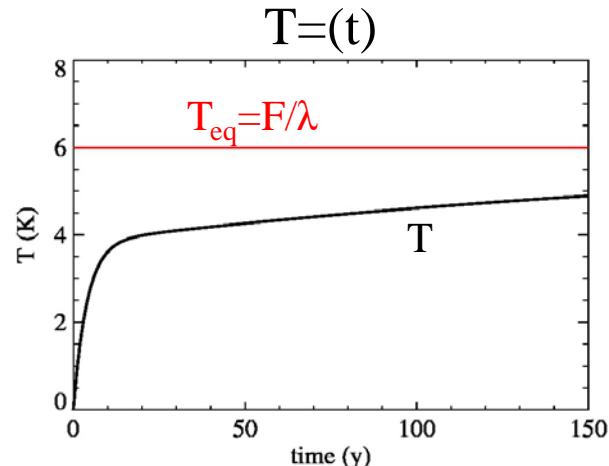
$$T = \frac{F}{\lambda} - \frac{F}{\lambda} a_f e^{-t/\tau_f} - \frac{F}{\lambda} a_s e^{-t/\tau_s}$$

Calibration of $F, \lambda, C, C_0, \gamma$ from AB4CO₂ only



Linear fit (*Gregory et al, 2004*)

↓
 F, λ

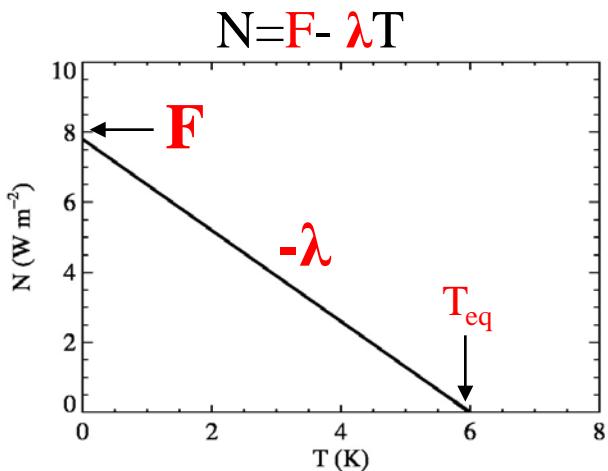


Parameter calibration, method

Analytical solution
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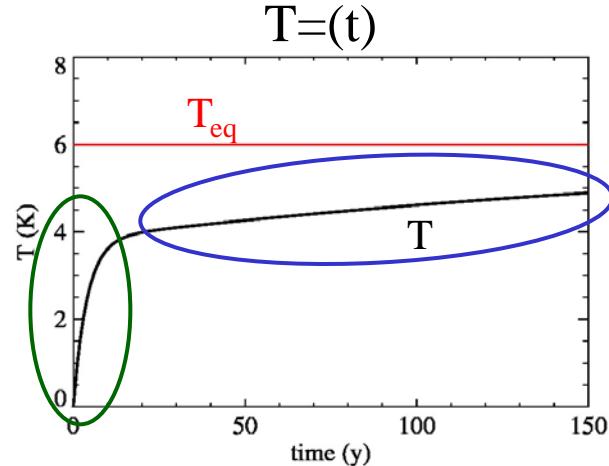
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Linear fit (*Gregory et al, 2004*)

F, λ



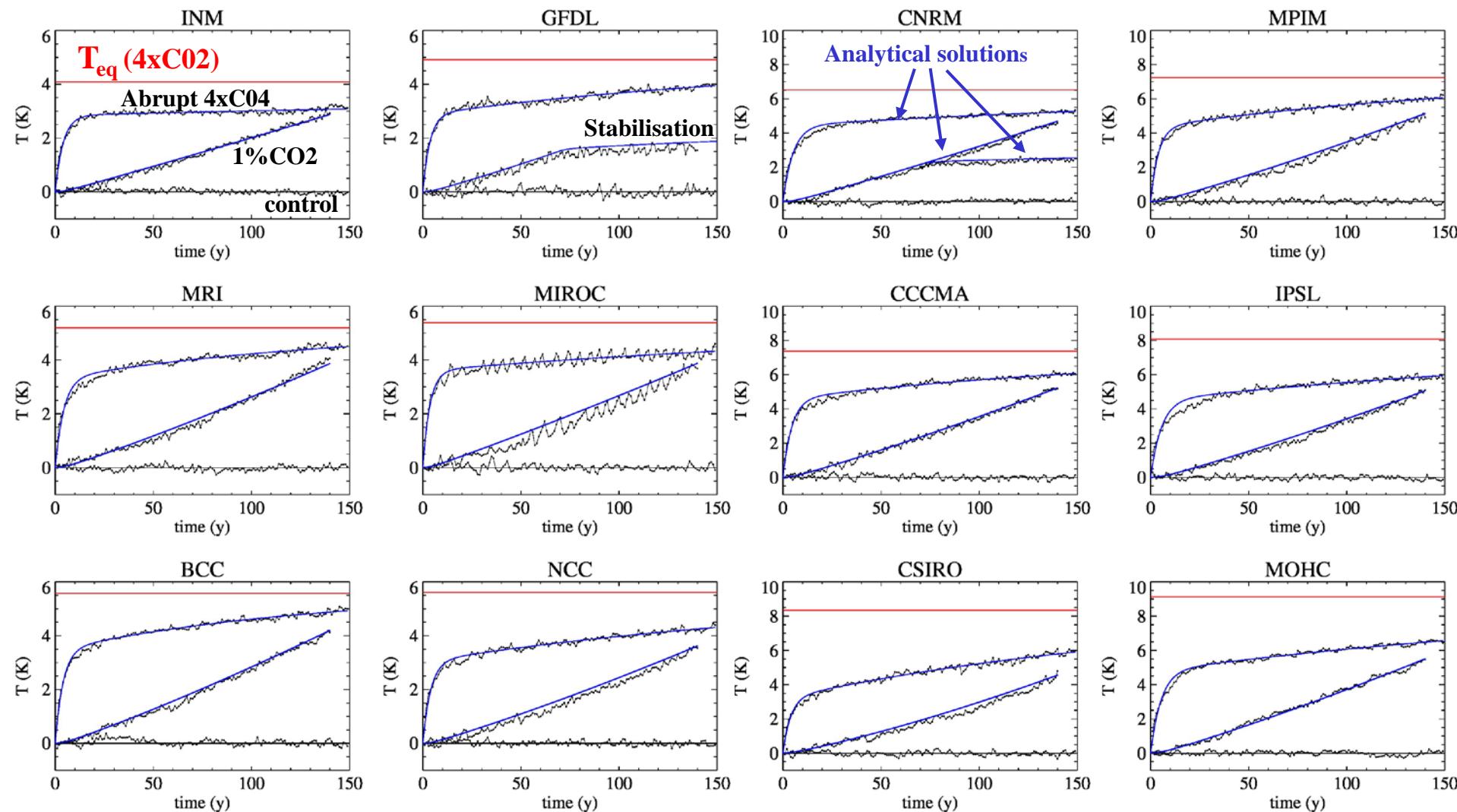
2 Fits (30-150 yr, first 10 yr)

a_f, a_s, τ_s, τ_f

C, C_0, γ

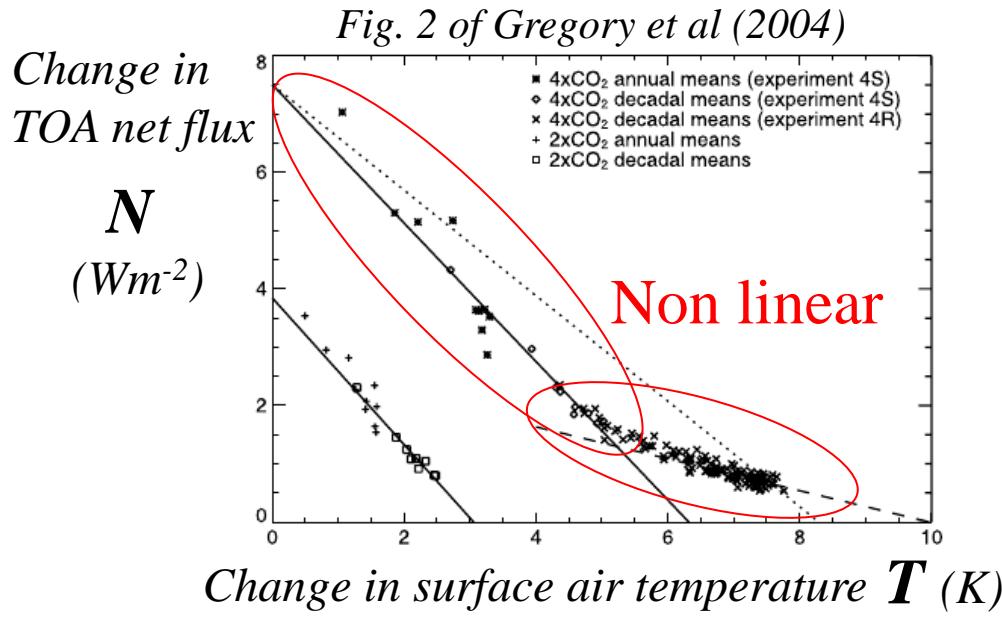
CMIP5 AOGCMs global temperature, results

(Geoffroy et al 2012a, submitted)



- Good representation of T_{AB4CO_2} and T_{UNPCO_2}

$N=f(T)$: limitations of the linear hypothesis.



$$\mathbf{N} = \mathbf{F} - \lambda \mathbf{T}$$

Limited validity

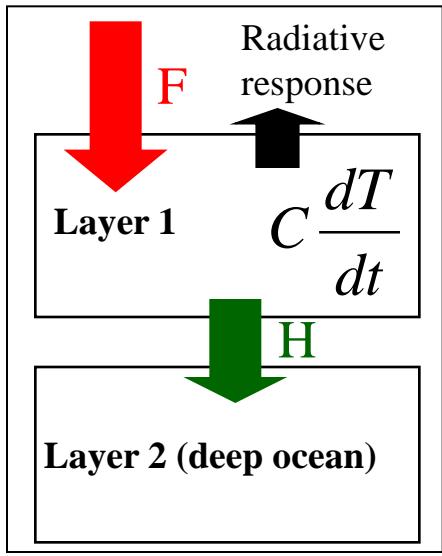
Bias in F and T_{eq} \rightarrow bias in other parameters

2-box EBM with efficacy of deep ocean heat uptake ϵ

(Winton *et al.*, 2010).

(Held *et al.*, 2010)

Underlying assumptions



3 « forcings » → Decomposition of the radiative response in
3 radiative responses:

$$F - \lambda T_{eq} = 0$$

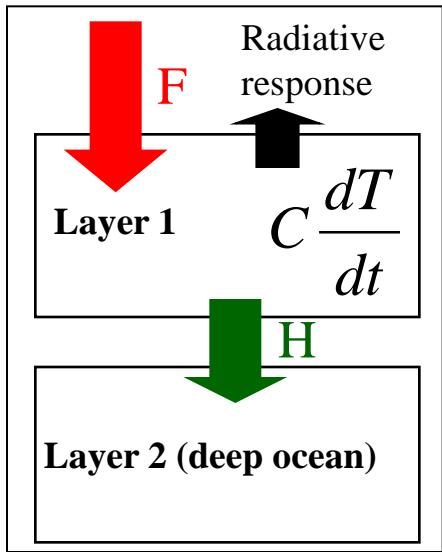
$$- C \frac{dT}{dt} - \lambda T_U = 0$$

Assumption of additivity

$$- H - \lambda_D T_D = 0$$

$$T = T_{eq} + T_U + T_D$$

Underlying assumptions



3 « forcings » → Decomposition of the radiative response in
3 radiative responses:

$$F - \lambda T_{eq} = 0$$

$$- C \frac{dT}{dt} - \lambda T_U = 0$$

$$- H - \lambda_D T_D = 0$$

Assumption of additivity

$$T = T_{eq} + T_U + T_D$$

- the pattern of the temperature response associated with deep ocean heat uptake is different from the equilibrium pattern
- the strength of the feedbacks varies geographically

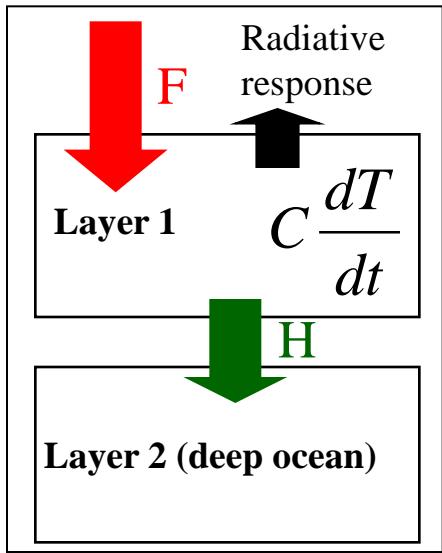


T_D associated with a different feedback parameter $\lambda_D = \frac{\lambda}{\varepsilon} \neq \lambda$

(Winton et al., 2010).

Efficacy (Hansen et al, 2005)

Underlying assumptions



3 « forcings » → Decomposition of the radiative response in
3 radiative responses:

$$F - \lambda T_{eq} = 0$$

$$-C \frac{dT}{dt} - \lambda T_U = 0$$

$$-H - \frac{\lambda}{\varepsilon} T_D = 0$$

Assumption of additivity

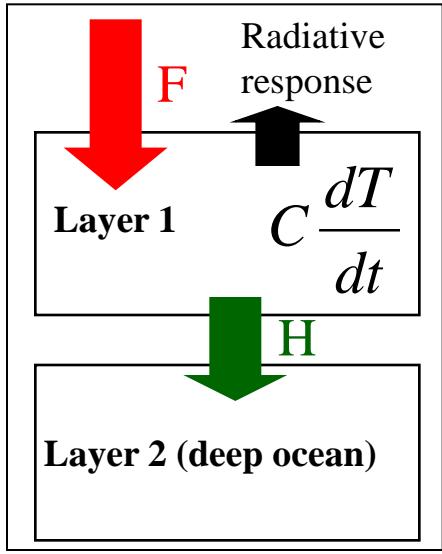
$$T = T_{eq} + T_U + T_D$$

$$C \frac{dT}{dt} = \underbrace{F - \lambda T - (\varepsilon - 1)H - H}_{N}$$
$$C \frac{dT_0}{dt} = H$$

2-box EBM
with ε

(Held et al, 2010)

Underlying assumptions



3 « forcings » → Decomposition of the radiative response in 3 radiative responses:

$$F - \lambda T_{eq} = 0$$

$$-C \frac{dT}{dt} - \lambda T_U = 0$$

$$-H - \frac{\lambda}{\varepsilon} T_D = 0$$

Assumption of additivity

$$T = T_{eq} + T_U + T_D$$

$$N = F - \lambda_t T$$

↑

Time-dependant:

$$\lambda_t = \lambda - (\varepsilon - 1) \frac{H}{T}$$

\downarrow

$$C \frac{dT}{dt} = \underbrace{F - \lambda T - (\varepsilon - 1)H - H}_{N}$$

$$C \frac{dT_0}{dt} = H$$

2-box EBM
with ε

(Held et al, 2010)

Parameter calibration

6 parameters: F , λ , ε , C , C_0 , γ

Iterative method

$$N = F - \lambda T - (\varepsilon - 1)H$$

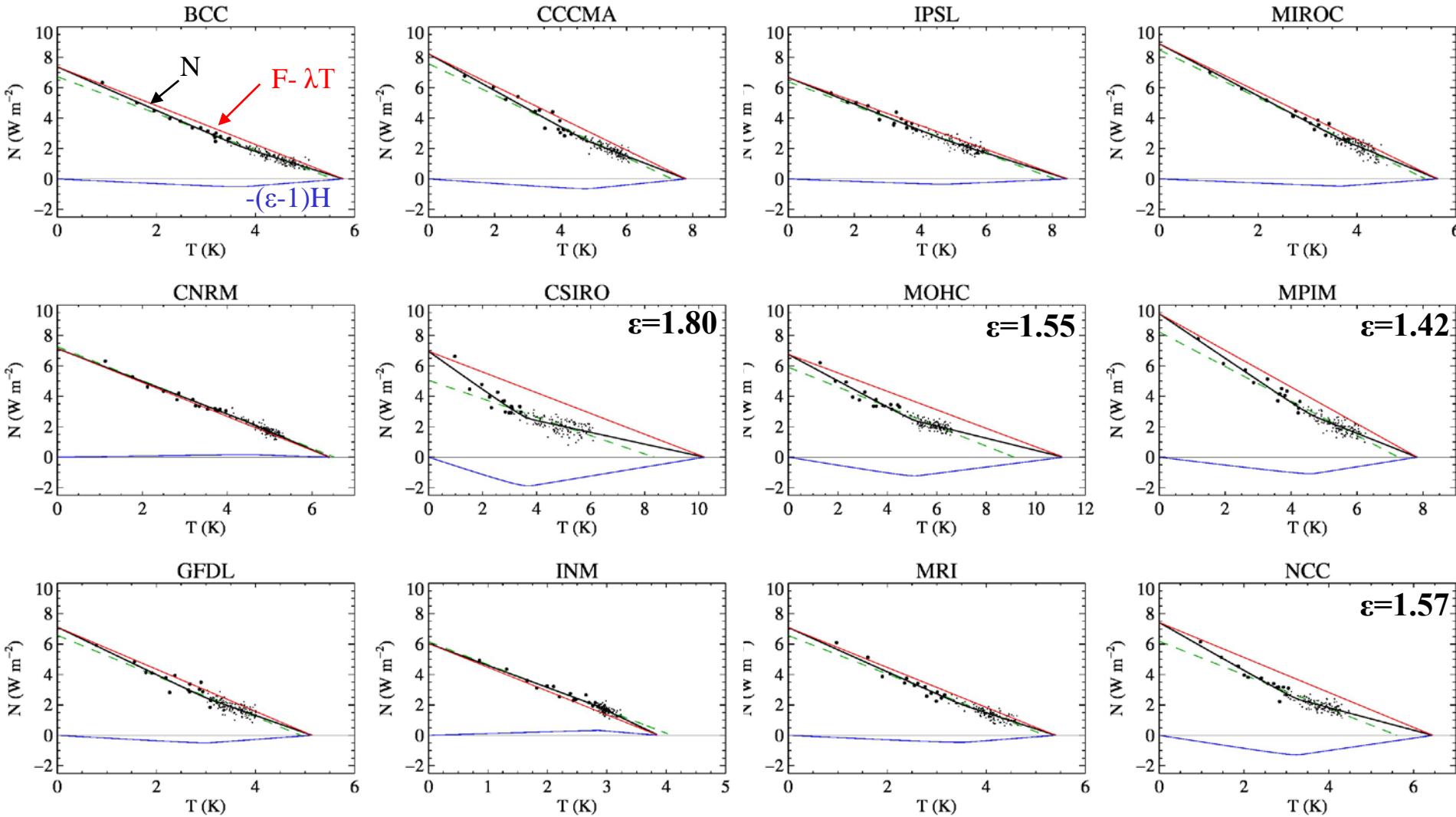
Multiple regression $\rightarrow F_{\text{4xCO}_2}$, λ and ε

Initial value of H : solutions of the EBM with $\varepsilon=1$

$$N = F - \lambda T - (\varepsilon - 1)H$$

Results, N=f(T)

(Geoffroy et al 2012b, submitted)



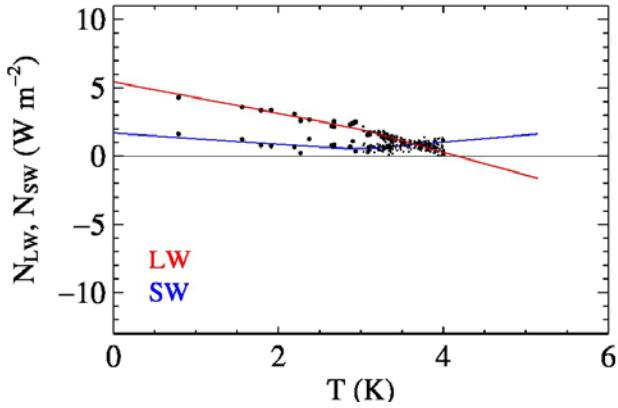
- N-T evolution well represented for all AOGCMs → Better estimation of ECS, F, γ ...
- On average, $T_{eq4xCO_2} \uparrow$ of 0.5 K (max: ~2K), $F_{4xCO_2} \uparrow$ of 0.6 Wm⁻² (max: ~ 2 Wm⁻²)

Decomposition of N

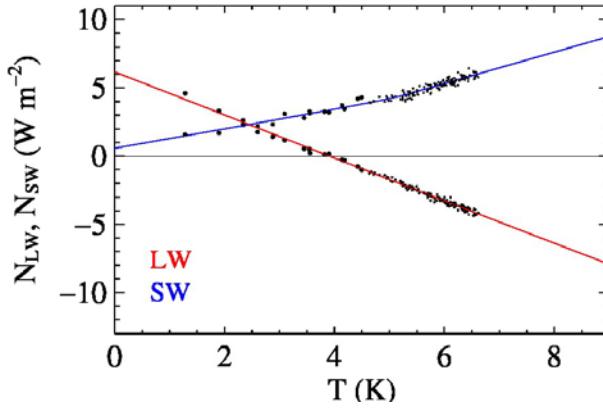
$$N^i = F^i - \lambda^i T - (\lambda^i - \lambda_D^i) \frac{\epsilon}{\lambda} H$$

$i = [\text{LW}, \text{SW}]$.

GFDL



MOHC



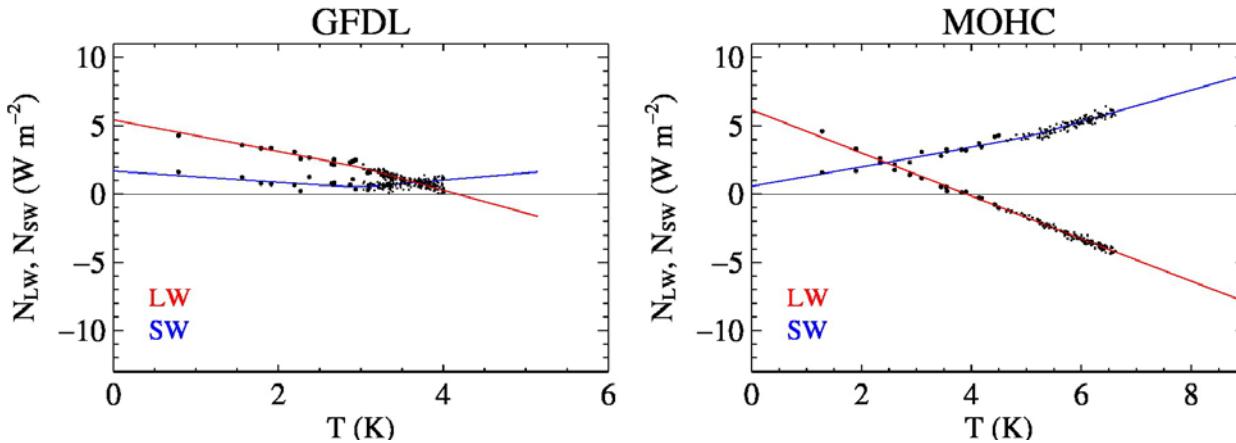
Difference with the linear fit:

	GFDL	MOHC
$F^{\text{LW}} (\text{W m}^{-2})$	-0.68	0.05
$F^{\text{SW}} (\text{W m}^{-2})$	+1.21	0.78

Decomposition of N

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	GFDL	MOHC
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PRP in transition (*Colman and Mc Avaney, 2011*)

- PRP over the **first 15 years** and the **last 10 years** of CNRM-CM5 AB4CO₂
- + **linear fit** (ϵ close to 1 for CNRM-CM5) of each partial radiative flux:

	CO ₂	T	W	A	C	O ₃	Sum	$N=f(T)$
$F (\text{W m}^{-2})$	8.6	-0.73	-0.55	-0.15	0.64	-0.44	7.4	7.4
$\lambda (\text{W m}^{-2} \text{K}^{-1})$	0.04	-3.85	2.11	0.45	0.17	0.00	-1.08	-1.10

Contribution of each component to F and λ

The $N=f(T)$ fit method can be extended to each component of a N decomposition such as LW/SW, cloudy/clear sky fluxes and **partial radiative fluxes (PRP, kernel)**.

Contributors to the spread of the multimodel T response

Objective: use the EBM framework to quantify the contribution of each parameter ($F_{4\text{CO}_2}$, λ , ε , C , C_0 , γ) to the spread of the CMIP5 AOGCMs temperature response (1% CO₂ yr⁻¹ experiment)

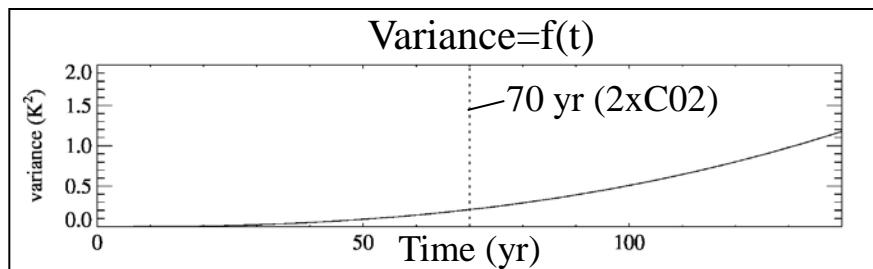
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Method: analysis of variance.

$$f(F_{4\text{CO}_2}^i, \lambda^j, \varepsilon^k, C^l, C_0^m, \gamma^n) = f_0 + f_1(F_{4\text{CO}_2}^i) + f_2(\lambda^j) + f_3(\varepsilon^k) + f_4(C^l) + f_5(C_0^m) + f_6(\gamma^n) + I$$

Results



Geoffroy et al (2012c, to be submitted)

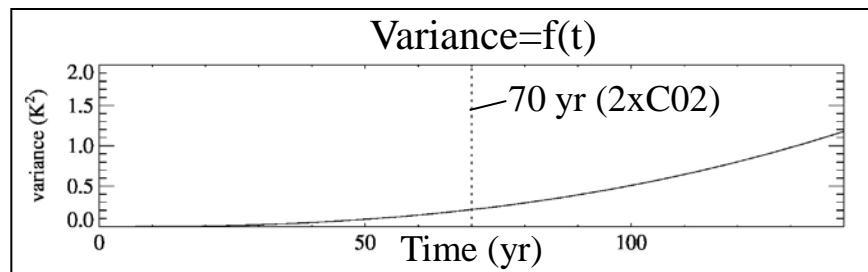
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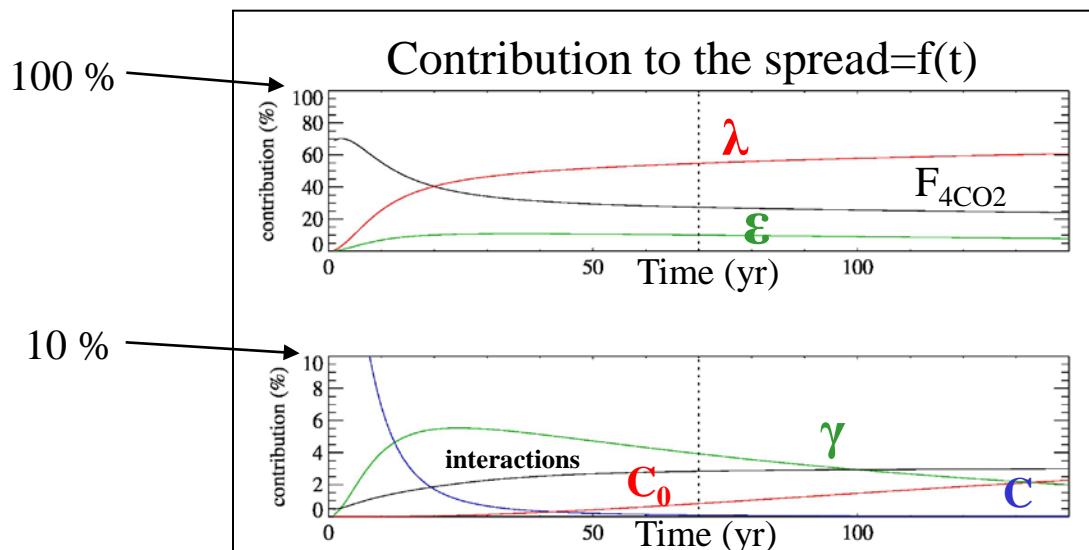
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Results



Geoffroy et al (2012c, to be submitted)



Contribution to the TCR spread

λ: 55 %

F: 27 %

ε: 10 %

γ: 4 %

C₀: 0.8 %

C: 0.1 %

Inter: 2.8%

Radiative parameters:
92% of the spread

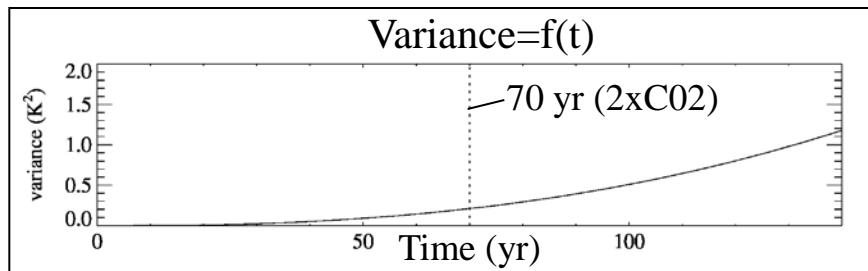
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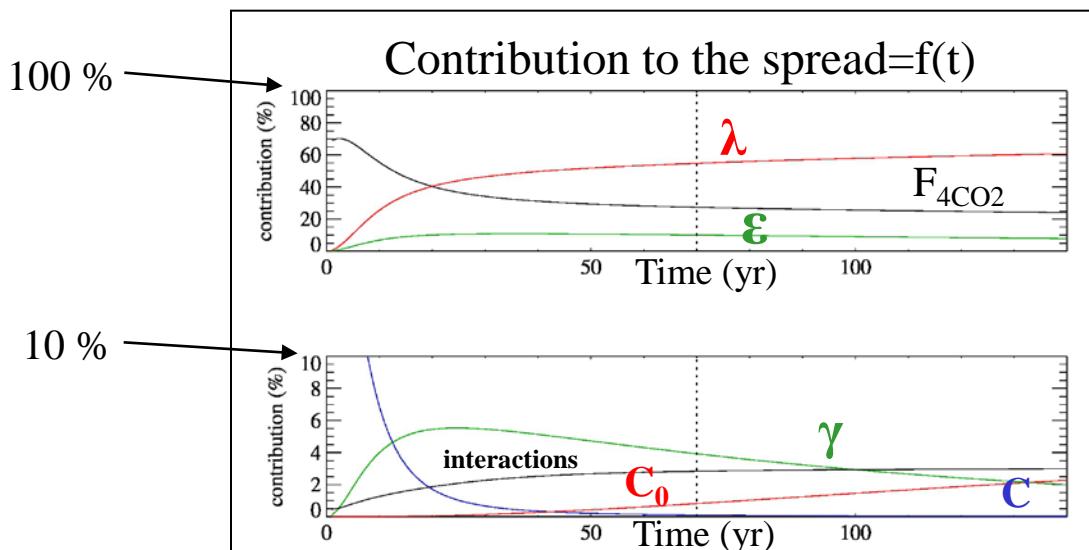
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Results



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F : 27 %

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Radiative parameters:
92% of the spread

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C : 0.1 %

Inter: 2.8%

ECS

λ : 87 %

F : 12 %

Inter: 1 %

Conclusion and perspectives

- Analytical solution of the 2-box EBM → **Physically-based method** for calibration
- **Good representation of the temperature response** for AB4CO₂ and 1% CO₂ yr⁻¹.
- **Efficacy of deep ocean heat uptake** allows to take in account the time-evolution of the feedback strength in transient regime due to the impact of the deep ocean heat uptake on the temperature pattern and the spatial heterogeneity of the strength of the feedbacks.
 - **Good representation of the evolution of the radiative flux imbalance as a function of the temperature response** for all AOGCMs.
 - **Determination of an AOGCM first-order parameters** (link with other properties e.g. MLD?)
- This framework combined with an analysis of variance method has been used to quantify the **contribution of the parameters to the spread of CMIP5 AOGCM response** (1% CO₂ yr⁻¹ runs). The radiative parameters are the **main contributors to the spread: equilibrium feedback parameter (55%), forcing (27 %), efficacy of deep ocean heat uptake (10%)**.
- The calibration method can be extended to **partial radiative flux** (PRP or kernels method) and then compute the contribution of each component (temperature, water vapor, surface albedo, clouds) to the fast adjustment of the radiative forcing, to the equilibrium feedback parameter and to the deep ocean heat uptake feedback parameter.

-
- Geoffroy O., D. Saint-Martin, D. J. L. Oliviè, A. Volodire, G. Bellon, S. Tytéca: « *Transient climate response in a two-box energy-balance model. Part I: analytical solution and parameter calibration using CMIP5 AOGCM experiments* », submitted to Journal of Climate.
 - Geoffroy O., D. Saint-Martin, G. Bellon, A. Volodire, D. J. L. Oliviè, S. Tytéca: « *Transient climate response in a two-box energy-balance model. Part II: representation of the efficacy of deep-ocean heat uptake and validation for CMIP5 AOGCMs* », submitted to Journal of Climate.
 - Geoffroy O., D. Saint-Martin, A. Ribes: « *Contribution of global thermal properties intermodel differences to the spread of CMIP5 AOGCMs transient response* », to be submitted to GRL.

Parameters, T_{eq} and time scales

2-box EBM with ε free parameter

Model	$\mathcal{F}_{4x\text{CO}_2}$ (W m ⁻²)	λ (W m ⁻² K ⁻¹)	ε	$T_{4x\text{CO}_2}$ (K)	C (W y m ⁻² K ⁻¹)	C_0 (W y m ⁻² K ⁻¹)	γ (W m ⁻² K ⁻¹)	τ_f (y)	τ_s (y)
BCC (BCC-CSM1-1)	7.4	1.28	1.27	5.8	8.4	56	0.59	4.1	152
CCCMA (CanESM2)	8.2	1.06	1.28	7.8	8.0	77	0.54	4.5	139
CNRM (CNRM-CM5.1)	7.1	1.12	0.92	6.4	8.3	95	0.51	5.2	266
CSIRO (CSIRO-Mk3-6-0)	7.0	0.68	1.82	10.2	8.5	76	0.71	4.2	316
GFDL (GFDL-ESM2M)	7.1	1.38	1.21	5.1	8.8	112	0.85	3.6	233
INM (INMCM4)	6.0	1.56	0.83	3.9	8.5	271	0.67	4.0	546
IPSL (IPSL-CM5A-LR)	6.7	0.79	1.14	8.5	8.1	100	0.57	5.5	327
MIROC (MIROC5)	8.9	1.58	1.19	5.6	8.7	158	0.73	3.6	338
MOHC (HadGEM2-ES)	6.8	0.61	1.54	11.1	7.5	98	0.49	5.4	457
MPIM (MPI-ESM-LR)	9.4	1.21	1.42	7.8	8.5	78	0.62	4.0	220
MRI (MRI-CGCM3)	7.1	1.31	1.25	5.4	9.3	68	0.59	4.4	181
NCC (NorESM1-M)	7.4	1.15	1.57	6.5	9.7	121	0.76	4.1	328
Multimodel mean	7.4	1.14	1.29	7.0	8.5	109	0.64	4.4	300
Standard deviation	1.0	0.32	0.27	2.1	0.6	58	0.11	0.7	113

Radiative parameters and T_{eq}

2 box EBM
 $\epsilon=1$

Model	\mathcal{F}_{4xCO_2} (W m ⁻²)	λ (W m ⁻² K ⁻¹)	T_{4xCO_2} (K)
BCC (BCC-CSM1-1)	6.7	1.21	5.6
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MRI (MRI-CGCM3)	6.6	1.26	5.2
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Multimodel mean	6.8	1.11	6.5
Standard deviation	1.0	0.31	1.6

2 box EBM
 ϵ free
parameter

Model	\mathcal{F}_{4xCO_2} (W m ⁻²)	λ (W m ⁻² K ⁻¹)	ϵ	T_{4xCO_2} (K)
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INM (INMCM4)	6.0	1.56	0.83	3.9
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Multimodel mean	7.4	1.14	1.29	7.0
Standard deviation	1.0	0.32	0.27	2.1

Thermal inertia parameters and time scales

2 box EBM
 $\epsilon=1$

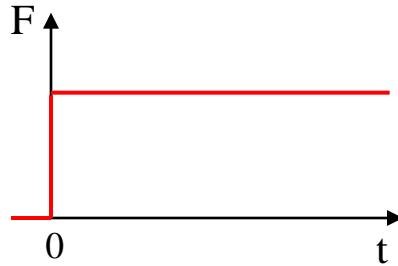
Model	C (W y m $^{-2}$ K $^{-1}$)	C_0 (W y m $^{-2}$ K $^{-1}$)	γ (W m $^{-2}$ K $^{-1}$)	τ_f (y)	τ_s (y)
BCC (BCC-CSM1-1)	7.6	53	0.67	4.0	126
CCCMA (CanESM2)	7.3	71	0.59	4.5	193
CNRM (CNRM-CM5.1)	8.4	99	0.50	5.2	289
CSIRO (CSIRO-Mk3-6-0)	6.0	69	0.88	3.9	200
GFDL (GFDL-ESM2M)	8.1	105	0.90	3.6	197
INM (INMCM4)	8.6	317	0.65	4.0	698
IPSL (IPSL-CM5A-LR)	7.7	95	0.59	5.5	286
MIROC (MIROC5)	8.3	145	0.76	3.5	285
MOHC (HadGEM2-ES)	6.5	82	0.55	5.3	280
MPIM (MPI-ESM-LR)	7.3	71	0.72	3.9	164
MRI (MRI-CGCM3)	8.5	64	0.66	4.3	150
NCC (NorESM1-M)	8.0	105	0.88	4.0	218
Multimodel mean	7.7	106	0.70	4.3	257
- without INM	7.6	87	0.70	4.3	217
Standard deviation	0.8	71	0.13	0.7	150
- without INM	0.8	26	0.14	0.7	60

2 box EBM
 ϵ free
parameter

Model	C (W y m $^{-2}$ K $^{-1}$)	C_0 (W y m $^{-2}$ K $^{-1}$)	γ (W m $^{-2}$ K $^{-1}$)	τ_f (y)	τ_s (y)
BCC (BCC-CSM1-1)	8.4	56	0.59	4.1	152
CCCMA (CanESM2)	8.0	77	0.54	4.5	139
CNRM (CNRM-CM5.1)	8.3	95	0.51	5.2	266
CSIRO (CSIRO-Mk3-6-0)	8.5	76	0.71	4.2	316
GFDL (GFDL-ESM2M)	8.8	112	0.85	3.6	233
INM (INMCM4)	8.5	271	0.67	4.0	546
IPSL (IPSL-CM5A-LR)	8.1	100	0.57	5.5	327
MIROC (MIROC5)	8.7	158	0.73	3.6	338
MOHC (HadGEM2-ES)	7.5	98	0.49	5.4	457
MPIM (MPI-ESM-LR)	8.5	78	0.62	4.0	220
MRI (MRI-CGCM3)	9.3	68	0.59	4.4	181
NCC (NorESM1-M)	9.7	121	0.76	4.1	328
Multimodel mean	8.5	109	0.64	4.4	300
Standard deviation	0.6	58	0.11	0.7	113

Analytical solution (temperature response T)

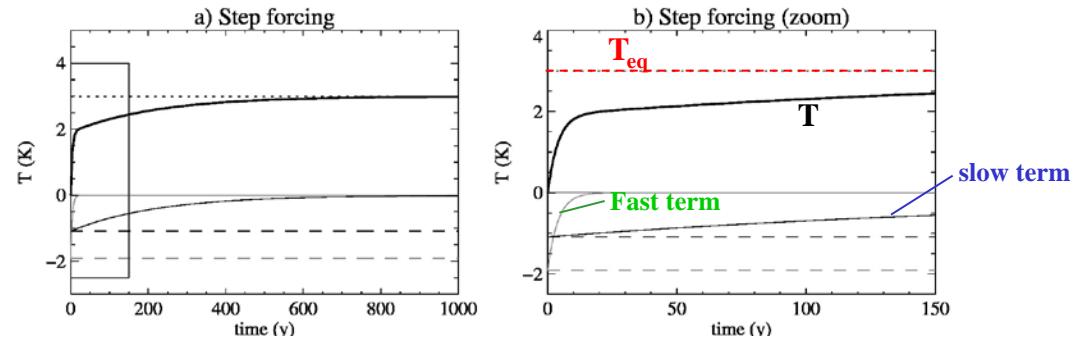
Abrupt forcing (e.g. AB4CO₂)



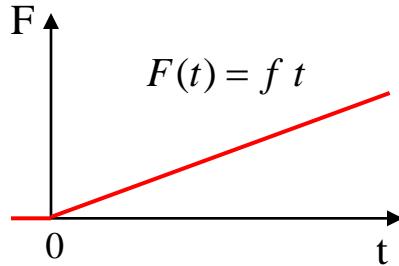
Simulations CMIP5:
Abrupt 4xCO₂

$$T = \frac{F}{\lambda} - \frac{F}{\lambda} a_f e^{-t/\tau_f} - \frac{F}{\lambda} a_s e^{-t/\tau_s}$$

T_{eq} **Fast term** **slow term**

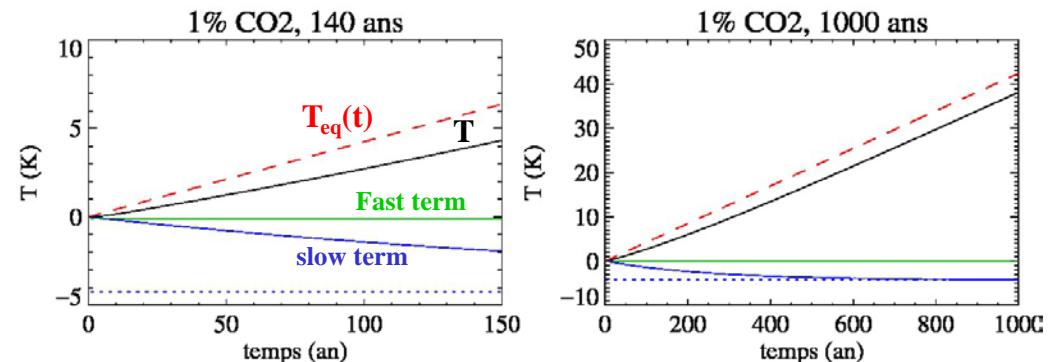


Linear forcing (e.g. 1% yr⁻¹ CO₂)



$$T = \frac{f}{\lambda} t - \frac{f}{\lambda} a_f \tau_f (1 - e^{-t/\tau_f}) - \frac{f}{\lambda} a_s \tau_s (1 - e^{-t/\tau_s})$$

T_{eq(t)} **Fast term** **slow term**

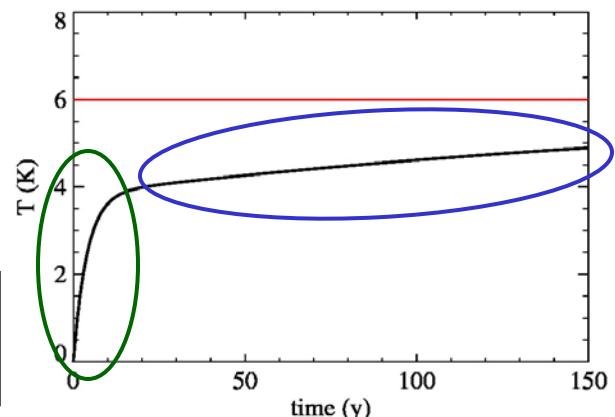


Parameter calibration, method

Analytical solution
(step forcing case)

$$T = \frac{F}{\lambda} - \frac{F}{\lambda} a_f e^{-t/\tau_f} - \frac{F}{\lambda} a_s e^{-t/\tau_s} \quad (1)$$

Calibration of $F_{4\text{CO}_2}$, λ , C , C_0 , γ from AB4CO2



→ Linear fit $N=f(T)$ (Grégory et al, 2004): →

F, λ

$$\rightarrow (1) \xrightarrow{t \gg \tau_f} \log\left(\frac{T_{eq} - T}{T_{eq}}\right) = \frac{-t}{\tau_s} + \log(a_s)$$

Fit over
30-150 yr

a_s, τ_s

→ Analytical relationship

$$a_f + a_s = 1$$

a_f

$$\rightarrow (1) \Rightarrow \tau_f = \frac{-t}{\log\left(\left(\frac{T_{eq} - T}{T_{eq}} - a_s e^{-t/\tau_s}\right) / a_f\right)}$$

Mean of the
first 10 yr

τ_f

$$\rightarrow \text{Analytical relationships} \quad C = \frac{\lambda}{\frac{a_f}{\tau_f} + \frac{a_s}{\tau_s}} \longrightarrow C$$

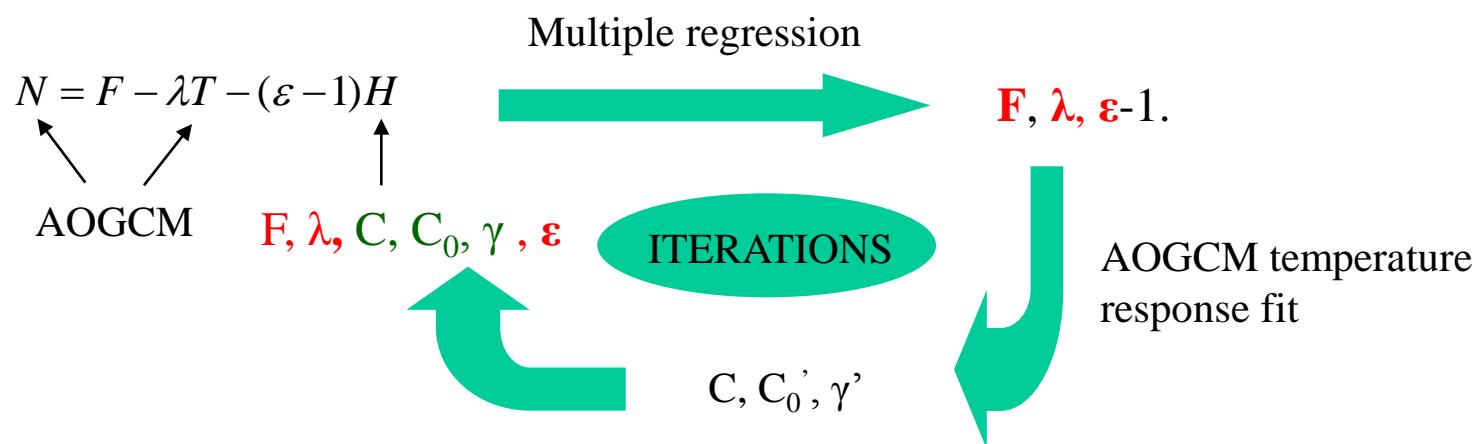
$$C_0 = \lambda(\tau_f a_f + \tau_s a_s) - C \longrightarrow C_0$$

$$\gamma = \frac{C_0}{(\tau_f + \tau_s - \frac{C + C_0}{\lambda})} \longrightarrow \gamma$$

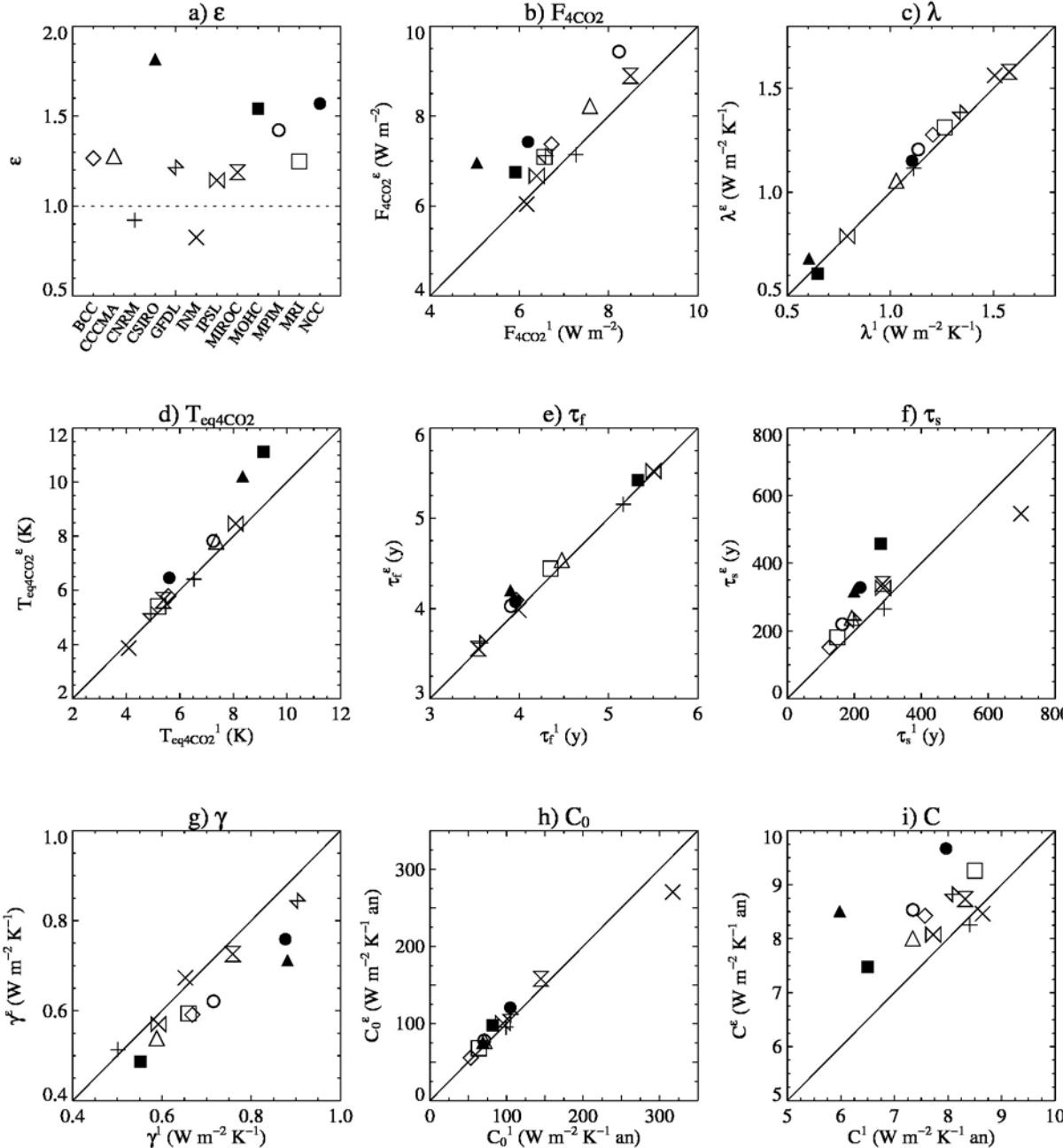
Parameter calibration, method

6 parameters: $F, \lambda, \varepsilon, C, C_0, \gamma$

$$\begin{cases} C \frac{dT}{dt} = F - \lambda T + \gamma'(T - T_0) \\ C_0' \frac{dT_0}{dt} = \gamma'(T - T_0) \end{cases}$$
$$\gamma' = \varepsilon \gamma$$
$$C_0' = \varepsilon C_0$$



Comparison of parameters EBM -1 vs EBM- ε



- X-axis: EBM-1 estimation
- Y-axis: EBM- ε estimation

$\varepsilon > 1:$
 $\rightarrow \begin{cases} F \nearrow, \text{ECS} \nearrow \\ \gamma \downarrow \\ \tau_{\text{slow}} \nearrow \\ C \nearrow \end{cases}$

Correlations ?

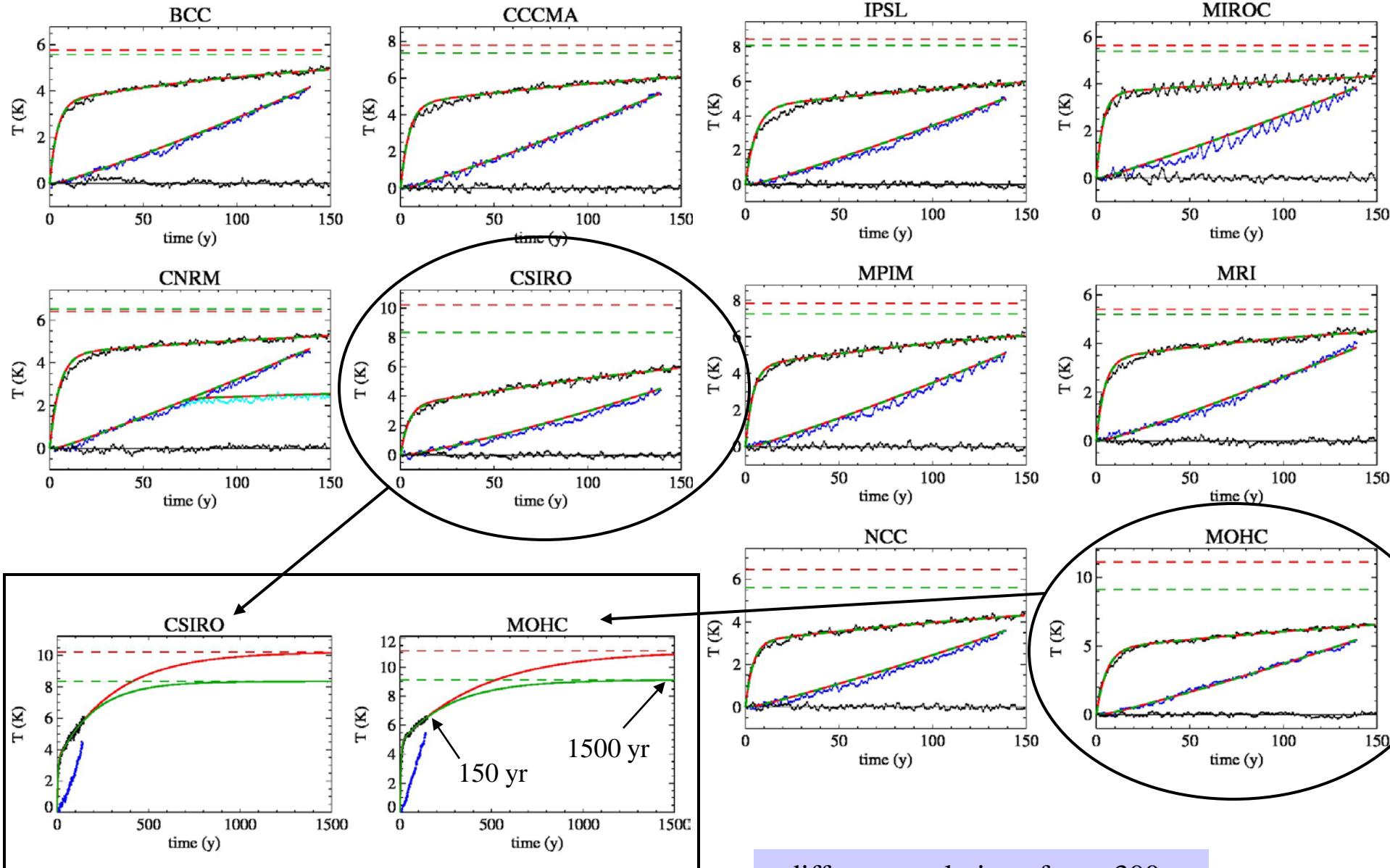
TABLE 3. Intermodel correlations between the equilibrium temperature at 4xCO₂ $T_{4\text{xCO}_2}$, and the physical parameters \mathcal{F} , λ , ε , γ , C_0 , C of the EBM- ε for the 12 CMIP5 AOGCMs.

	$T_{4\text{xCO}_2}$	\mathcal{F}	λ	ε	γ	C_0	C
$T_{4\text{xCO}_2}$	1	0.02	-0.86	0.64	-0.38	-0.45	-0.51
\mathcal{F}		1	0.23	0.18	0.06	-0.28	0.12
λ			1	-0.55	0.42	0.46	0.47
ε				1	0.14	-0.48	0.09
γ					1	0.29	0.62
C_0						1	0.06
C							1

- No corrélation between λ and γ .
Agreement with Gregory and Foster (2002).
Not with Raper et al. (2002)

- light corrélation ε - ECS.

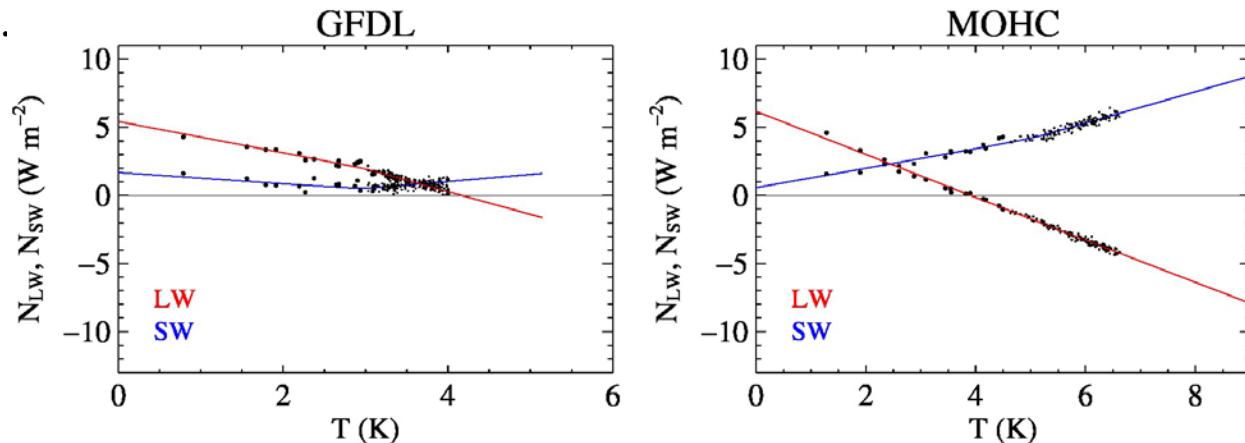
Results, net flux TOA $N=f(T)$



Decomposition of N

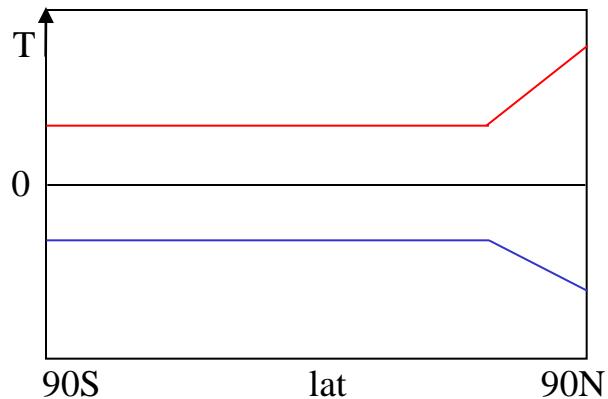
$$N^i = F^i - \lambda^i T - (\lambda^i - \lambda_D^i) \frac{\epsilon}{\lambda} H$$

Multimodel: i =[LW,SW].



Model	\mathcal{F}^{LW} (W m ⁻²)	\mathcal{F}^{SW} (W m ⁻²)	λ^{LW} (Wm ⁻² K ⁻¹)	λ^{SW} (Wm ⁻² K ⁻¹)	λ_D^{LW} (Wm ⁻² K ⁻¹)	λ_D^{SW} (Wm ⁻² K ⁻¹)
BCC	6.4	1.0	1.69	-0.42	1.68	-0.68
CCCMA	6.2	2.0	1.42	-0.37	1.38	-0.57
CNRM	5.1	2.1	1.62	-0.50	1.67	-0.46
CSIRO	7.4	-0.4	1.97	-1.29	1.81	-1.43
GFDL	5.4	1.7	1.37	0.01	1.68	-0.54
INM	6.8	-0.7	2.12	-0.55	2.65	-0.76
IPSL	3.4	3.3	1.92	-1.13	1.89	-1.20
MIROC	6.9	2.0	1.93	-0.35	1.70	-0.37
MOHC	6.2	0.6	1.56	-0.96	1.55	-1.16
MPIM	7.0	2.5	1.67	-0.46	1.50	-0.65
MRI	6.6	0.5	2.24	-0.93	2.16	-1.11
NCC	6.3	1.1	1.82	-0.67	1.67	-0.93
Mean	6.1	1.3	1.78	-0.63	1.78	-0.82
STDV	1.1	1.2	0.27	0.37	0.34	0.34

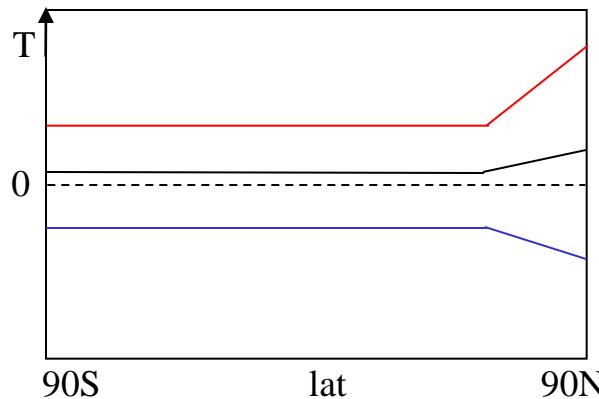
Flux de chaleur accumulé dans l'océan: introduction d'une efficacité de forçage ϵ



$$T = T_{\text{eq}} + T_U$$

Schéma: moyenne zonale de T

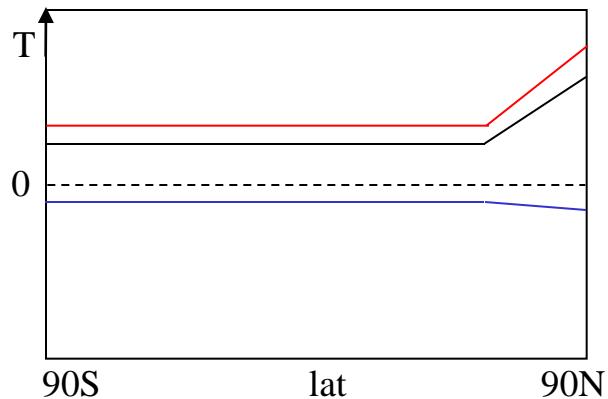
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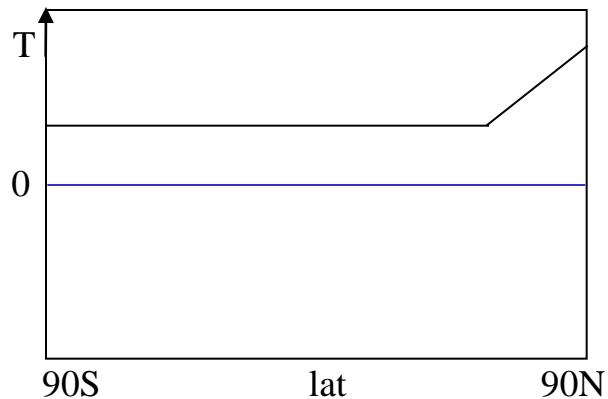
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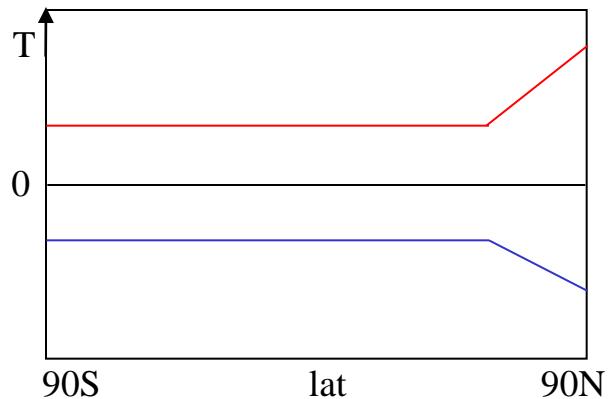
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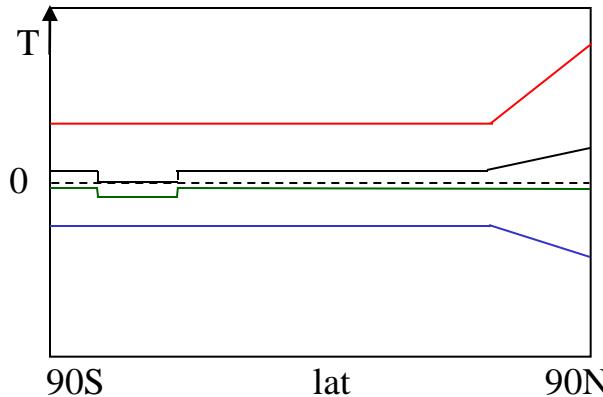
Flux de chaleur accumulé dans l'océan: introduction d'une efficacité de forçage ϵ



$$T = T_{\text{eq}} + T_U + T_D$$

Schéma: moyenne zonale de T

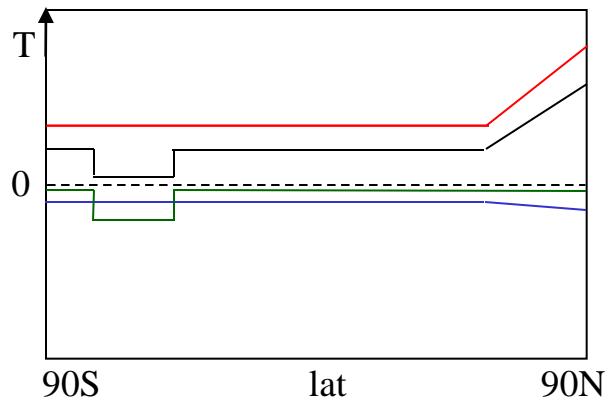
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$$T = T_{\text{eq}} + T_U + T_D$$

Schéma: moyenne zonale de T

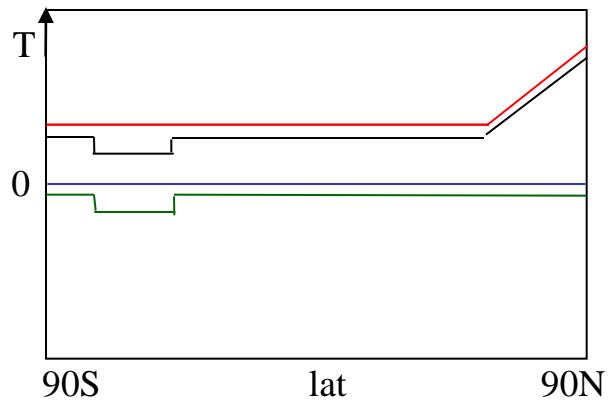
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$$T = T_{\text{eq}} + T_U + T_D$$

Schéma: moyenne zonale de T

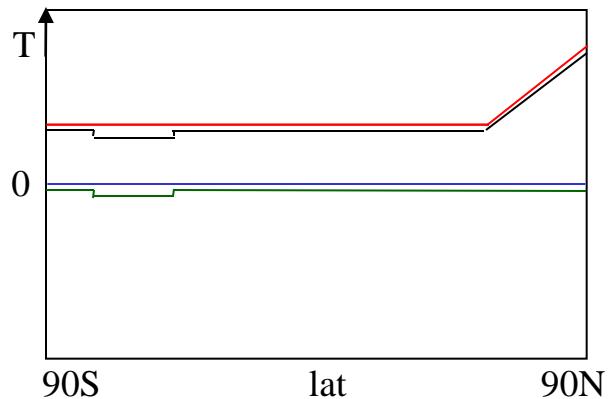
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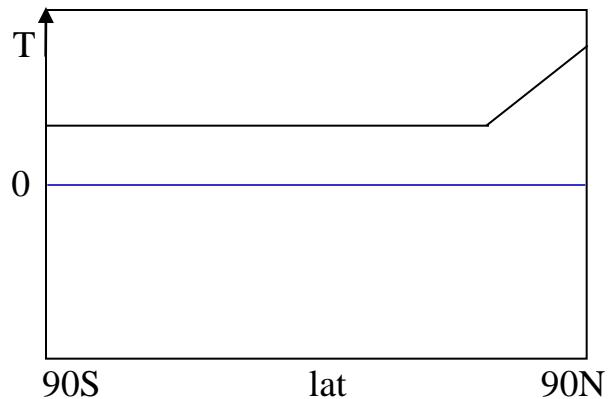
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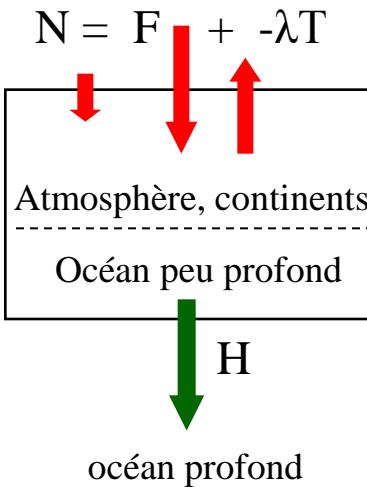
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$$T = T_{eq} + T_U + T_D$$

Schéma: moyenne zonale de T

1-box EBM



Energie accumulée
océan peu profond
(+continents
+atmosphère)

Déséquilibre
radiatif N

Energie accumulée dans l'océan
profond

$$C \frac{dT}{dt} = F - \lambda T + H$$

avec $H = \kappa \cdot T$

(Gregory and Mitchell, 1997)