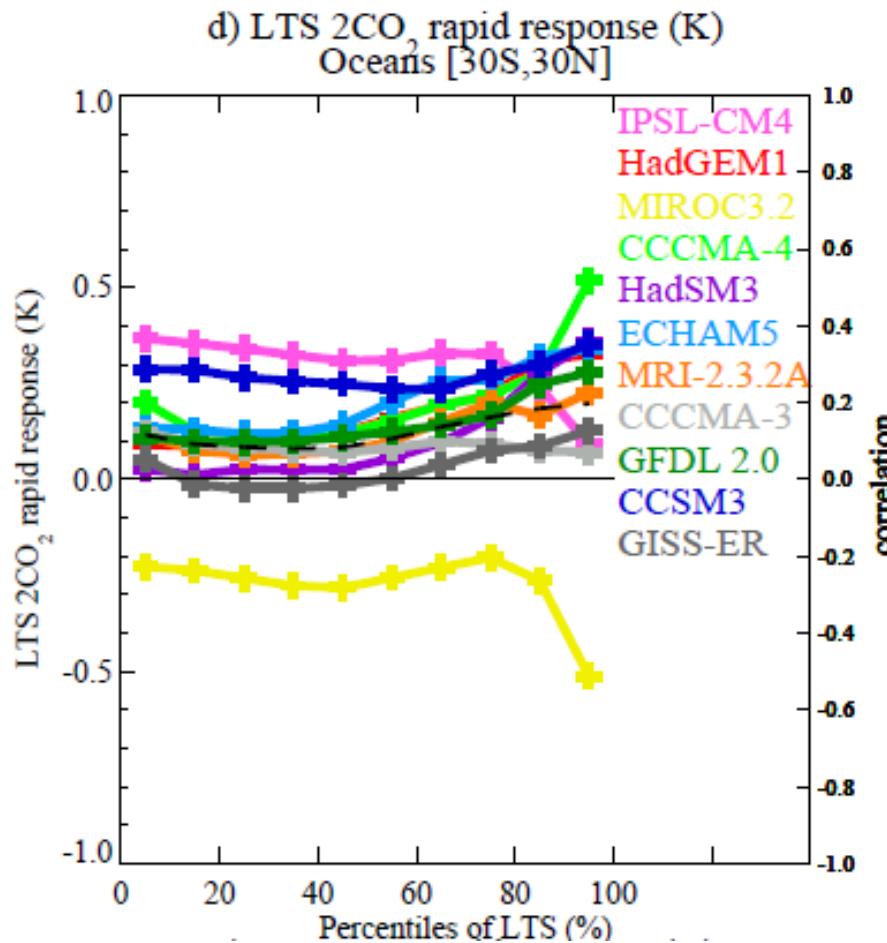


Importance of instantaneous radiative forcing to tropospheric adjustment

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Tropospheric adjustment of Lower Tropospheric Stability (LTS) to CO₂ doubling (CFMIP1/CMIP3 slab ocean experiments)



Webb et al. submitted to Climate Dynamics

Negative response of MIROC3.2 leads to large inter-model difference.²

Understanding the difference in LTS adjustment

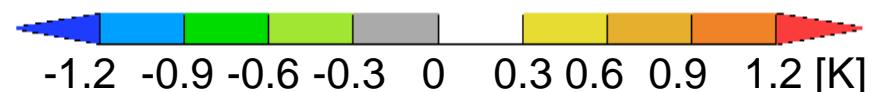
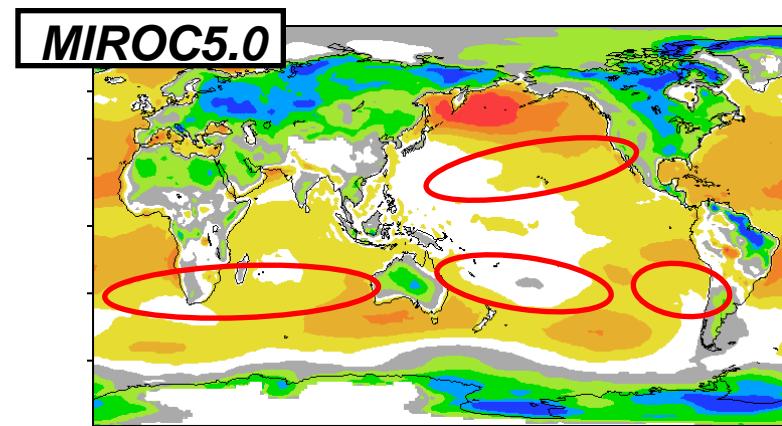
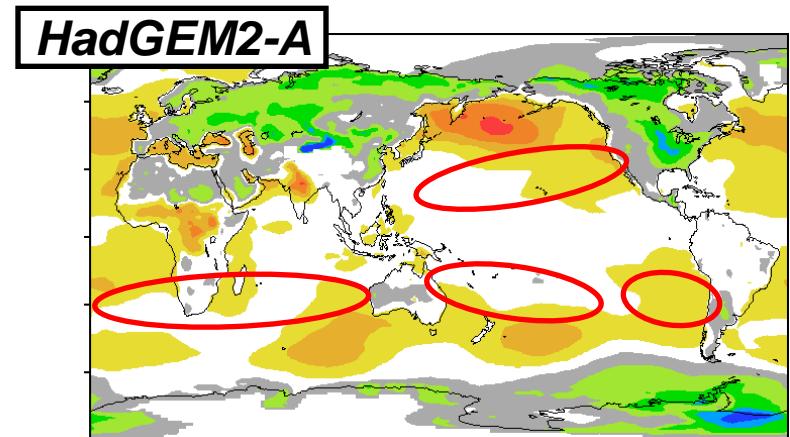
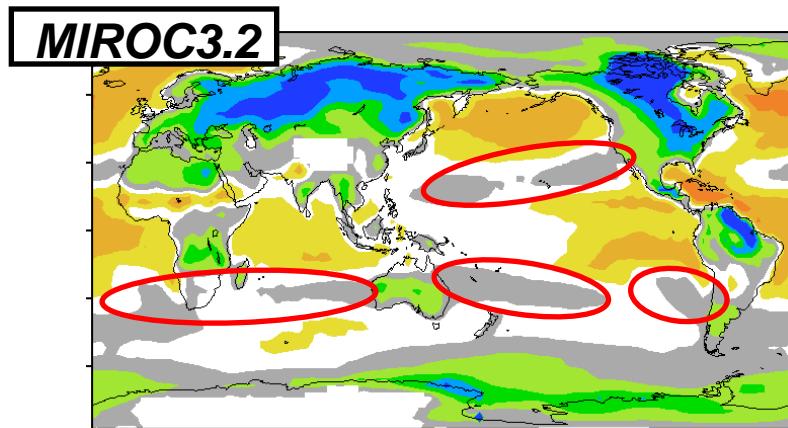
Additional experiments with 3 GCMs

Experimental design

- Trop. adjustment evaluated by “**4xCO₂ AMIP**” minus “**AMIP**”, 1979-2008.
; difference in 30 years average between the two experiments. (Hansen et al. 2005)
 - The above approach chosen to confirm the result of regression method.
(Gregory et al. 2004)
 - Using MIROC3.2, four members of initial value ensemble runs obtained for significance test (95%).
 - For HadGEM-2 and MIROC5.0, one member analyzed for each.

Tropospheric adjustment of LTS estimated from AMIP-type runs

LTS response ($4\times\text{CO}_2$ AMIP minus AMIP, Annual mean)

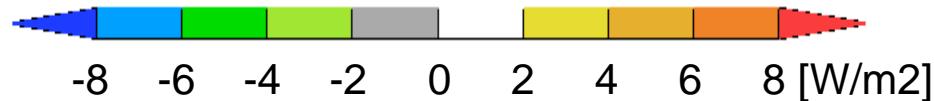
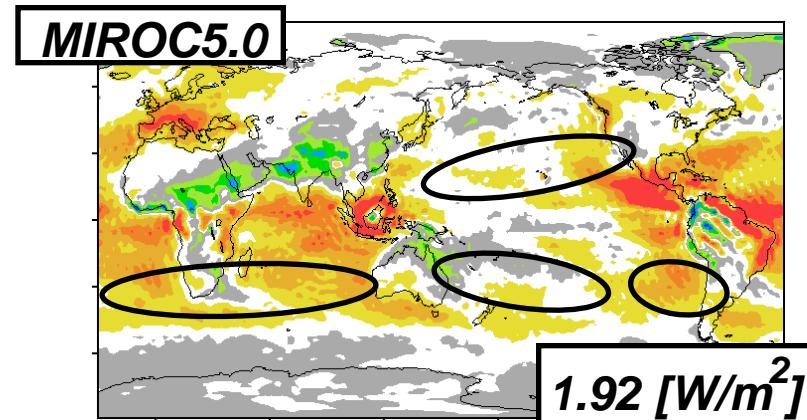
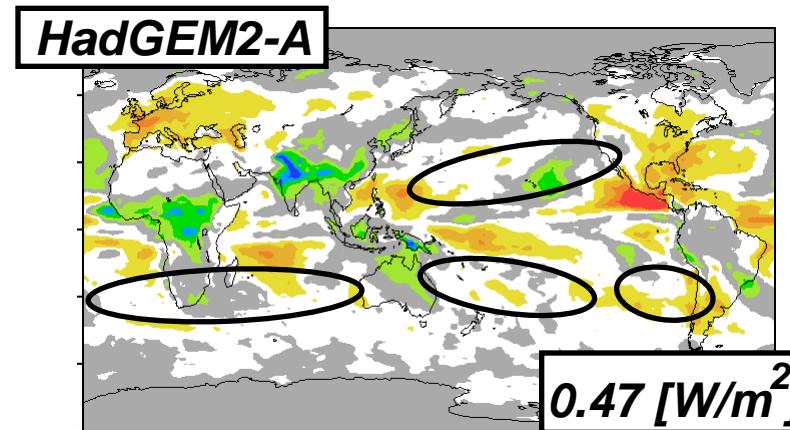
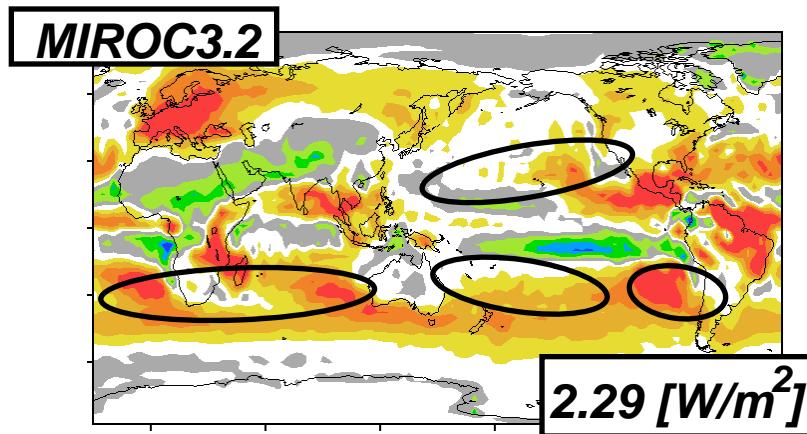


Negative adjustment found in MIROC3.2 (over subtropical ocean).



Tropospheric adjustment of Cloud Radiative Effect (CRE) from AMIP-type runs

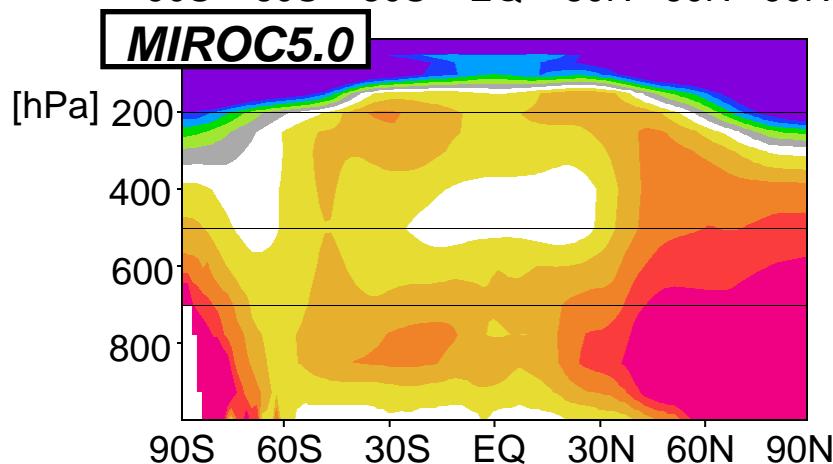
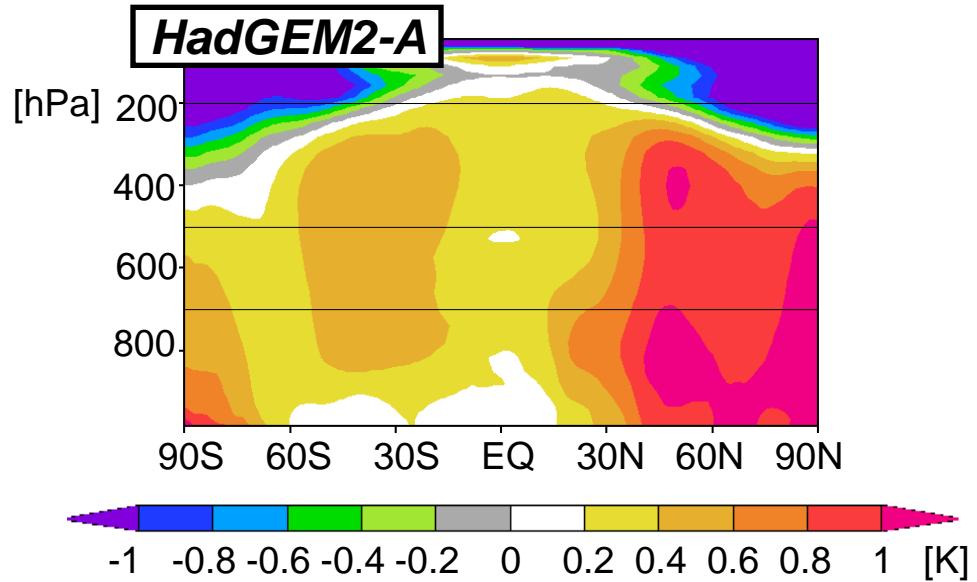
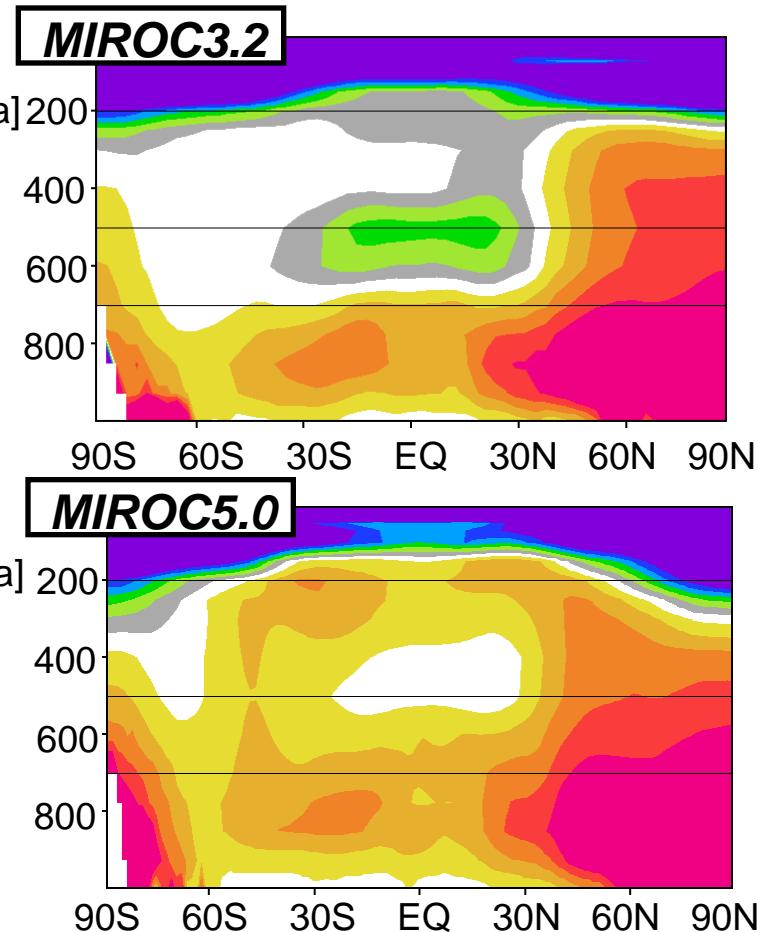
Shortwave CRE Response ($4\times\text{CO}_2 \text{ AMIP} \text{ minus AMIP}$, Annual mean)



Positive adjustment in MIROC3.2, consistent with the LTS adjustment.

Tropospheric adjustment of temperature profile from AMIP-type runs

Temperature response ($4\times\text{CO}_2$ AMIP minus AMIP, Zonal Annual mean)



What causes the negative adjustment in mid-troposphere in MIROC3.2 ?

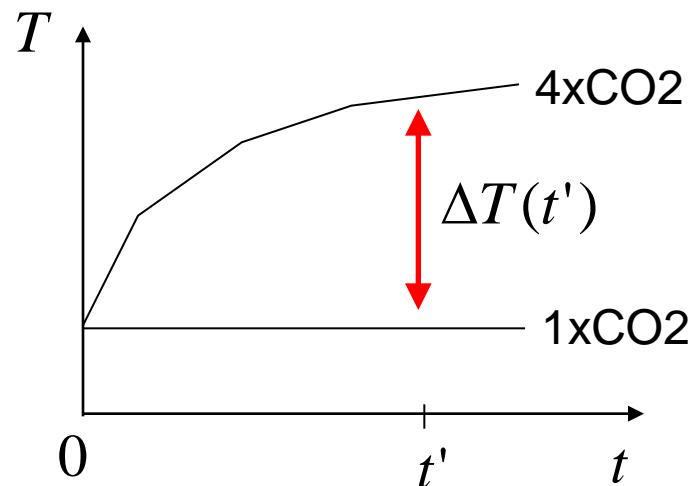
Analysis of temperature tendency terms

$$\frac{\partial T}{\partial t} = \text{Convective Cloud} + \text{Non-Convective Cloud} + \text{Longwave Heating} + \text{Shortwave Heating} + \text{Diffusion} + \text{Dynamics} + \text{Residual}$$

$$\equiv \sum_{i=1}^7 \dot{T}_i$$

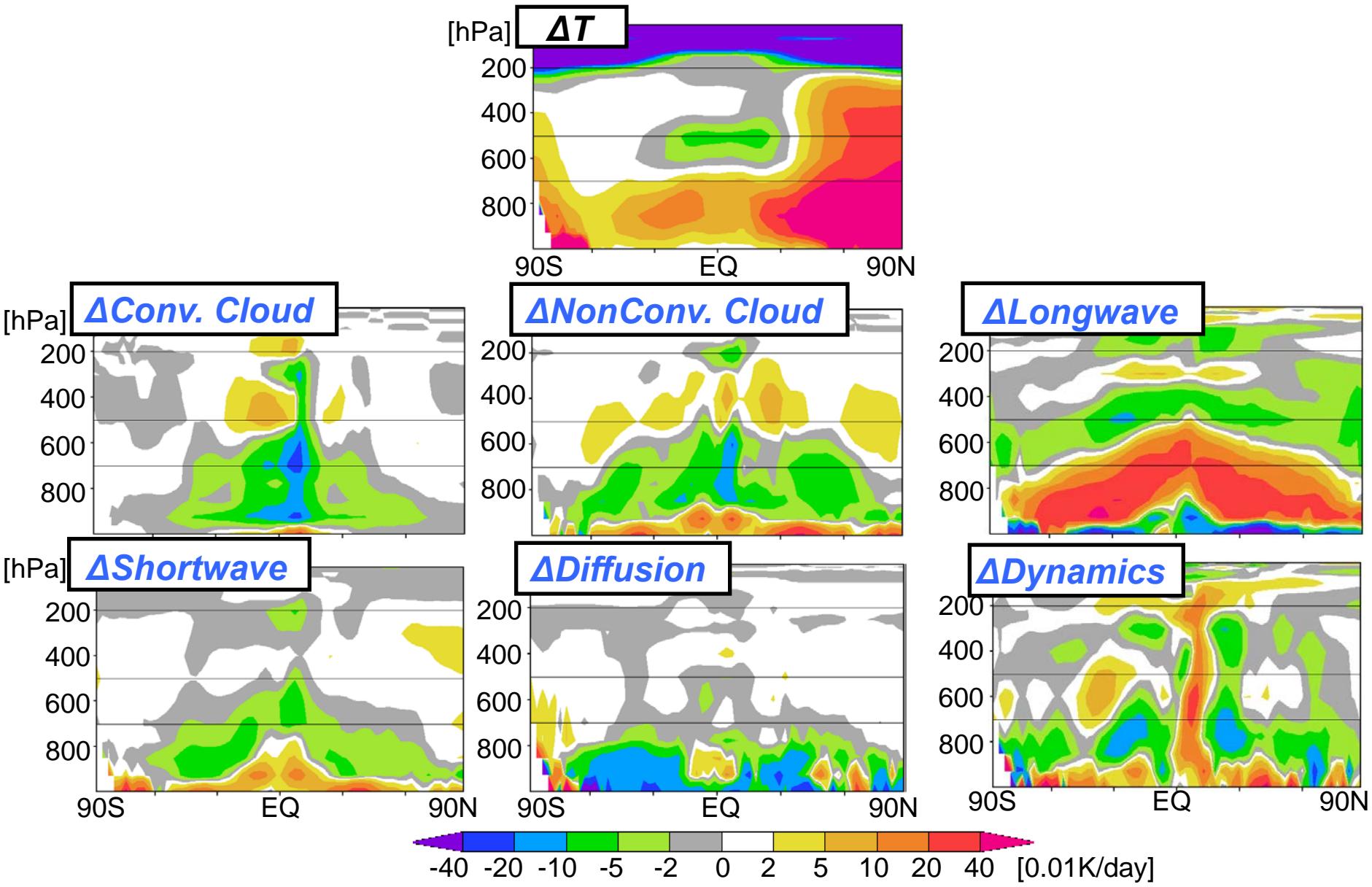
Integrating from $t=0$ to $t=t'$, and taking the difference between $4x\text{CO}_2$ and $1x\text{CO}_2$ by Δ ,

$$\Delta T(t') = \left(\sum_{i=1}^7 \bar{\Delta \dot{T}_i} \right) \times t'$$



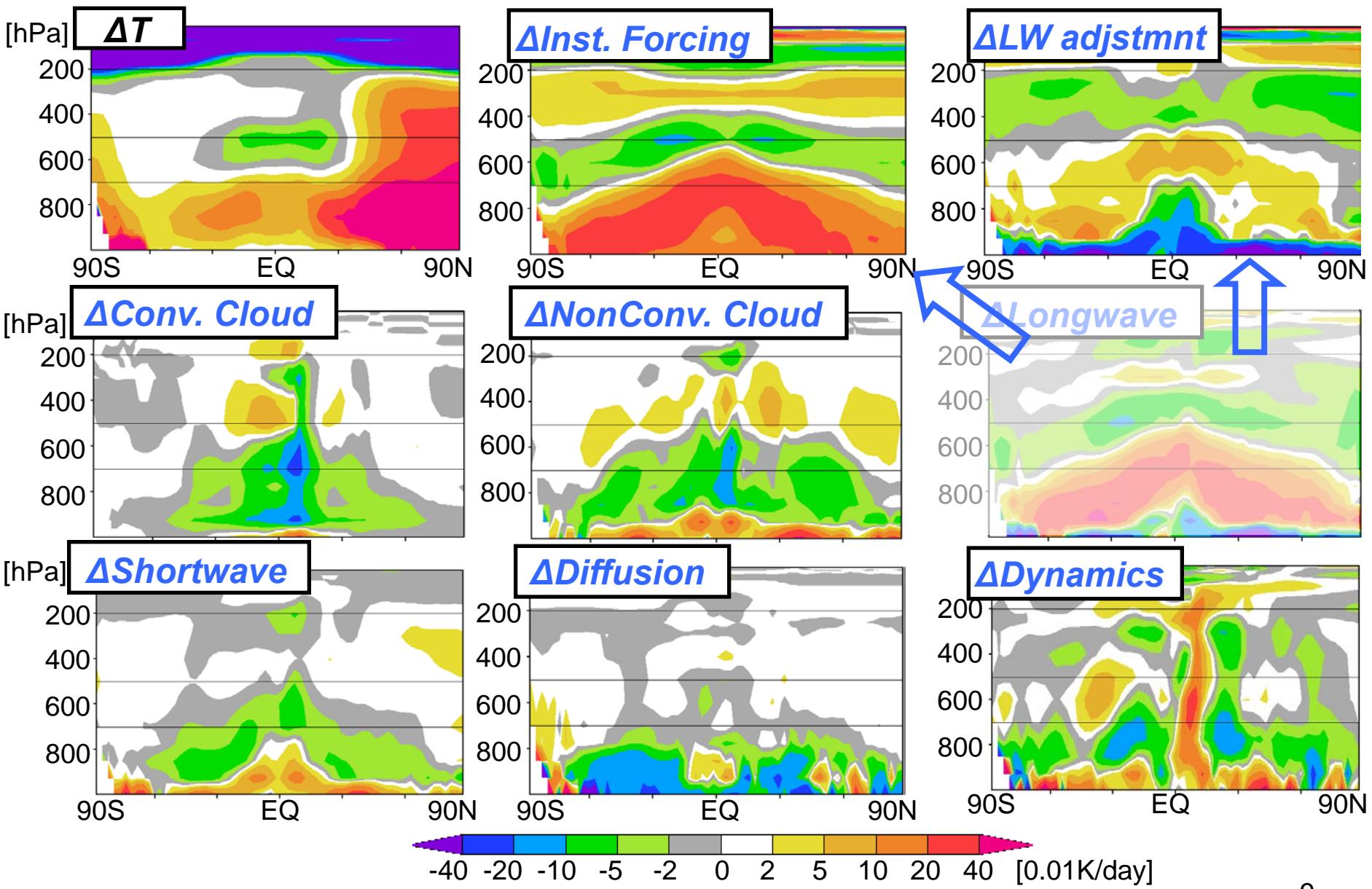
Consistency checked between T response, $\Delta T(t')$, and tendency terms response, $\bar{\Delta \dot{T}_i}$.

Tropospheric adjustment of temperature tendency terms in MIROC3.2



Contribution from multiple terms to the negative ΔT in mid-troposphere.

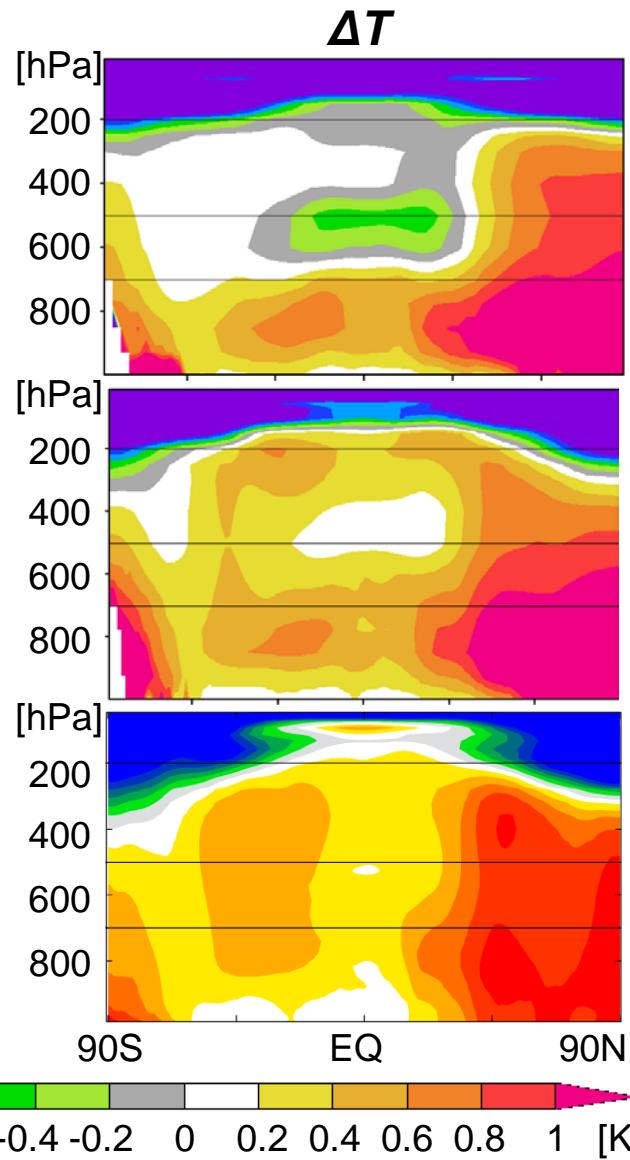
Tropospheric adjustment of temperature tendency terms in MIROC3.2



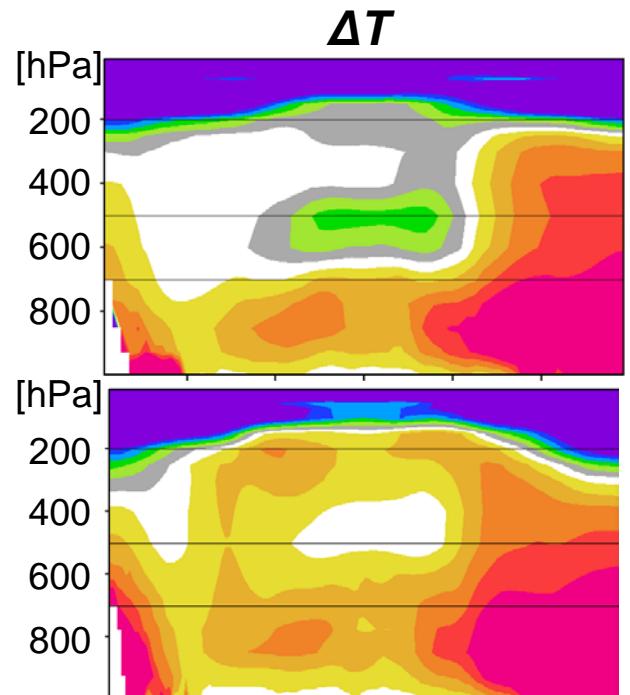
Contribution of $\Delta \text{Inst. Forcing}$ appears important to the negative ΔT .

Inter-model difference of $\Delta\text{Inst. Forcing}$; is it consistent with ΔT ?

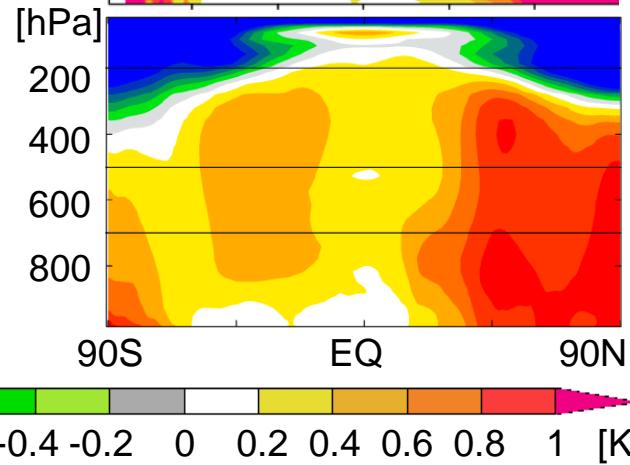
MIROC3.2



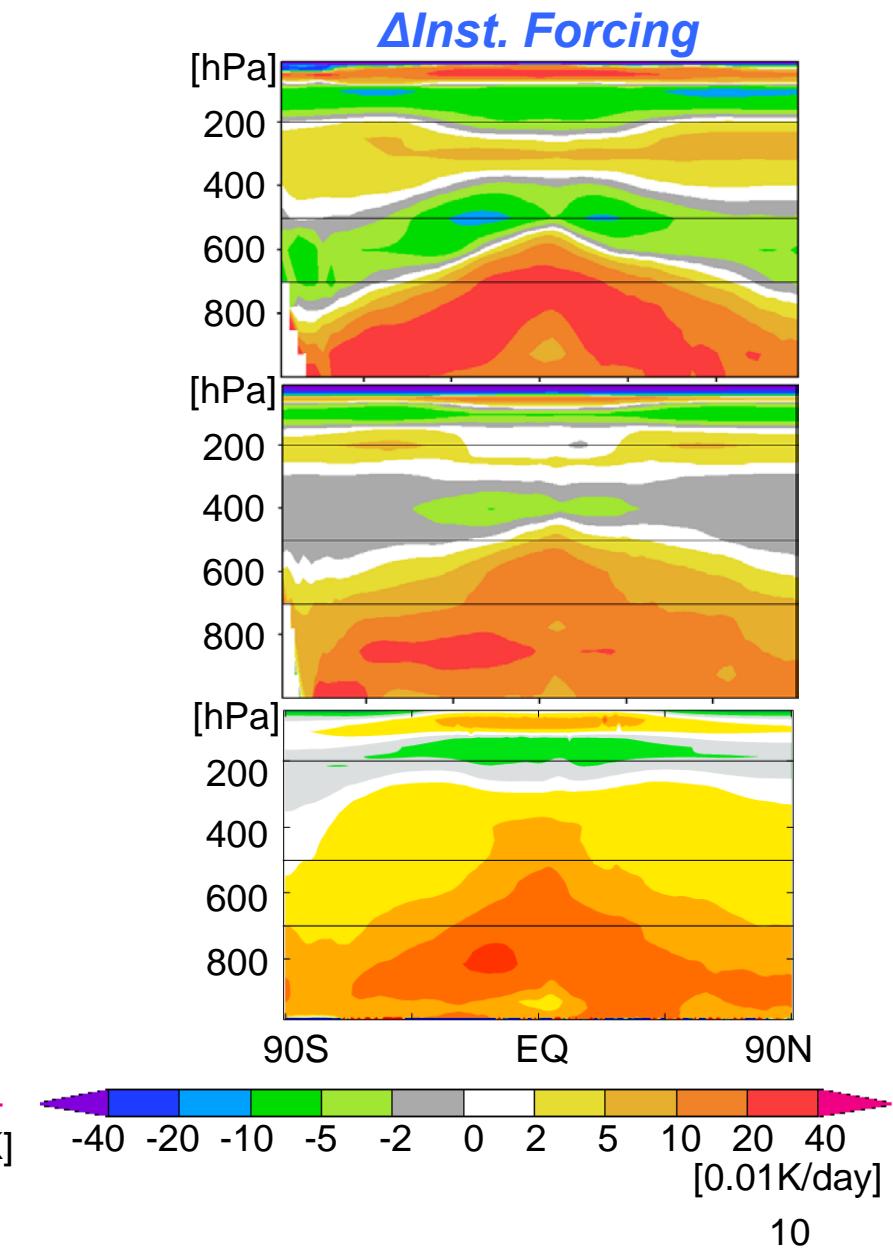
MIROC5.0



HadGEM2-A



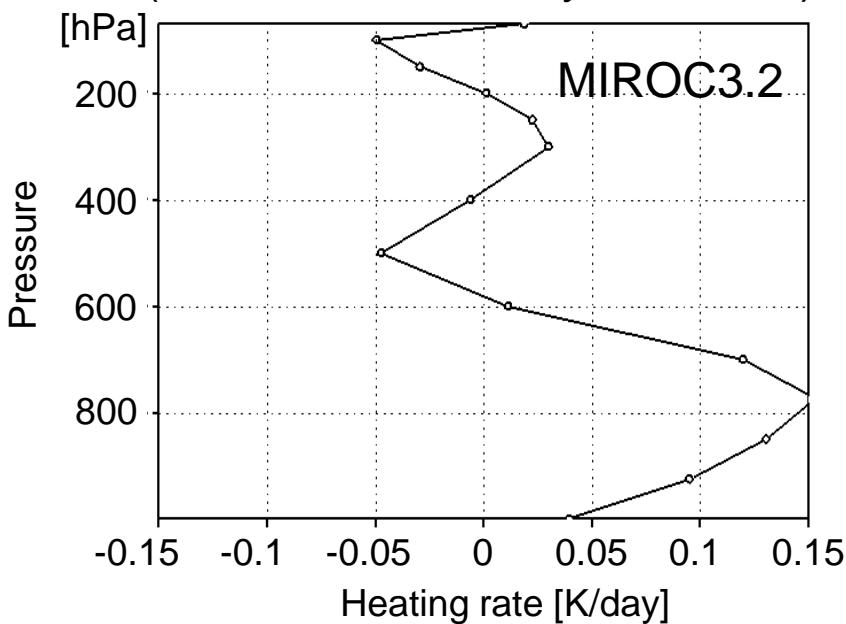
$\Delta\text{Inst. Forcing}$



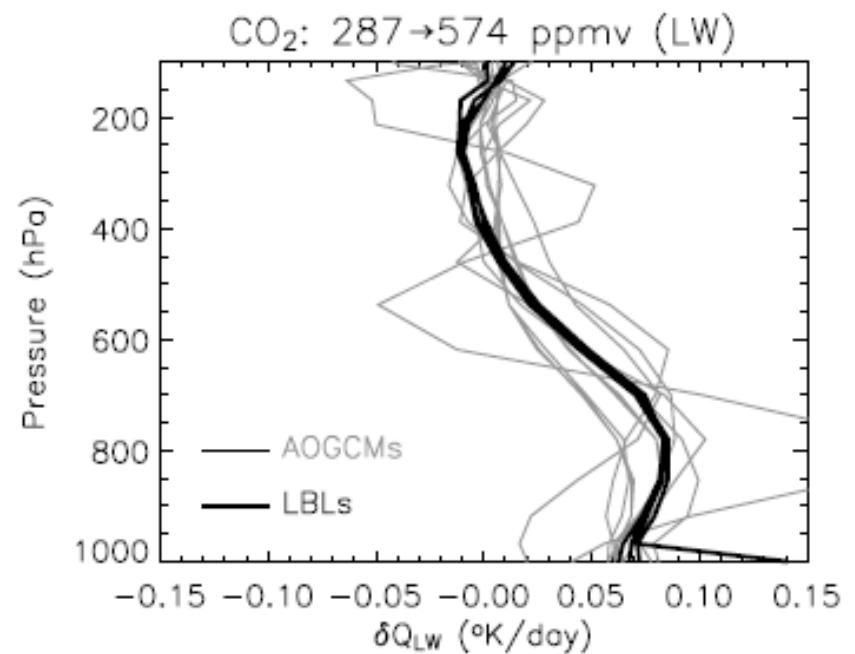
Inter-model difference of $\Delta\text{Inst. Forcing}$ appears consistent with ΔT .

Evaluation of $\Delta\text{Inst. Forcing}$ with Line-By-Line (LBL) calculations.

Instantaneous 4xCO₂ Heating rate x 0.5
(JJA mean, clear sky, 30N-60N)



Longwave heating rate perturbation



Adapted from Collins et al. (2006), JGR, Fig.10

Radiative Transfer Model Intercomparison Project (RTMIP),
AR4 models including MIROC3, compared with LBLs.

The negative $\Delta\text{Inst. Forcing}$ in mid-troposphere may be questionable.

Tropospheric adjustment of Lower Tropospheric Stability (LTS) to 2xCO₂ (CFMIP1/CMIP3 slab ocean experiments)

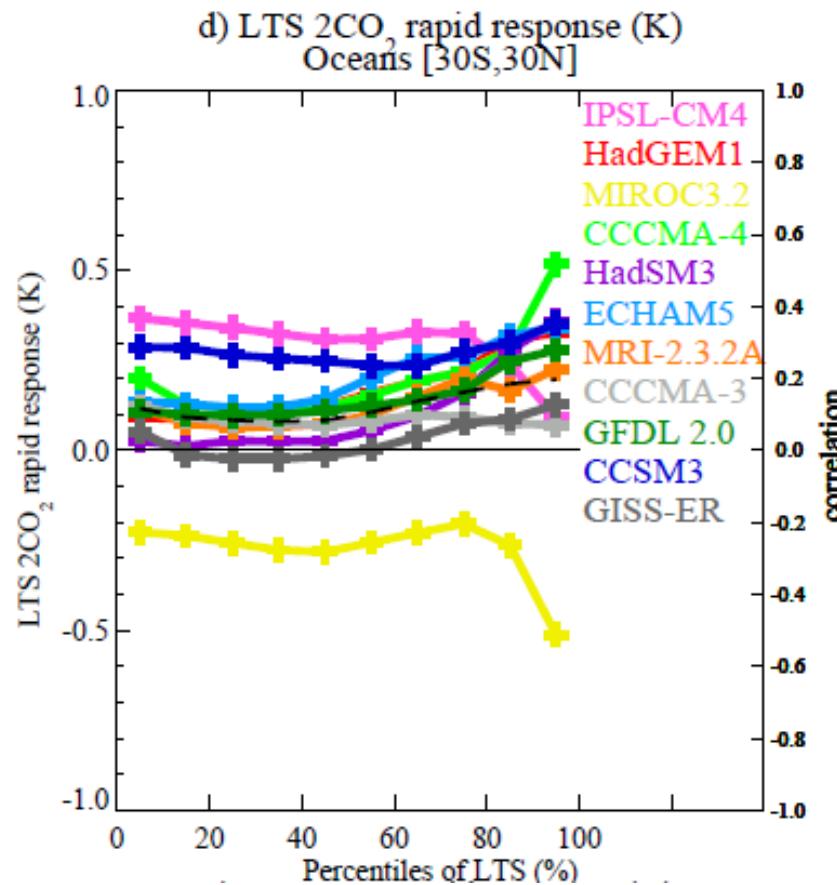


Fig. 9 Lower Tropospheric Stability (LTS) composites of net, shortwave and longwave CRE components of forcing over the low latitude oceans (30N/S) in the AR4 ensemble, and rapid responses of LTS, Estimated Inversion Strength (EIS) and pressure velocity at 500 hPa.

Webb et al. submitted to Climate Dynamics

Negative response of MIROC3.2 leads to large inter-model difference.

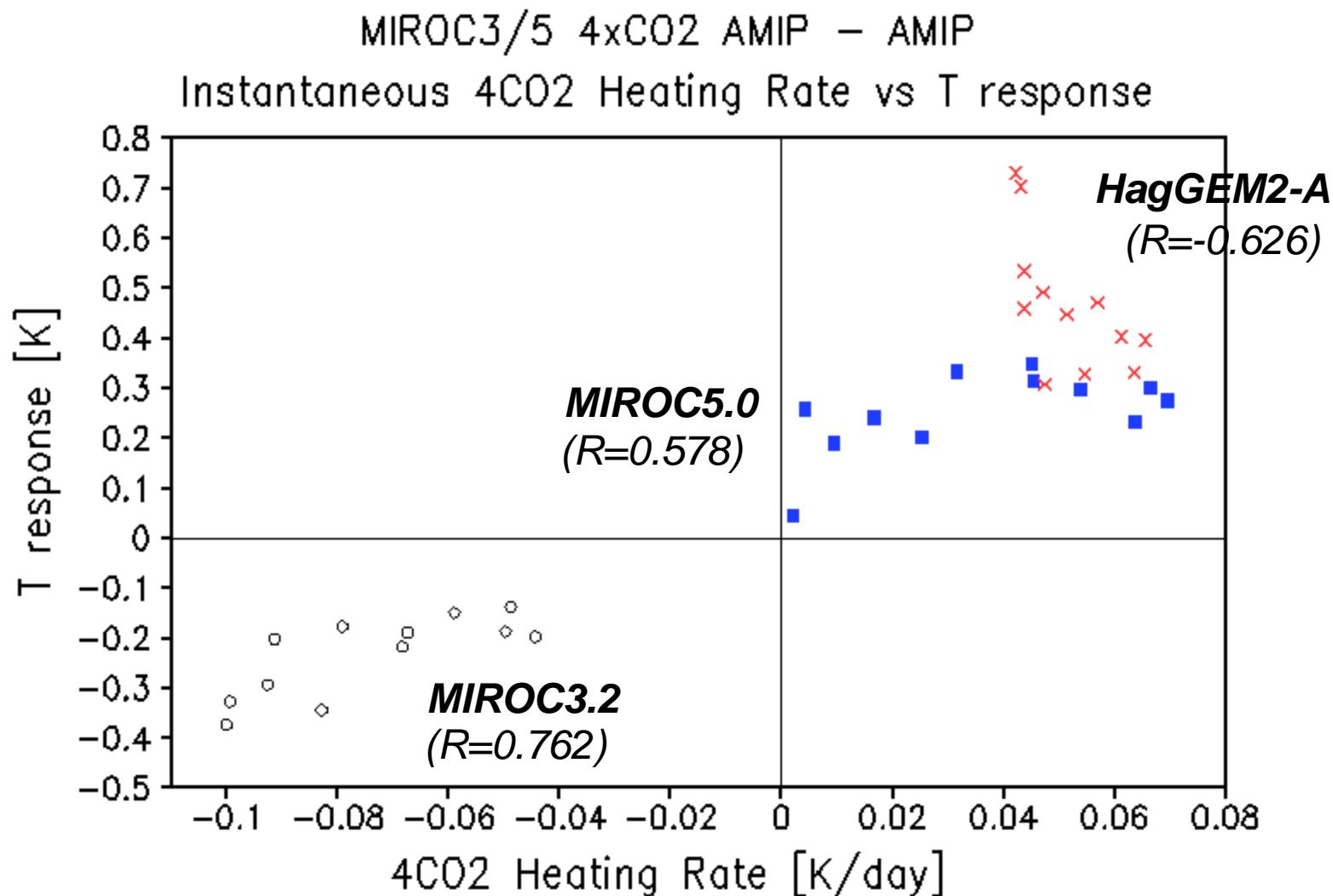
Summary

An example of model inter-comparison study on LTS adjustment presented.

- LTS reduction in MIROC3.2 reflects cooling in mid-troposphere , which can be explained by multiple factors, especially instantaneous radiative forcing.
- Comparison with LBL calculation suggests that the negative forcing in mid-troposphere is questionable.
- Inter-model difference in LTS adjustment is expected to decrease with the above constraint (at least in CMIP3).



Seasonal variation of $\Delta\text{Inst. Forcing}$; is it related to ΔT ?



Seasonal variation of $\Delta\text{Inst. Forcing}$ correlated with ΔT only in MIROCs.

Forcing, feedback and tropospheric adjustment

$$N = Q - Y \cdot \Delta Ts$$

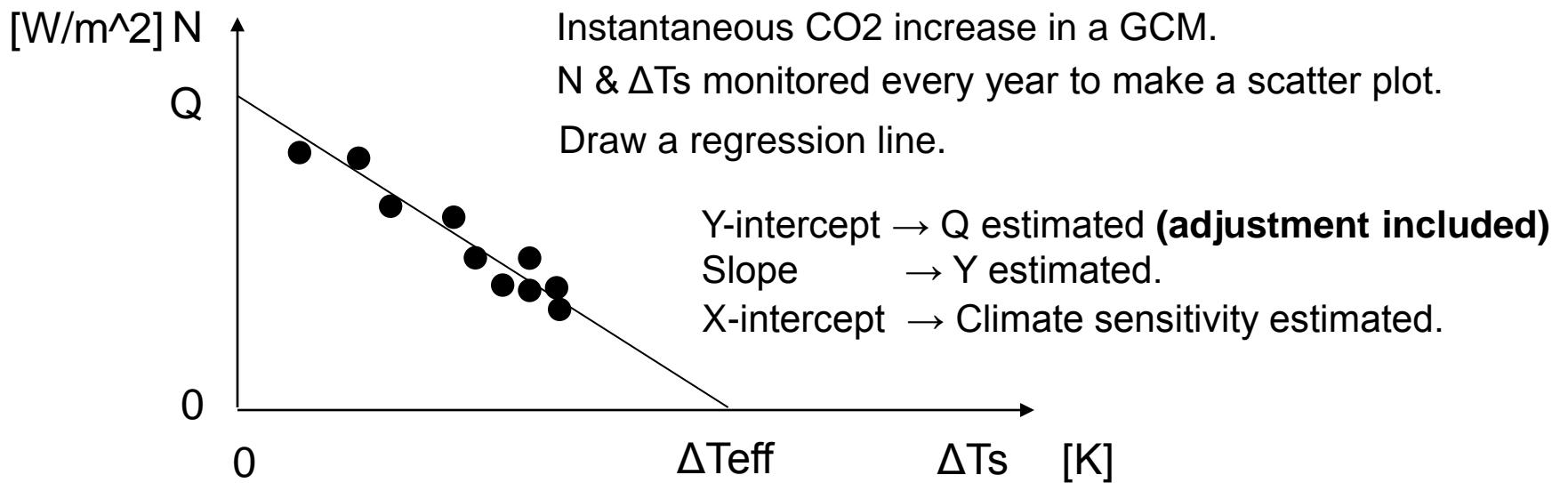
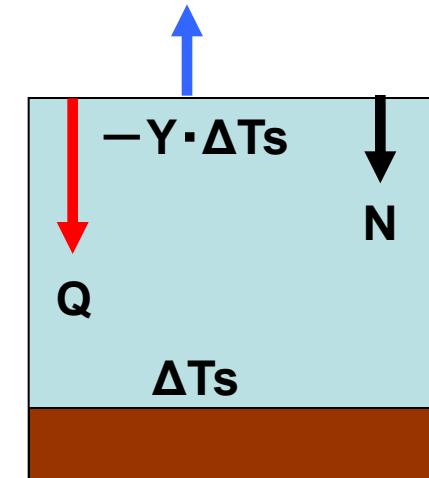
Gregory et al. (2004)

N: radiative imbalance at TOA (positive down, W/m²)

Q: radiative forcing

Y: climate feedback (W/m²/K)

ΔTs : global annual mean surface T change



Lower Tropospheric static Stability (LTS)

$\Delta\theta = \theta(p=700\text{mb}) - \theta(p=\text{sea level pressure}, T=\text{surface air temperature})$

Klein and Hartmann (1993)

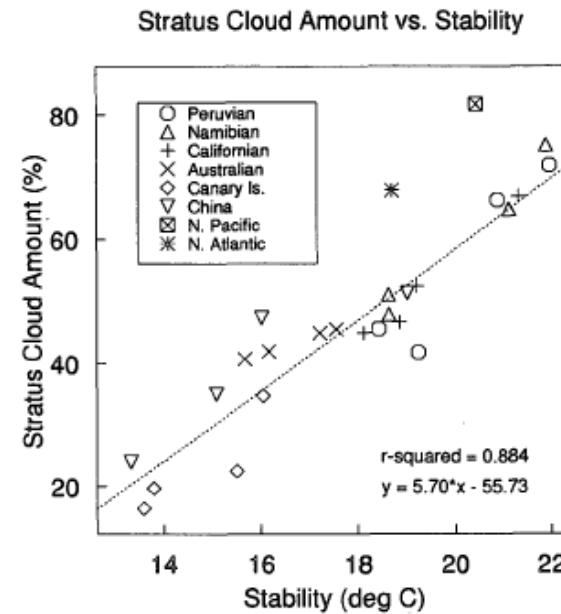
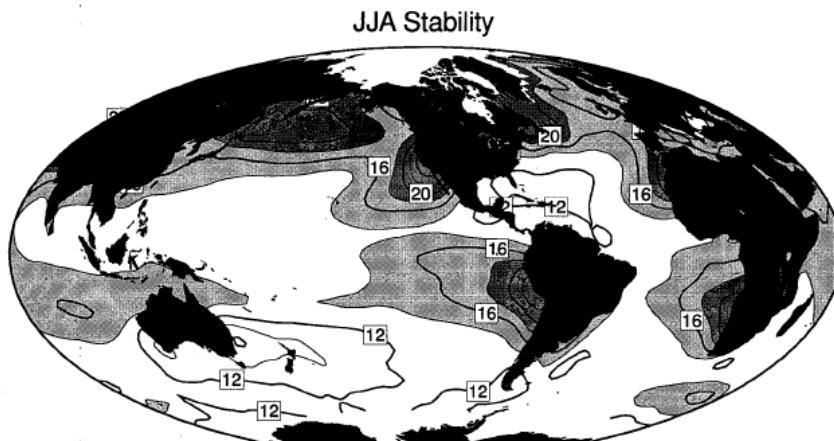
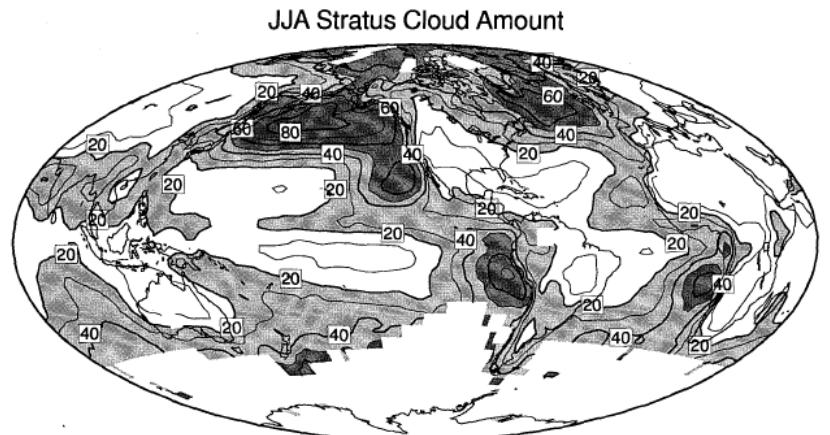
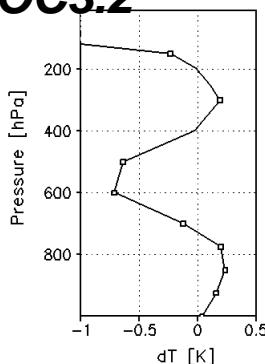


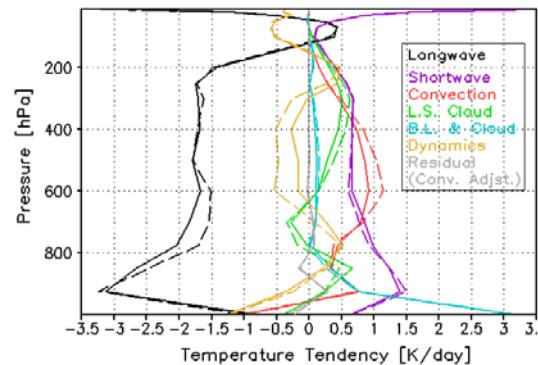
FIG. 13. Scatterplot of seasonally averaged stratocumulus cloud amount with seasonally averaged lower-tropospheric stability for the five subtropical oceanic regions and the Chinese stratus region. In addition, the June, July, and August seasonally averaged quantities are plotted for the North Pacific and North Atlantic but are not included in the regression.



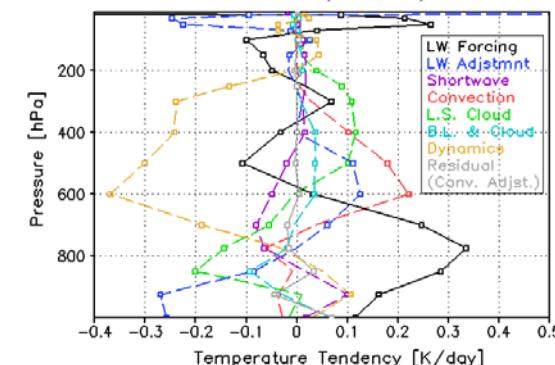
(c) MIROC3 4xCO₂ AMIP – AMIP
Temperature (152E,18N)



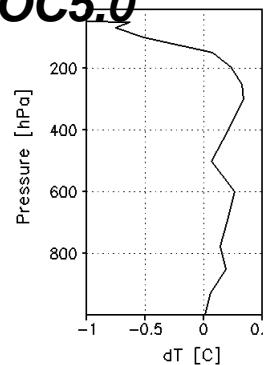
(b) MIROC3 AMIP & 4xCO₂ AMIP
T Tendencies (152E,18N)



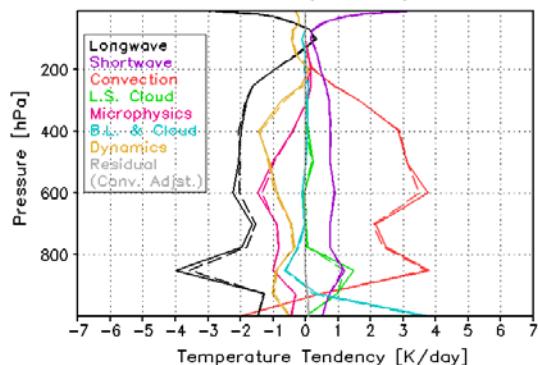
(d) MIROC3 4xCO₂ AMIP – AMIP
T Tendencies (152E,18N)



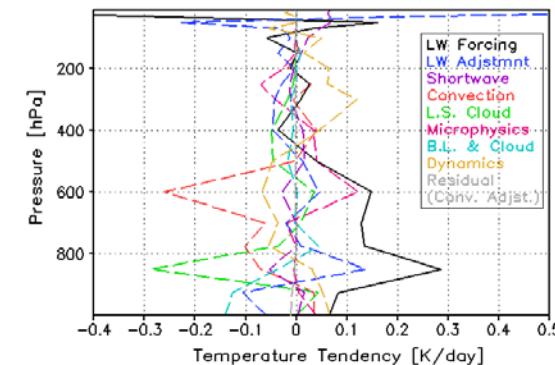
(c) MIROC5 4xCO₂ AMIP – AMIP
Temperature (152E,18N)



(b) MIROC5 AMIP & 4xCO₂ AMIP
T Tendencies (152E,18N)



(d) MIROC5 4xCO₂ AMIP – AMIP
T Tendencies (152E,18N)



MIROC5.0

HadGEM2